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O. Evrard, Jean-Patrick Laceby, Hugo Lepage, Yuichi Onda, Olivier Cerdan, et al.. Radiocesium transfer from hillslopes to the Pacific Ocean after the Fukushima Nuclear Power Plant accident: A review. Journal of Environmental Radioactivity, 2015, 148, pp.92-110. 10.1016/j.jenvrad.2015.06.018 . hal-01401008

HAL Id: hal-01401008 https://brgm.hal.science/hal-01401008v1

Submitted on 26 May 2020

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Radiocesium transfer from hillslopes to the Pacific Ocean after the Fukushima Nuclear Power Plant accident: A review

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Highlights

- The majority of radiocesium is transported in the particulate fraction.
- The contribution of the dissolved fraction is only relevant in base flows.
- Significant transfer of particulate-bound radiocesium occurs during heavy rainfall.
- Radiocesium deposited in floodplains may be remobilized, inducing contamination.
- Transdisciplinary approach is required to quantify radiocesium transfers.

1 Radiocesium transfer from hillslopes to the Pacific Ocean after the Fukushima

2 Nuclear Power Plant accident: A review

3

4 Abstract:

The devastating tsunami triggered by the Great East Japan Earthquake on March 11, 2011 5 6 inundated the Fukushima Dai-ichi Nuclear Power Plant (FDNPP) resulting in a loss of cooling 7 and a series of explosions releasing the largest quantity of radioactive material into the 8 atmosphere since the Chernobyl nuclear accident. Although 80% of the radionuclides from 9 this accidental release were transported over the Pacific Ocean, 20% were deposited over 10 Japanese coastal catchments that are subject to frequent typhoons. Among the radioisotopes released during the FDNPP accident, radiocesium (¹³⁴Cs and ¹³⁷Cs) is 11 considered the most serious current and future health risk for the local population. 12

13 The goal of this review is to synthesize research relevant to the transfer of FDNPP derived radiocesium from hillslopes to the Pacific Ocean. After radiocesium fallout deposition on 14 vegetation and soils, the contamination may remain stored in forest canopies, in vegetative 15 litter on the ground, or in the soil. Once radiocesium contacts soil, it is quickly and almost 16 17 irreversibly bound to fine soil particles. The kinetic energy of raindrops instigates the 18 displacement of soil particles, and their bound radiocesium, which may be mobilized and 19 transported with overland flow. Soil erosion is one of the main processes transferring particle-bound radiocesium from hillslopes through rivers and streams, and ultimately to the 20 21 Pacific Ocean. Accordingly this review will summarize results regarding the fundamental 22 processes and dynamics that govern radiocesium transfer from hillslopes to the Pacific 23 Ocean published in the literature within the first four years after the FDNPP accident.

24 The majority of radiocesium is reported to be transported in the particulate fraction and is 25 attached to fine particle size fractions. The contribution of the dissolved fraction to 26 radiocesium migration is only relevant in base flows and is hypothesized to decline over 27 time. Owing to the hydro-meteorological context of the Fukushima region, the most 28 significant transfer of particulate-bound radiocesium occurs during major rainfall and runoff 29 events (e.g. typhoons and spring snowmelt). There may be radiocesium storage within 30 catchments in forests, floodplains and even within hillslopes that may be remobilized and 31 contaminate downstream areas, even areas that did not receive fallout or may have been 32 decontaminated.

33 Overall this review demonstrates that characterizing the different mechanisms and factors 34 driving radiocesium transfer is important. In particular, the review determined that 35 quantifying the remaining catchment radiocesium inventory allows for a relative comparison of radiocesium transfer research from hillslope to catchment scales. Further, owing to the 36 variety of mechanisms and factors, a transdisciplinary approach is required involving 37 geomorphologists, hydrologists, soil and forestry scientists, and mathematical modellers to 38 comprehensively quantify radiocesium transfers and dynamics. Characterizing radiocesium 39 40 transfers from hillslopes to the Pacific Ocean is necessary for ongoing decontamination and 41 management interventions with the objective of reducing the gamma radiation exposure to 42 the local population.

43 Key Words (Max 6): Fukushima, Cesium, Soil, Sediment, Forest, Modelling

44

45 Graphical Abstract:



47 Introduction

The Great East Japan Earthquake on March 11, 2011 generated a giant tsunami inundating the 48 49 Fukushima Dai-ichi Nuclear Power Plant (FDNPP). After the loss of cooling at three FDNPP reactors, a 50 series of explosions occurred resulting in a massive release of radioactive contaminants into the 51 atmosphere. The FDNPP accident resulted in the largest atmospheric release of radioactive material since Chernobyl (Chino et al., 2011; Steinhauser, 2014; Thakur et al., 2013). Although 80% of the 52 53 fallout from this atmospheric release was over the ocean, 20% of fallout was over terrestrial Japan (Groëll et al., 2014; Kawamura et al., 2011; Kobayashi et al., 2013). The terrestrial fallout from the 54 55 FDNPP accident occurred in a different environmental context than previous major accidents, 56 occurring predominantly in a coastal mountainous region subject to frequent typhoons (Hilton et al., 57 2008; Shinomiya et al., 2014).

Radiocesium (¹³⁴Cs and ¹³⁷Cs) is considered the most serious risk to the current and long term health of the local population (Kitamura et al., 2014; Saito et al., 2015). The half-life of the two most abundant radiocesium isotopes is 2 y (¹³⁴Cs) and 30 y (¹³⁷Cs). Therefore, FDNPP derived radiocesium will be an ongoing, long-term source of gamma radiation exposure to the local population. Accordingly, this review will focus on radiocesium transfer from hillslopes to the Pacific Ocean.

Shortly after the FDNPP accident, multiple other radioisotopes were measured in the environment. 63 The wet deposition of hazardous iodine (¹³¹I, half-life of 8 d) was detected along with multiple short-64 lived radionuclides (Shozugawa et al., 2012; Tagami et al., 2011; Tazoe et al., 2012) (Table 1). 65 However, these short-lived isotopes disappeared quickly and were primarily below detection limits 66 by late April 2011 (Fujiwara et al., 2012). Long-lived isotopes (e.g., 239 Pu (T_{1/2} = 24100 y), 240 Pu (T_{1/2} = 67 6560 y)) were also detected (Table 1), although only in trace levels (Evrard et al., 2014b; Schneider et 68 al., 2013; Schwantes et al., 2012). Steinhauser (2014) comprehensively reviewed radionuclides that 69 may have been released by the FDNPP accident which have received less attention. These 70 radionuclides include less volatile elements (⁸⁹Sr, ⁹⁰Sr, ¹⁰³Ru, ¹⁰⁶Ru), pure beta-emitters (³H, ¹⁴C, ³⁵S), 71

- 72 gaseous radionuclides (⁸⁵Kr, ¹³³Xe, ¹³⁵Xe) or radionuclides with very long half-lives (e.g., ³⁶Cl, ⁹⁹Tc, ¹²⁹I,
- 73 ²³⁶U)(Table 1).

