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## **Airborne and land-based controlled-source electromagnetic surveying in challenging electromagnetic environments – application to geothermal exploration in a volcanic island**

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### **Summary**

Exploring for underground resources using land-based electromagnetic methods can be very challenging due to the presence of strong human-generated and “geological” noise. In such context, some passive EM techniques like the Magneto-Telluric method may not be applicable at all and a dedicated toolbox of EM techniques capable of dealing with these issues is required. We focus here on the challenges encountered while exploring for geothermal resources in volcanic islands, namely a highly heterogeneous near-surface creating strong static-shifts, the presence of a sea/land transition zone and a high degree of urbanization over the area of interest. We will show that a combined airborne and land-based controlled-source EM approach provide a robust approach and illustrate these aspects with an EM dataset acquired over the Martinique Island, in the French Lesser Antilles. We believe this work sets the foundations for more robust geophysical exploration workflows in such challenging EM environments.

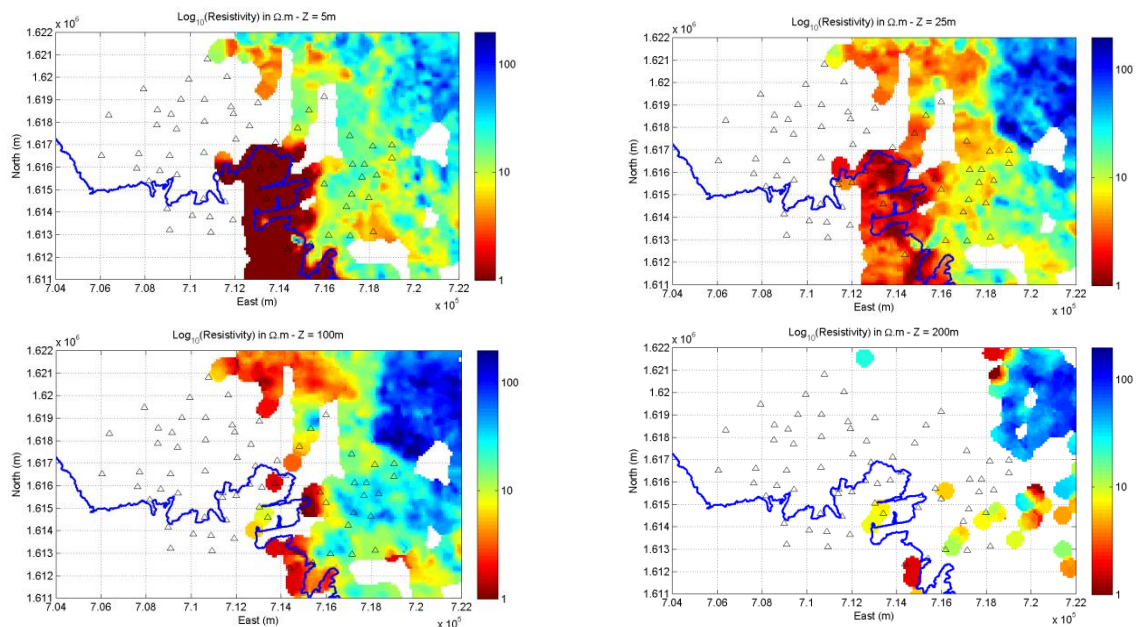


## Airborne EM survey results

A heliborne EM survey covering the whole Martinique has been carried out using the SkyTEM system, which operates in a dual transmitter mode. The survey was flown primarily in the N-S direction with a 400 m framework, locally reduced to 200 to 100 m spacing. The spacing between each EM sounding along the flight lines is approximately 30 m. The nominal height of the loop was on average about 64 m above ground level, due to the variable topography of the island.

Because of the presence of many urbanized areas and strong power lines that drastically deteriorate EM responses, the area of interest was only partially covered and particular attention was paid to properly remove noise from the data at the processing stage. The airborne EM soundings are available every 10m along flight lines. Data were then inverted using the Spatially Constrained Inversion (Viezzoli et al., 2008). Each EM sounding was inverted to obtain a 1-D earth model constrained vertically and laterally by nearby soundings.