 Table 1. Radionuclides detected in environmental samples after the FDNPP accident

 Half-life

Radionuclide	Half-life (d)	Half-life (y)	Reference
¹²⁹ Te	0.05		Tazoe et al. (2012)
¹³²	0.10		Shozugawa et al. (2012)
¹³⁵ Xe	0.38		Steinhauser (2014)
⁹¹ Sr	0.40		Shozugawa et al. (2012)
¹⁴⁰ La	1.70		Shozugawa et al. (2012); Tagami et al. (2011)
^{133m} Xe	2.20		Steinhauser (2014)
²³⁹ Np	2.35		Shozugawa et al. (2012)
¹³² Te	3.30		Shozugawa et al. (2012))
¹³³ Xe	5.25		Steinhauser (2014)
¹³¹	8.00		Shozugawa et al. (2012); Tagami et al. (2011); Tazoe et al. (2012)
^{131m} Xe	11.90		Steinhauser (2014)
¹⁴⁰ Ba	12.80		Shozugawa et al. (2012)
¹³⁶ Cs	13.20		Tagami et al. (2011)
^{129m} Te	33.60		Tagami et al. (2011); Tazoe et al. (2012)
⁹⁵ Nb	35.00		Shozugawa et al. (2012); Tagami et al. (2011)
¹⁰³ Ru	39.40		Steinhauser (2014)
⁵⁹ Fe	44.60		Shozugawa et al. (2012)
⁸⁹ Sr	50.53		Steinhauser (2014)
⁹¹ Y	58.50		Shozugawa et al. (2012)
⁹⁵ Zr	64.00		Shozugawa et al. (2012)
³⁵ S	87.37		Steinhauser (2014)
^{110m} Ag	252.00		Shozugawa et al. (2012); Tagami et al. (2011); Tazoe et al. (2012)
¹³⁴ Cs		2.06	Tagami et al. (2011); Tazoe et al. (2012)
⁸⁵ Kr		10.75	Steinhauser (2014)
³ H		12.32	Kakiuchi et al. (2012); Steinhauser (2014)
²⁴¹ Pu		14.33	Steinhauser (2014)
¹³⁷ Cs		30.20	Shozugawa et al. (2012); Tagami et al. (2011); Tazoe et al. (2012)
²³⁸ Pu		87.70	Steinhauser (2014)
²⁴¹ Am		432.60	Steinhauser (2014)
¹⁴ C		5730.00	Steinhauser (2014)
²⁴⁰ Pu		6561.00	Steinhauser (2014)
²³⁹ Pu		24110.00	Shozugawa et al. (2012); Steinhauser (2014)
⁹⁹ Tc		2.10E+05	Steinhauser (2014)
³⁰ Cl		3.00E+05	Steinhauser (2014)
1 236		1.57E+07	Steinhauser (2014)
²³⁰ U		2.34E+07	Steinhauser (2014)
²³³ U 238		7.04E+08	Steinhauser (2014)
- U		4.47E+09	Steinhauser (2014)

75 Researching less abundant radionuclides may provide useful and complementary information on the 76 events that occurred at FDNPP (Achim et al., 2014; Le Petit et al., 2014). However, owing to the potential gamma radiation exposure, there has been an increasing focus in the literature on the 77 78 transfer, fate, dynamics and consequences of radiocesium in the fallout impacted region. Before the 79 accident, there were on average six publications a year with Fukushima in the title (Fig. 1). After the 80 accident (from 2011 to 2014), this average increased over 50 fold to approximately 350 publications 81 a year. Of these publications, the percent of published research including radiocesium as a topic 82 increased from 8% in 2011 to 32% in 2014 indicating that radiocesium, in the fallout impacted region, 83 is of increasing significance to the scientific community.



Figure 1. Annual publications with Fukushima in the title (A - top), and Fukushima in the title with Cesium/Caesium included the topic (B - bottom) from a Web of Science search conducted early in January 2015.

85 Initially radiocesium was deposited on terrestrial vegetation and soils. Depending on local conditions, the contamination may remain stored in the forest canopy, in the vegetative litter on the ground or 86 in the soil. Once radiocesium contacts soil, it is quickly and almost irreversibly bound to fine soil 87 particles (He and Walling, 1996; Sawhiney, 1972; Tamura, 1964). During rainfall, the kinetic energy of 88 89 raindrops instigates the displacement of soil particles and their bound radiocesium, which are mobilized and transported with rainfall runoff (Renard et al., 1991; Torri et al., 1987)(Fig. 2). Soil 90 91 erosion is considered to be the main process transferring particle-bound radiocesium from hillslopes 92 through rivers and streams, and ultimately to the Pacific Ocean (Yoshida and Kanda, 2012)(Fig. 2).



93

Figure 2. Potential pathways for radiocesium transfer from hillslopes to the Pacific Ocean following the FDNPP
 accident denoting major processes such as erosion (E), deposition (DP), decontamination (DC) and desorption
 (DS) (Images courtesy of the Integration and Application Network, University of Maryland Center for
 Environmental Science (ian.umces.edu/symbols/)).

98

99 The transfer of particulate-bound radiocesium may be substantial during typhoons typically occurring 100 between June and October in Japan. Importantly, the fine particles, to which radiocesium is 101 preferentially bound to, are the particles that are preferentially eroded and likely to be transported the farthest during erosive rainfall events (Govers, 1985; Malam Issa et al., 2006; Motha et al., 2002).
Fundamentally, soil erosion and riverine transport processes are the primary factors governing the
long-distance transfer of radiocesium from hillslopes to the Pacific Ocean (Kitamura et al., 2015; Saito
and Onda, 2015; Yamada et al., 2015).

106 The goal of this review is to synthesize research relevant to the transfer of FDNPP derived 107 radiocesium outlined in the previous paragraph. The main objective is to comprehensively review 108 literature investigating the fundamental processes and dynamics governing radiocesium transfer 109 from hillslopes to the Pacific Ocean published within 4 y of the FDNPP accident. This review will 100 conclude with recommendations for future research and for the design and implementation of 111 effective countermeasures.

112 **1. Fallout radiocesium deposition**

113 Although radionuclides from the FDNPP accident were detected worldwide (Evrard et al., 2012; Long et al., 2012; Masson et al., 2011; Wetherbee et al., 2012), the majority of terrestrial deposition 114 115 occurred in the Fukushima Prefecture. The main radiocesium plume is located northwest of the 116 FDNPP, with a secondary plume in the Abukuma River Valley southwest from Fukushima City past 117 Koriyama (Fig. 3). There are multiple approaches to quantify catchment radiocesium inventories 118 including both airborne and soil sampling surveys (Gonze et al., 2014; Saito et al., 2015; Yoshida and 119 Takahashi, 2012). As this review examines the transfer of radiocesium from hillslopes to the Pacific 120 Ocean, we used soil sampling surveys to quantify the total, mean, and standard deviation of radiocesium (¹³⁴Cs+¹³⁷Cs) fallout in 14 coastal catchments of the Fukushima region (Fig. 3, Table 2). 121 122 This quantification of catchment inventories is based on the spatial interpretation conducted by 123 Chartin et al. (2013) of 2200 soil samples collected by the Japanese Ministry of Education, Culture, 124 Sports, Science and Technology (MEXT, 2012) between June and July 2011.



Figure 3. Labelled coastal catchments impacted by the fallout from the FDNPP accident, major rivers (grey lines), the location of the FDNPP (black circle) and major cities (black squares), and the spatial distribution of radiocesium fallout (134 Cs + 137 Cs) in kBq/kg provided by Chartin et al. (2013) who spatially interpreted (ordinary kriging) the radiocesium distribution of 2200 soil samples taken by the Japanese Ministry of Education, Culture, Sports, Science and Technology (MEXT) between June and July 2011 decay corrected to June 14, 2011.

Table 2. Radiocesium (¹³⁴Cs+¹³⁷Cs) inventory received by the major coastal catchments in Fukushima region (Coastal catchment data derived from the Japanese Ministry of Land, Infrastructure, Transport and Tourism (http://nlftp.mlit.go.jp/ksj-e/index.html)) and the radiocesium inventories were calculated with data from Chartin et al. (2013) spatially interpreted radiocesium inventory calculations from Fig. 3 converted to total Cs from Bg/kg.

		Total	% of	Deposition
	Area	Inventory	Total	Density
Catchment	(km²)	(TBq)	Inventory	(TBq km ^{⁻²})
Abukuma	5338	734.9	30%	0.14
Ukedo	423	643.9	26%	1.52
Niida	265	283.4	12%	1.07
Kuma	75	161.4	7%	2.14
Maeda	45	149.9	6%	3.32
Ota	77	111.0	5%	1.44
Odaka	68	81.4	3%	1.20
Mano	171	78.1	3%	0.46
Tomioka	62	58.6	2%	0.95
Kido	263	48.6	2%	0.18
Natsui	696	36.3	1%	0.05
Uda	101	27.6	1%	0.27
Same	601	19.2	1%	0.03
Ide	40	16.2	1%	0.14

126

127 The Abukuma catchment received the most radiocesium fallout, followed by the Ukedo and Niida 128 catchments (Table 2). Radiocesium inventories for the 14 coastal catchments ranged between 734.9 TBq in the Abukuma to 16.2 TBq in the Ide catchment. The Abukuma catchment received 129 130 approximately 30% of the fallout received by these 14 catchments, followed by 26% for the Ukedo 131 and 12% for the Niida. Although the Abukuma catchment received the highest amount of fallout, the 132 radiocesium was distributed over a broad spatial area and the mean deposition density in the Abukuma catchment was 0.14 TBq km². Comparatively, the smaller coastal catchments had the 133 highest mean radiocesium inventories, particularly the Meada (3.32 TBq km⁻²), Kuma (2.14 TBq km⁻²), 134 Ukedo (1.52 TBq km⁻²), Ota (1.44 TBq km⁻²), Odaka (1.20 TBq km⁻²), and Niida (1.07 TBq km⁻²) 135 catchments. 136

137 Although quantifying catchment inventories with airborne and soil based sampling techniques or 138 different spatial interpretation approaches may yield different results (Gonze et al., 2014), what is important to the research and management community is the spatial distribution of radiocesium. In particular, characterizing the deposition of radiocesium, especially the different catchments and landscapes receiving the fallout (e.g different forest types), is important for understanding the sources and spatial distribution of radiocesium available for transfer from hillslopes to the Pacific Ocean.

144 2.1 Forest deposition

145 Significant radiocesium deposition occurred on forested landscapes. Approximately 66% of high levels of radioactive fallout (> 1000 kBq.m⁻²) took place over a mixture of deciduous and coniferous 146 147 forests (Hashimoto et al., 2012). As the accident occurred during late winter and deciduous trees had 148 no leaves, the fallout more directly contaminated the ground surface (soil and litter). In coniferous 149 forests, with their needles intact, the fallout radiocesium was partly intercepted by the forest canopy. In particular, Kato et al. (2012a) estimated that between 62% and 65% of ¹³⁷Cs was 150 151 intercepted by the canopies of cypress and cedar forests, respectively. When examining both deciduous and coniferous forests, Koizumi et al. (2013) modelled that ~10% of the ¹³⁷Cs deposited on 152 153 forests within the exclusion zone was trapped by biomass.

The majority of research has focused on Japanese Cedar (*Cryptomeria japonica*) followed by the mixed forests and coniferous forest with broad leaf forests receiving the least attention (Fig. 4). The research, in particular research examining Japanese Cedar, has been widespread throughout the fallout impacted region and across northern Japan (Fig. 4). As the forested areas within 20 km of the FDNPP (Koizumi et al., 2013) are approximately 50% broadleaf and 50% needle-leaf forest, it may be advantageous for future research to examine radiocesium dynamics equally in broadleaf, coniferous and mixed forests within the highly contaminated coastal catchments.

After canopy interception, the radiocesium transfer pathway from the canopy to the soil is complex (Fig. 5). Radiocesium may transfer through leaf surface tissues or remain on the foliar surface. A portion of foliar radiocesium (17 – 59%) was measured inside leaves within a year, whereas the

remaining radiocesium adhered to the leaf surface may be removed with wash off over time (Koizumi et al., 2013; Nishikiori et al., 2015; Tanaka et al., 2013a). Results were similar for deciduous (46% removed after washing) and coniferous canopies (41%) (Koizumi et al., 2013) indicative of a potential long-term radiocesium source where canopies intercepted significant quantities of radiocesium.



Figure 4. Localisation of samples examining radiocesium processes and dynamics in forests after the FDNPP Accident. Multiple samples may have been obtained at one location.

During the first year after the accident, foliar contamination was mainly due to the radiocesium deposition. After the initial deposition, radiocesium may translocate between old and new vegetative tissues (Nishikiori et al., 2015; Yoshihara et al., 2014b). In most studies, a decrease of contamination was observed between old and young needles in coniferous forests (Akama et al., 2013; Kanasashi et al., 2015; Nishikiori et al., 2015; Yoshihara et al., 2013). In contrast, Yoshihara et al. (2014b) reported that new foliar growth in 2013 contained more radiocesium than 2011 samples in coniferous forests.

- 175 The variation between older and younger foliar components of coniferous forests contaminated by
- the accident is complex and may impact the radiocesium export from these forests.
- 177



Figure 5. An illustration of the major radiocesium transfer pathways in forested (coniferous – to the left – and deciduous – to the right) landscapes. The major transfer pathways to the soil include throughfall (TF), canopy interception (CI), litter fall (LF) and stem flow (SF). There is also potential for the flora to uptake (UT) radiocesium (Images courtesy of the Integration and Application Network, University of Maryland Center for Environmental Science (ian.umces.edu/symbols/)).

- 178 For deciduous forests, research into foliar contamination is inconclusive. Tanaka et al. (2013a) did not
- 179 observe any contamination in new leaves. Nishikiori et al. (2015) reported that although leaves were
- absent during the accident, contamination was detected in leaves sampled in 2012. Yoshihara et al.
- 181 (2014a) found no significant variation in leaves sampled in 2011 and 2012. Additional factors may
- 182 explain tree contamination such as root uptake, stem transfer, or surficial contamination due to
- 183 wind-eroded particle deposition.
- 184 Characterizing radiocesium transfer from forests to hillslopes is important for understanding 185 cumulative long-term radiocesium export from forest environments. During rainfall events, 186 radiocesium is more likely to be transferred by throughfall than by stemflow, although both

pathways increase ground contamination (Kato et al., 2012a; Loffredo et al., 2014). The effective
half-life of radiocesium in trees was calculated to be 620 d for cypress and 890 d for cedar canopies
(Kato et al., 2012a). Accordingly, there will likely be a medium to long-term canopy to ground
stemflow and throughfall transfer of radiocesium (Kato et al., 2012a).