Figure 2 shows resistivity maps recovered at 5m, 25m, 100m and 200m depth over the area of interest. It clearly delineates a shallow conductive anomaly ( $< 1 \Omega.m$ ,  $< 25m$ ) around the bay area representative of the seawater intrusion zone. On the other hand, the eastern part of the survey is highly resistive at all depth ( $> 100 \Omega.m$ ) highlighting the presence of thick andesitic lavas.



**Figure 2:** Resistivity maps recovered at 25m, 100m, 250m and 1000m depth from the 1D modelling/inversion of the airborne EM and CSEM datasets. Black triangles represent the CSEM recording stations.

## Land/shallow marine CSEM survey results

### a) CSEM data acquisition and processing

The CSEM survey was acquired using a double LEMAM injection setup for the transmitter (LEMAM standing for Long Electrode Mise-à-la Masse). It uses the metallic casings of boreholes as long electrodes for current injection. Here, two deep exploratory boreholes ( $\sim 400m$ ) connected with a 4km long electric cable were used. Due to the good electrical coupling of the casings with the earth, the current injected was about 30A. To adequately characterize the subsurface, a broad band of CSEM frequencies (fundamental frequencies from 0.125 Hz up to 512 Hz increasing with a factor 4) was acquired with a minimum of 150 cycles at low-frequencies to ensure proper stacking of any random noise. The waveform was a square wave to ensure powerful harmonics to be measured as well. A total



number of 100 CSEM recording stations were deployed on the Lamentin bay area (figure 1), 80 onshore and 20 offshore. They were Metronix ADU06 or ADU07 acquisition systems, MFS07 or MFS06 magnetic coils and 2 orthogonal electric dipoles from 20 to 100m long.

Raw data have been processed in the frequency domain using a proprietary code. After automatic and manual quality control of the processed data, 20% of the data was rejected due to high levels of noise, mainly due to the weak CSEM signal at stations at large distances from the transmitter ( $> 6\text{km}$ ). The remaining 80% of the data were considered of good quality and subsequently inverted. In total, 42 frequencies (= fundamental frequencies + frequencies up to the 13th harmonics  $\times 7$  transmitter frequencies) have been processed to provide a broadband signal and allow an easy removal of noisy frequencies bands (e.g. industrial noise at 50Hz and harmonics). At this stage, we only worked with the amplitude and phase of the maximum axis of the polarization ellipses of the horizontal electric field as their spectra are a lot smoother and hence easier to quality control than individual horizontal components (which can sometimes cross zero and hence be confused with noise). It surely comes at the expense of losing some information, especially the “3D effects”, but we consider these as second order effects that will be studied at a later stage.

### **b) CSEM static-shift corrections**

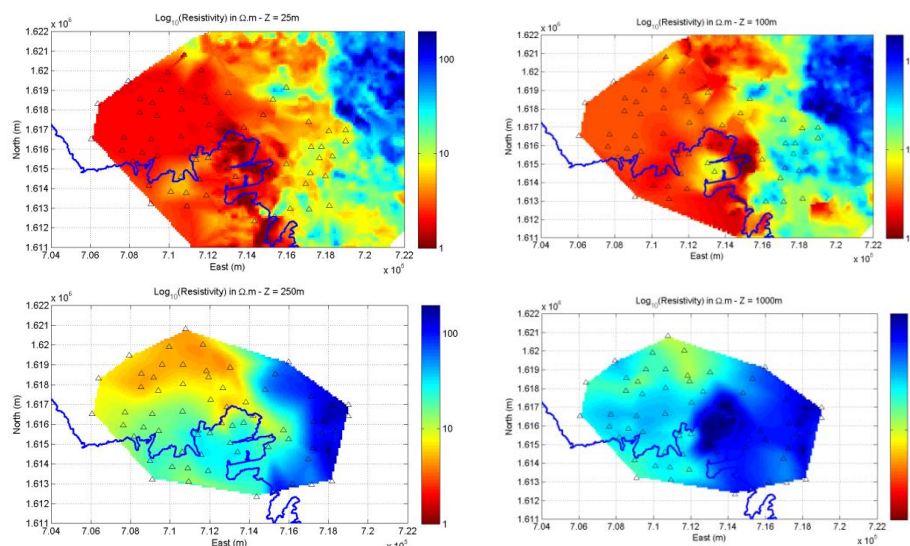
As evidenced by the airborne EM data (figure 2), the near-surface ( $<100\text{m}$ ) can be very heterogeneous and its electrical resistivity can increase from a few  $\Omega\cdot\text{m}$  up to a few 10's of  $\Omega\cdot\text{m}$  in a very short distance ( $<500\text{m}$ ), especially in the eastern side of the survey area where resistive silicified sediments can be found amongst conductive seawater filled sediments. Such a short contrast of resistivity will give rise to a static-shift effect on the CSEM measurements (MacLennan and Li, 2013), similar to what can be observed on MT data.