191 Litterfall is a significant radiocesium transfer pathway. In fact, litterfall was reported to transfer as 192 much radiocesium as throughfall (Teramage et al., 2014a). Further, Hasegawa et al. (2013) reported 193 that radiocesium concentrations in litter were greater than those measured in the soil. In the future, 194 litterfall will likely become the dominant pathway for radiocesium transfer from the canopy to the 195 forest floor as hydrological pathways transfer less radiocesium with time owing to decreased 196 availability (Teramage et al., 2014a). In addition, Kuroda et al. (2013) found that the tree bark was 197 highly contaminated. More than 94% of the stem contamination was located in the bark, indicative 198 of another potential long-term source of radiocesium derived either from surficial contamination or 199 through contaminated phloem sap.

200 The quantity of radiocesium in the litter depends on multiple factors such as tree species, location of 201 the forest in the radioactive pollution plume, or sampling date. Matsunaga et al. (2013) did not 202 observe specific migration of radiocesium from litter to soil after the first rainfall event (July-August 2011). Nakanishi et al. (2014) reported a decrease of ¹³⁷Cs contamination in litter from 67% to 13% 203 204 during the first year, whereas contamination in soil increased from 26% to 80%. Takahashi et al. 205 (2015) found that the proportion of radiocesium in leaf litter, relative to soil, decreased from ~90% in 206 2011 to 18-14% in 2012. Fujii et al. (2014) also noted a significant decrease in the percentage of 207 radiocesium in the leaf litter, decreasing from between 44-65% in 2011 to 14-27% in 2012. The 208 transfer of radiocesium from litter to soil is controlled by rainfall and litter decomposition. Research 209 estimates that around 30%-40% of litter in Japanese forests is decomposed annually (Hashimoto et 210 al., 2012; Matsunaga et al., 2013).

211 In the Fukushima region, forests are a radiocesium reservoir, retaining significant quantities of 212 radiocesium within both the canopy and the leaf litter (Hashimoto et al., 2012; Koarashi et al., 213 2012a). Although forested landscapes are not considered to be significant erosion sources 214 (Yoshimura et al., 2015a), there is an abundance of radiocesium within these systems, often stored in 215 organic matter, that may be transferred downstream during significant rainfall events. Accordingly, 216 forest landscapes may potentially become perennial sources of radiocesium in the Fukushima region. 217 Understanding both intra- and inter- forest radiocesium transfers will be fundamental to 218 understanding radiocesium transfer and export by soil erosion and wash-off. Incorporating 219 differences between deciduous and coniferous forests and their different rates of radiocesium 220 transfer is important for modelling hillslope exports and cumulative radiocesium transfers from 221 hillslopes to the Pacific Ocean.

222 2.2 Soil deposition

Upon contact with the soil, it is well known that radiocesium is strongly and rapidly fixated to fine soil particles (He and Walling, 1996; Sawhiney, 1972). Cesium is an alkali metal and accordingly it possesses similar chemical properties to potassium and may substitute for potassium in interactions with clay minerals resulting in an almost irreversible fixation to fine soil particles (Kitamura et al., 2015).

Following the Fukushima accident, researchers have extensively examined radiocesium fixation to fine particles (Akai et al., 2013; Saito et al., 2014; Tanaka et al., 2013c). These examinations were undertaken primarily because of the high organic matter and clay content of soils in the Fukushima region that may influence the depth migration of radiocesium in the soil profile. In particular, there has been a research focus on the physical and chemical foundation of the fixation of radiocesium to fine particles and the depth migration of radiocesium in the soil profile. Characterizing the fixation potential of radiocesium to fine soil particles is fundamental to understanding potential radiocesium 235 migration pathways including the mobilization and transport of radiocesium, and ultimately its
 236 transfer through rivers and streams to the Pacific Ocean.

237 Compared to the forest radiocesium research, there has not been a similar regional coverage of soil 238 sampling across northern Japan (Fig. 6). The majority of samples examining soil transfer processes 239 including depth migration or fixation potential were taken within the 14 coastal catchments with a 240 well distributed coverage across the fallout impacted region. The focus of soil sampling is mainly on 241 paddy fields with a substantial number samples having unclassified land uses. The inclusion of land 242 use details would be beneficial to future research examining radiocesium transfer processes.



Figure 6. Location of soil sampling locations where research examined radiocesium dynamics or process such as depth migration in the soil profile or fixation to the mineral or soil fractions. More than one sample may occur at any given location.

- 243 Immediately after the FDNPP accident, Tanaka et al. (2013c) reported that the radiocesium was
- water soluble, though once radiocesium adsorbed on to soil and rock particles it quickly became
- essentially insoluble. The increase in insolubility was also reported by Shiozawa (2013) who indicated

that the strong fixation of radiocesium to clay particles during the initial 2–3 months after the accident reduced the migration of radiocesium in the depth profile. The strong fixation to clay particles was again reported by Saito et al. (2014) and Kikawada et al. (2014) who indicated that virtually no radiocesium was extractable from soil samples with simple water solutions.

250 Using X-ray diffraction, Saito et al. (2014) reported that micaceous minerals (e.g. muscovite, biotite, 251 and illite) were responsible for the strong radiocesium soil fixation. Conversely, Kozai et al. (2012) 252 utilized desorption experiments to determine that radiocesium was almost irreversibly fixated to the non-micaceous soil minerals. Niimura et al. (2015) reported that radiocesium fixation is due to a 253 254 physical adherence to the rough surfaces of the soil mineral particles. These different approaches to 255 understanding radiocesium fixation indicate that the physical processes behind this strong fixation 256 are difficult to characterize and are dependent upon multiple parameters such as experiment 257 characteristics, mineral contents, and/or the physical-chemical characteristics of both the fallout 258 radiocesium itself and the soil medium.

259 The extractability of radiocesium was found to be higher in forest soils with higher contents of clay 260 and silt particles, and a higher cation exchange capacity (Koarashi et al., 2012b). Using extraction 261 processes with K₂SO₄, Koarashi et al. (2012b) reported that only 2.1 to 12.8% of Fukushima-262 originated radiocesium found in the surface mineral soils were retained as easily exchangeable ions by abiotic components in forest sites. Potential uptake and retention of ¹³⁷Cs by microorganisms in 263 264 the forest surface soils (0-3 cm) were found to be minor and less important than ion-exchange 265 adsorption on non-specific sites provided by abiotic components (Koarashi et al., 2012b). In the long-266 term, microorganisms in forest soils may contribute to the mobilization of radiocesium by 267 decomposition of the soil organic matter (e.g. leaf litter) (Koarashi et al., 2012b).

Research prior to the FDNPP accident clearly indicates that in undisturbed soils an exponential decline of radiocesium in the soil profile is expected along with a very slow radiocesium depth migration in the soil profile (He and Walling, 1997). In fact, the radiocesium peak from nuclear

weapons testing in the 1960s is now only 2-8 cm below the soil surface in undisturbed soils(Jagercikova et al., 2015).

273 Owing to the strong and irreversible bond to fine soil particles, the majority of FDNPP derived 274 radiocesium currently remains stored within the top 5 cm of the soil profile in undisturbed soils 275 (Lepage et al., 2014b; Matsuda et al., 2015; Matsunaga et al., 2013; Mishra et al., 2015; Takahashi et 276 al., 2015). Anthropogenic activities increase the depth of migration of radiocesium in the soil profile 277 (Lepage et al., In Press; Matsunaga et al., 2013). Erosion and subsequent deposition processes may 278 also result in the increase of radiocesium with depth in the soil profile similarly to carbon and other 279 minerals in depositional landscapes (Van Oost et al., 2012). Lepage et al. (In Press) demonstrated that 280 90% of radiocesium remained in the upper 5 cm of the soil profile after deposition in non-cultivated 281 soils, whereas within cultivated soils the contamination was homogeneous throughout the tilled soil 282 layer (generally down to 15-30 cm). Endo et al. (2013) demonstrated that radiocesium 283 concentrations were not dependent on depth in rice paddy field soils, whereas they declined 284 exponentially in an uncultivated soil. Sakai et al. (2014) and Tanaka et al. (2013b) both indicated that 285 radiocesium from the FDNPP accident was measureable at 15 cm depth in rice paddy fields. Mishra 286 et al. (2015) reported that the vertical migration down the soil profile was slower in forest soils than in grassland soils. In a study of 85 soil sampling sites, Matsuda et al. (2015) found that the depth 287 penetration of radiocesium was higher than the standard range (0.2-1.0 kg m⁻² y⁻¹) depicted by He 288 289 and Walling (1997).

290 Multiple researchers have postulated different hypotheses for the observed depth migration of 291 radiocesium in the soil profile. Fujiwara et al. (2012) reported that the elevated rates of radiocesium 292 migration down the soil profile are potentially indicative of the mobile form of radiocesium 293 immediately after the accident. Kato et al. (2012b) also observed increased downwards migration of 294 radiocesium immediately and attributed this phenomenon to high clay content that increased the 295 aggregate stability of topsoil resulting in increased water and radiocesium infiltration deeper into the

soil profile. Conversely, Satou et al. (2015) investigated the vertical migration in coarse-grained beach
material impacted by the tsunami. The authors demonstrated that the majority of radiocesium was
located at ~3-6 cm depth. This migration is likely due to coarser grained material indicative of a
potential particle size impact on the depth migration of radiocesium.

300 The depth migration of radiocesium also depends on soil properties. Koarashi et al. (2012a) 301 demonstrated that clay and organic matter content are key factors controlling radiocesium depth 302 migration. Organic matter inhibits the fixation of radiocesium on the soil mineral fraction by limiting 303 its access to adsorption sites on clay minerals through direct competition for potential adsorption 304 sites. As organic matter may limit the adsorption of radiocesium, it may also influence the rate of 305 radiocesium depth migration. The impact of organic matter is particularly notable in forests. The high 306 organic content of forest soils increases the potential for radiocesium depth migration (Koarashi et 307 al., 2012a). Nakanishi et al. (2014) demonstrated that 0.1% of the radiocesium deposited on the 308 forest floor indeed migrated below 10 cm depth in 1 y.

309 The timing and extent of rainfall may also drive the migration of radiocesium deeper within the soil 310 profile. The rainfall that occurred immediately after the accident fallout may have been a factor in 311 the radiocesium depth migration in forest soils (Teramage et al., 2014b) along with rainfall volume in 312 general (Fujii et al., 2014). Further, rainfall was listed as a potential factor for the depth migration in 313 the beach sand samples (Satou et al., 2015). The average rainfall in the Fukushima region is more 314 than 2 times greater than Chernobyl and the rainfall erosivity is almost 3 times greater (Laceby et al., 315 submitted). Both the elevated rainfall intensity and volume may have resulted in greater depth 316 penetration of radiocesium in the Fukushima region compared to that of Chernobyl.

Migration of radiocesium was also correlated with low bulk densities as highly porous structures may facilitate water and potentially dissolved radiocesium infiltration deeper down the soil profile (Koarashi et al., 2012a; Takahashi et al., 2015). The radiocesium interception potential (RIP) is often used to quantify this process. The RIP describes the capacity of the mineral content of soils to adsorb

321 radiocesium (Wauters et al., 1996). Fan et al. (2014b) reported the mean RIP from 13 soil samples was 6.1 (with all discussed RIP values noted as mol.kg⁻¹). Nakao et al. (2014) noted a higher mean RIP 322 323 of 7.8 for 97 soil samples. Takahashi et al. (2015) reported a mean RIP of 4.1 for 7 soil samples and 324 demonstrated that RIP was negatively correlated with a penetration coefficient (α), indicating that 325 the radiocesium penetrates deeper in soils with low RIP values. A similar relationship was found 326 between RIP values and the organic content of the soil with Takeda et al. (2014) confirming deeper 327 radiocesium penetration in soils with elevated organic content. In particular, Takeda et al. (2014) 328 reported a mean RIP for 46 samples of 2.4. Further, Koarashi et al. (2012a) reported a strong 329 negative correlation (r = -0.79, p < 0.005) between total organic carbon (TOC) content divided by clay content (TOC/clay) and the percentage of retention of ¹³⁷Cs, suggesting that the presence of organic 330 matter inhibits the adsorption of ¹³⁷Cs on clay minerals. 331

332 Along with TOC, RIP depends on the presence of minerals having selective adsorption sites for Cs 333 (micas, micaceous clays) (Takeda et al., 2014). Indeed, the type of clay may be more important than 334 the total clay content as there were higher RIP values in soils containing clays from the micaceous, 335 vermiculitic and chloritic groups (>10 mol/kg)(Takahashi et al., 2015) than in soils containing clays of 336 amorphous and kaolinitic groups (Nakao et al., 2014). In the Fukushima region where some soils 337 developed on volcanic ash deposits, they contain higher amounts of poorly crystalline minerals (e.g., 338 allophane, imogolite), which do not provide these adsorption sites (Takeda et al., 2014). There have 339 even been positive relationship between RIP and native K content which further reflects the fact that the ¹³⁷Cs retention ability of soil clays is controlled primarily by the amount of micaceous minerals 340 (Nakao et al., 2014). Ultimately, owing to the variety of relationships between TOC, clay content and 341 342 clay species, quantifying the dominant factors influencing RIP in the Fukushima region is more complicated than in the Chernobyl region. 343

The research on radiocesium deposition on soils has focussed on understanding post-deposition dynamics. In particular, understanding the RIP and concomitant fixation to soil is important to

understand post-deposition processes with important ramifications for management. For example,
owing to the strong fixation potential to soils, radiocesium will be maintained in the upper sections
of undisturbed soil allowing for potential decontamination through the removal of the soil surface
(e.g. top 5 cm) in undisturbed sites. Further, as radiocesium is strongly fixated to fine mineral soils, it
potentially may be eroded during rainfall events and its migration and transfer from hillslopes to the
Pacific Ocean will be governed by the geomorphological and hydrological factors driving soil erosion,
mobilization, and transport processes.

353 2. Radiocesium hillslope transfers

354 As radiocesium is tightly bound to soil particles, the transfer of soil-bound radiocesium can occur 355 with the physical mobilization of soil particles during erosion processes (Yoshimura et al., 2015a). The 356 amount of radiocesium wash-off generated by erosion processes is dependent upon the volume of 357 soil eroded and the radiocesium concentration in the soil, with the latter being dependent on the 358 amount of radiocesium within the soil (Yoshimura et al., 2015a). In the coastal catchments of the 359 Fukushima Prefecture, sheetwash erosion on cultivated slopes dominates, as slope gradients on 360 cultivated hillslopes are low and their length is insufficient to allow runoff to concentrate and 361 generate rills or gullies (Vandaele et al., 1996).