To correct for this effect, we have used the results from the airborne EM survey as inductive methods do not suffer from such effects (Sternberg, 1988). The CSEM static shift correction is calculated by fitting the high frequency amplitude response to the modelled response computed on the resistivity model obtained from the inversion of the airborne EM data, typically for frequencies greater than 100Hz. For CSEM stations located on the eastern side of the survey area and where shallow layers are rather heterogeneous (figure 2), significant static shift corrections are required. However, on the western part of the survey, shallow sediments are rather conductive and homogeneous and CSEM data do not require significant static-shift corrections.

### **c) 1D CSEM/AEM modelling/inversion results**

Once data have been processed and static-shift corrected, we inverted them by determining the resistivity of a multi-layered earth model for each CSEM station. During this process, we fixed the resistivity of the shallow layers with the resistivity obtained from the inversion of the airborne EM data. When not available, we used the high frequency CSEM dataset to estimate it. Fortunately, such points are located in the western area of the survey where shallow layers are rather homogeneous and conductive and hence unlikely to be biased by large static-shifts corrections.

Because of a lack of a fast and accurate algorithm to compute the electric field at various frequencies accounting for the complex transmitter geometry and current injection parameters (e.g. cable geometry, metallic boreholes casings), we have approximated the transmitter as a finite length dipole at the ground surface. 1D modelling of the current flowing in the metallic casings shows that this approximation is valid as the shallow layers are very conductive and most of the current exits the borehole in the shallow section. Figure 3 shows the resistivity maps recovered at 25m, 100m, 250m and 1000m depth over the area of interest.



*Figure 3: Resistivity maps recovered at 25m, 100m, 250m and 1000m depth from the 1D inversion of the CSEM dataset with static shift corrections from the AEM data. Black triangles represent the CSEM recording stations.*

### AEM/CSEM data interpretation

As evidence in the geothermal exploration wells, the very shallow conductive layers ( $< 5 \Omega.m$ , depth  $< 100m$ ) match well with the presence of unconsolidated sediments as well as highly altered volcano-detritic sediments. In the western part of the survey area, a deep conductive layer ( $5-10 \Omega.m$ , depth  $100m-500m$ ) is present and corresponds well to the presence of altered volcanic-detritic sediments. This could be the caprock of a geothermal system made of medium temperature alteration minerals ( $< 100 \text{ degC}$ ) like smectite (Ussher et al, 2000). At greater depths, the resistivity increases in the  $10-100 \Omega.m$  range and would correspond well to the former high temperature geothermal reservoir itself with high temperature alteration minerals like illite/chlorite. In the eastern part of the survey area, resistivities are high at all depths ( $> 100 \Omega.m$ ) and correspond well with the presence of the volcanic substratum composed of highly resistive low permeability sub-marine breccia.

### Conclusions

Exploring for underground resources using land-based electromagnetic methods can be very challenging in volcanic islands. We have shown that a combined airborne and land-based controlled-source EM approach can provide an attractive alternative to image resistivity variations at depth in such a context. In particular, it does efficiently tackle the issue of static-shift correction of CSEM data in the presence of near surface heterogeneities. We believe this work sets the foundations for a more robust geophysical exploration workflow in challenging EM environments.

### Acknowledgements

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