362 To examine soil and radiocesium export, Yoshimura et al. (2015a) established seven Universal Soil 363 Loss Equation (USLE) run-off plots in Kawamata Town. They examined one forest, two cultivated, two 364 uncultivated, and two grassland sites and reported that land use is important when modelling the 365 export of radiocesium. In particular, one uncultivated farmland plot dominated the radiocesium and sediment runoff. The farmland plots had the highest rates of erosion (1 kg m⁻² of soil loss) and ¹³⁷Cs 366 367 export (16 kBq m⁻²), followed by grassland and the forest plots (< 0.1 kg m⁻² of soil loss and < 2 KBq m⁻ ² of ¹³⁷Cs). In particular, Yoshimura et al. (2015a) reported an exponential relationship between 368 369 vegetation cover and soil loss.

370 Quantifying hillslope erosion is difficult for Japanese paddy fields as they may be both sources and 371 sinks of sediment and sediment-bound radiocesium (Tanaka et al., 2013b; Wakahara et al., 2014). 372 When paddies are puddled for the rice growing season, they are flooded with water potentially 373 containing suspended sediment and adsorbed radiocesium (Endo et al., 2013; Wakahara et al., 2014; 374 Yoshikawa et al., 2014). Further, when the paddies are drained, they are directly connected to the 375 stream network and may become a significant source of sediment and radiocesium, as water and 376 sediments discharged from paddy fields are often transferred directly to stream systems (Wakahara 377 et al., 2014).

Tanaka et al. (2013b) investigated the impact of contaminated irrigation water on rice paddy radiocesium inventories. Although they did not report a significant impact of irrigation water on the radiocesium inventory in the paddy, they concluded that significant rainfall events may increase radiocesium migration and transfer. In a non-hillslope erosion oriented study, Yoshikawa et al. (2014) found that the irrigation of paddy fields only contributed an additional 0.03-0.05% of ¹³⁷Cs to the preirrigation paddy field ¹³⁷Cs inventory. The authors concluded that the additional ¹³⁷Cs load from irrigation is negligible in comparison to the existing soil ¹³⁷Cs inventory.

Wakahara et al. (2014) examined the discharge of radiocesium from rice paddy fields reporting that a normally cultivated paddy field exported 1240 Bq m⁻² of ¹³⁷Cs after puddling and 102 Bq m⁻² after irrigation. Comparatively, a decontaminated paddy exported 48 Bq m⁻² of ¹³⁷Cs after puddling and 317 Bq m⁻² after irrigation. Although decontamination removed 95% of the fallout radiocesium from this particular paddy, the export of radiocesium from the decontaminated paddy was greater after irrigation than the non-decontaminated paddy. The authors concluded that water inflows during events and irrigation with sediment-bound ¹³⁷Cs contributed to this counter-intuitive result.

There has been limited research on the radiocesium migration from forests. Kitamura et al. (2014) modelled that forest soils, which occupy a large portion of the region, export only a small fraction of soil and radiocesium. Pratama et al. (2015) modelled that subcatchments of the Abukuma River with 395 greater proportions of forest coverage have the lowest radiocesium export. Nagao et al. (2013) 396 confirmed this finding in their monitoring research through indicating that, while forest land uses 397 cover 64% of their research catchments, they only accounted for 24% of soil and 41% of ¹³⁷Cs export. 398 Although forests disproportionately contribute less radiocesium than their initial inventories, the 399 gross amounts exported are high owing to their large area (Kitamura et al., 2014).

400 Fig. 7 illustrates the limited extent of direct hillslope erosion sampling in the Fukushima region. The 401 research on hillslope erosion was located in the Abukuma catchment. This limited research reflects a 402 similar lack of research in monsoonal regions on soil erosion owing to the difficulties in conducting 403 complex field observation studies during intense rainfall conditions or replicating these conditions in 404 laboratory environments. More research at the hillslope scale is important to quantify the erosion 405 and concomitant radiocesium transfer during runoff events. Different erosion and runoff dynamics 406 exist for each land use, complicating our understanding of the detachment and transfer of 407 contaminated sediments, particularly in catchments with mixed land use (Kinouchi et al., 2015). As 408 radiocesium is bound to soil, its initial mobilization and transfer will be directly dependent on the 409 various factors governing soil erosion (for example see Morgan et al. (1998); Nearing et al. (1989); 410 Renard et al. (1991)). Thoroughly characterizing potential soil erosion processes from different 411 sources at the hillslope scale is thus fundamental to understanding the potential transfer from 412 hillslopes to the Pacific Ocean. In particular, more research is required to properly characterize the 413 dominant sources of particulate radiocesium in the Fukushima region and the relative contribution of 414 these sources to radiocesium fluxes to the Pacific Ocean.

415 **3. Factors influencing riverine radiocesium transfer**

As radiocesium mobilized from hillslopes and transported downstream is selectively attached to the finest eroded soil particles (Kinouchi et al., 2015), it is important to develop a quantitative understanding of the particle size redistribution of radiocesium to predict potential contamination of downstream areas that are more densely populated and not-affected by the initial fallout (Kinouchi

- 420 et al., 2015; Ueda et al., 2013). Characterizing the particle size distribution of radiocesium will
- 421 improve our understanding of the potential transfer of radiocesium (e.g. fine particle fractions



Figure 7. Sample locations for research examining hillslope radiocesium transfer with USLE-based and calculations of suspended sediment yields form rice paddy fields.

422 transfer farther and possibly stored longer) and also potential storage within stream systems (e.g. 423 coarse particles are deposited sooner). Along with the particulate fraction, it is important to understand the dissolved/particulate partitioning of transported radiocesium in rivers and streams. 424 425 Radiocesium is transferred in both dissolved and particulate fractions that behave differently. 426 Examining the rainfall/climate is also important to determine the impacts of climate on radiocesium 427 transfer and storage. Importantly the potential storage of radiocesium between hillslopes and the 428 Pacific Ocean may become a significant focal point of future management efforts. Prior to examining research that quantifies radiocesium transfers and export from the coastal catchments, each of these 429 430 dynamics impacting radiocesium transfers will be discussed.

In the fallout affected catchments, the main radioactive pollution plume is drained by the Abukuma River, along with several small coastal catchments (Fig. 3). The majority of research has focused on understanding the transfer of fallout contamination in the Abukuma River basin (Fig. 8). For the smaller coastal catchments, there has been a focus of research in the Niida and Mano rivers with decreasing research in the catchments close to and south of the FDNPP. In particular there has been limited research published on the catchments with the highest mean radiocesium inventories (the Maeda, Kuma, Ukedo and Ota catchments).



Figure 8. Location of stream sampling locations for samples examining dissolved/particulate (DP) fractions and their distribution coefficient (KD), particle size (PS), monitoring (M), and lag deposits (LD).

438 **4.1** Radiocesium distribution coefficient (dissolved/particulate)

- 439 To quantify radiocesium migration in catchments, it is necessary to understand the ratio of
- 440 particulate and dissolved radiocesium transported in river systems. Dissolved and particulate-bound
- radiocesium migrate differently in waterways and their ratio provides fundamental information for

understanding future migration (Yoshimura et al., 2015b). The dissolved/suspended partitioning of
radionuclides is typically described with a distribution coefficient, K_d, and is expressed as a ratio of
radiocesium activity concentration in particulate and dissolved fractions under equilibrium conditions
(Ueda et al., 2013).

446 In samples from the Abukuma River, Sakaguchi et al. (2015) reported that the particulate fraction transported 80% of ¹³⁷Cs (Table 3). Conversely, Tsuji et al. (2014) indicated that in the Abukuma and 447 Ota catchments, only 64% of the radiocesium was transported with the particulate fraction. When 448 examining sediments across the fallout impacted region, Yoshimura et al. (2015b) found that 449 450 between 58-100% of radiocesium was transported with the particulate fraction, with the exclusion of 451 one estuary site. Similarly, Yoshikawa et al. (2014) reported that between 60-95% of radiocesium was 452 transported with the particulate fraction in three small subcatchments (<0.1 km²) of the Abukuma 453 River.

454 Differences between these results are likely related to the timing of sampling in relation to significant 455 rainfall events. For example, in the Same and Natsui catchments, Nagao et al. (2013) observed that in normal flow conditions, the particulate fraction transported 40% of the ¹³⁷Cs, compared to ~100% 456 457 during a significant rainfall event (i.e. Typhoon Roke). Ueda et al. (2013) also reported significant 458 differences between flood and base flow conditions with the particulate partition contributing over 459 90% of radiocesium in the Niida River during Typhoon Roke. In this catchment, the particulate 460 fraction only accounted for 40% of the radiocesium migration in base flow conditions. In addition, 461 Shinomiya et al. (2014) found that 92-97% of radiocesium is transferred with the particulate fraction 462 during significant rainfall events (Typhoon Guchol), compared to between 12-96% in base flow. From modelling research, Mori et al. (2014), determined that the particulate fraction transported more 463 than 95% of the ¹³⁷Cs, slightly above the range calculated by Yamashiki et al. (2014) (82-93%). This 464 465 research clearly demonstrates that the particulate fraction dominates the transfer of radiocesium in the Fukushima region during major rainfall events. 466

- 467 Differences between the dissolved/particulate radiocesium partitioning may also be influenced by
- the time elapsed since the FDNPP accident. For example, in the Niida catchment, Nagao et al. (2014)

Author(s)	Catchment(s)	Sampling Period	Radionuclide(s)	Particulate Fraction of transported ¹³⁷ Cs (%)	Average K(d)
Nagao et al. (2013)	Same / Natsui	July 2011 - Dec. 2011	¹³⁴ Cs & ¹³⁷ Cs	99 Flood Conditions 33 Base Flow	2.5 x 10 ⁶
Nagao et al. (2014)	Niida	May 2011 - Nov. 2012	¹³⁴ Cs & ¹³⁷ Cs	47 (Before Sept. 2011) 84 (After Sept. 2011)	N/A
Sakaguchi et al. (2015)	Abukuma	June 2011 - Dec. 2012	¹³⁷ Cs	80	8.8 x 10 ⁵
Shinomiya et al. (2014)	Experimental Catchment (Abukuma)	June 2012	¹³⁴ Cs, ¹³⁷ Cs	92-97 Flood Conditions 12-96 Base Flow	N/A
Tsuji et al. (2014)	Abukuma, Ota	Sept. 2012 - May 2013	¹³⁴ Cs & ¹³⁷ Cs	64	N/A
Ueda et al. (2013)	Niida	July 2011 - Nov. 2011	¹³⁴ Cs & ¹³⁷ Cs	>90 Flood Conditions 40 Base Flow	3.5 x 10⁵
Yamashiki et al. (2014)	Abukuma	June 2011 - May 2012	¹³⁴ Cs & ¹³⁷ Cs	82-93	N/A
Yoshikawa et al. (2014)	Abukuma	April 2012 – Sept. 2012	¹³⁷ Cs	60–83 Flood Conditions 69–95 Base Flow	1.2 x 10 ⁵ Base 10.6 x 10 ⁵ Event
Yoshimura et al. (2015b)	Abukuma, Mano, Same, Fujiwara, Niida, Ota, Odaka, Asami, Ukedo,	Dec. 2012	¹³⁴ Cs & ¹³⁷ Cs	58-100 ^ª	3.6 × 10⁵

Table 3. Summary of research examining radiocesium in particulate/dissolved and K_d

^aexcluding an estuary site.

reported that the particulate fraction averaged 47% between May and September 2011 compared to 469 84% between December 2011 and November 2012. The differences were attributed to rainfall and 470 471 declining availability of leachable sources of radiocesium that were likely washed off in greater 472 concentrations early after the accident. When researching dissolved radiocesium fractions in stream 473 and ground water, Iwagami et al. (Submitted) also reported a fast flush of radiocesium before October 2011, resulting from the initial wash-off, followed by a slow decline in radiocesium 474 475 concentrations. Moving forward, it is likely that the particulate phase will constitute a greater 476 proportion of the radiocesium transfer.

In the Niida catchment, Ueda et al. (2013) reported that the mean K_d was 3.5 × 10⁵ (Table 3). 477 Yoshimura et al. (2015b) calculated a very similar mean K_d of 3.6 x 10^5 when comparing the 478 479 particulate material sampled from time-integrated samplers (Phillips et al., 2000) and water sampled with a vacuum pump sampler. Nagao et al. (2013) demonstrated that the K_d range in the Same and 480 Natsui Rivers was between $0.4 \times 10^6 - 5.0 \times 10^6$. Yoshikawa et al. (2014) reported that the K_d in three 481 small (<1.2 km²) was 1.2×10^5 in base flow compared to 10.6×10^5 in event conditions. Sakaguchi et 482 al. (2015) indicated the K_d in the Abukuma River was 8.8 x 10⁵. In general, these results are an order 483 of magnitude greater than those reported by the International Atomic Energy Agency after the 484 485 Chernobyl accident (2.9 x 10⁴) (UNSCEAR, 2000) indicative of a greater percentage of particulate-486 bound radiocesium migrating through these catchments compared to those sampled in areas impacted by fallout from the Chernobyl accident. In Ukrainian rivers with similar volumes of total 487 488 suspended solids as in Fukushima rivers, the dissolved fraction was found to contain ~75% of the sampled ¹³⁷Cs (1-2 y post-accident) (Sansone et al., 1996). Although a significant component of 489 490 radiocesium migrates with the dissolved fraction in normal flow conditions in the Fukushima region, 491 it contributes only a limited amount to the total radiocesium migration to downstream areas 492 (Kinouchi et al., 2015). The challenge is that the particulate radiocesium concentrations are often 493 irregular and driven by suspended solid fluctuations resulting from heavy precipitation (Tsuji et al.,

2014) and it is difficult and costly to continuously monitor suspended solids, particularly during highflow conditions.

496 Ueda et al. (2013) demonstrated a high correlation between the discharge of radiocesium and fluvial 497 water discharge. Correlations were also found between particulate/dissolved radiocesium 498 concentrations and catchment radiocesium inventories, suggesting that it may be possible to predict 499 radiocesium concentrations (Yoshimura et al., 2015b). As this correlation is empirical and may 500 change owing to regional hydro-meteorological dynamics, further research is required before using this relationship directly in modelling approaches. Importantly, K_d was found to be related to the 501 502 proportion of large particles (e.g. low clay content), the clay mineralogy of small particles, and the 503 aqueous phase cation concentration (Fan et al., 2014a). As radiocesium primarily bound to fine 504 sediment particles (Yamaguchi et al., 2012), it is important to understand the relationship between 505 radiocesium and particle size for modelling radiocesium transfer at larger scales.

506 4.2 Radiocesium particle size distributions

The fine soil fraction, to which radiocesium is bound, is preferentially eroded at the hillslope scale 507 508 (mean particle size = 25.5 μ m) (Yoshimura et al., 2015a). This fraction (e.g. silt and clay) is also 509 preferentially transported farther distances owing to a smaller settling velocity (Iwasaki et al., 2015) 510 resulting in the accumulation of radiocesium in downstream areas potentially unaffected by the 511 initial fallout (Tanaka et al., 2014). Indeed the distribution of radiocesium with regards to particle size 512 will have a significant impact on the radiocesium deposition in downstream areas. It is therefore 513 important to examine whether sediment generation and transport processes impact the radiocesium 514 fixation to fine particles.

515 River sediments were found to have a high ¹³⁷Cs adsorption capacity (Fan et al., 2014a). Multiple 516 studies demonstrated that radiocesium concentrations were elevated in the fine particle size 517 fractions (Fan et al., 2014a; Sakaguchi et al., 2015; Tanaka et al., 2015; Tanaka et al., 2014). Sakaguchi 518 et al. (2015) found that the maximum ¹³⁷Cs concentration in suspended sediment was found in the silt particle size fraction (3-63 μ m). This result was also reported in Fan et al. (2014a). In the former case, ¹³⁷Cs in each fraction did not demonstrate any systematic variation across multiple sampling sites (Sakaguchi et al., 2015). In the latter, Fan et al. (2014a) found that there was a weak relationship between particle size and radiocesium for fine sediment (<63 μ m), though a stronger relationship was derived for the coarser grain fractions (e.g. >63 μ m).

524 The sorption of radiocesium to sediments was found to be directly related to their physical-chemical 525 properties, including cation exchange capacity, organic matter content and mineralogy (Fan et al., 526 2014a). What was unexpected with these findings was that the radiocesium concentrations were not 527 always enriched in the fine clay material (<3 μ m). A below expected quantity of radiocesium in the 528 fine clay material was also reported by Tanaka et al. (2014) who found that the silt and sand fractions 529 transported 95% of the sediment-sorbed radiocesium. This study did find an enrichment of ¹³⁷Cs in 530 the clay material, though it was found to only constitute a small proportion of the sampled bedload 531 material. Fan et al. (2014a) also reported a low clay mineral content in the <2 μ m particle size 532 fraction. The lack of radiocesium enrichment in the finest particle size fraction may be the result of 533 sampling technique (e.g. sampling bed load material) or other fluvial processes that may result in the 534 majority of very fine material being exported from the catchment.

For modelling and management of radiocesium transfers in the contaminated catchments, it may be silt, rather than clay that requires emphasis. For example, silt was found to be the significant source of bed contamination (excluding initial fallout) whereas during later stages of rainfall flow events, clay materials were found to be the dominant supply of ¹³⁷Cs to the lower river reaches (Kurikami et al., 2014). Further characterization of the relationship between radiocesium and particle size is fundamental for understanding radiocesium transfer and storage in fallout affected catchments.

541 **4.3 Radiocesium mobilization: Rainfall regime**

Rainfall is the driving mechanism behind soil erosion by water (Nearing et al., 1989; Wischmeier and
Smith, 1958). The regional climate in conjunction with the topography of the fallout impacted region

544 exacerbates radiocesium transfer. The catchments of the Fukushima Prefecture are mountainous and characterized with a steep topography (Kinouchi et al., 2015; Minoura et al., 2014). The 545 546 temperate/monsoonal climate results in high seasonal and annual precipitation including significant 547 runoff events generated from snowmelt and intense rainfall from typhoons (Evrard et al., 2013; 548 Yamashiki et al., 2014). In the Fukushima region, the annual precipitation is three-fold greater than in 549 Chernobyl, resulting in a more rapid radiocesium washoff (Tsuji et al., 2014). This climate is one of 550 the fundamental mechanisms driving the migration and transfer of radiocesium throughout the 551 catchments that received the original fallout (Evrard et al., 2013).

552 Evrard et al. (2013) demonstrated that the region is highly reactive to both rainfall and snow fall 553 events. The authors demonstrated that there is a significant export of contaminated material from 554 coastal rivers to the Pacific Ocean. Chartin et al. (2013) found that the combination of rainfall and 555 snowmelt events results in a progressive migration of sediment bound radiocesium to the Pacific 556 Ocean. This sediment-based sampling clearly demonstrated that there has been a significant 557 dispersion of contaminated sediments in the 20 months after the FDNPP accident (Chartin et al., 558 2013). These authors highlight a massive, rapid and episodic transfer of radiocesium resulting from 559 the climate in general and the snowmelt/typhoon cycle in particular. Yamashiki et al. (2014) also 560 reported two significant runoff peaks in their research, one clearly coinciding with the snowmelt in 561 late spring (April-May) and the other coinciding with typhoons.

562 Overall, the literature clearly demonstrates that the majority of soil erosion and concomitant 563 radiocesium migration, coincides with extreme rainfall-runoff events (Minoura et al., 2014; Mouri et 564 al., 2014; Nagao et al., 2013; Onishi et al., 2014). In fact, most sediment and radiocesium migration is 565 thought to occur during short flood events that may occur 1-2 times per year (Yamaguchi et al., 566 2014). This importance of flooding and rainfall was incorporated by Kitamura et al. (2014) who only 567 modeled sediment migration when the flow rate is 10 times greater than the average annual flow.

568 In particular, the typhoons in the fallout impacted region transfer massive amounts of sediment and 569 radiocesium. For example, Typhoon Roke (Sept. 2011) destroyed paddy banks (Wakahara et al., 570 2014) and resulted in mud depositions that were ~10 cm thick (Minoura et al., 2014). Yamashiki et al. (2014) estimated that Typhoon Roke mobilized and transferred 6.18 TBq, or 61% of the total 571 572 radiocesium load to the coastal areas during their observation period. Kinouchi et al. (2015) reported 573 that three major flow events were responsible for the migration of 43% of the sediment between 574 June 2011 and Dec. 2012. Nagao et al. (2013) demonstrated that radiocesium concentrations were 575 an order of magnitude higher during high flow conditions after Typhoon Roke, which also resulted in 576 particulate transport of radiocesium being 100% compared to 21-56% in baseflows. The authors 577 concluded that these heavy rainfall events are one of the important drivers of radiocesium migration 578 from catchments to the coast resulting for between 30-50% of the annual radiocesium coastal 579 transfer. Ultimately, extreme rainfall events are fundamental to the migration of sediments and 580 radiocesium from the hillslopes to the Pacific Ocean.

581 4.4 Radiocesium storage: Reservoirs and dams

582 Not all mobilized sediment and radiocesium is exported to the Pacific Ocean. The migration of 583 sediment, particularly coarse particle size fractions, occurs over longer temporal scales. Further, 584 anthropogenic alterations to riverine morphology, such as reservoirs and dams, further delay the 585 migration and transfer of sediment and radiocesium downstream (Kurikami et al., 2014). Reservoirs 586 not only delay the migration of radiocesium and sediments, they also may act as a major sinks and 587 storage (Evrard et al., 2014a; Evrard et al., 2013) as only a select fraction of radiocesium is 588 transferred downstream, predominantly the dissolved or clay-bound fractions (Onishi et al., 2014). 589 These reservoirs may also act as sources of contaminated sediment during water releases.

590 Chartin et al. (2013) reported a significant layer of sediments with high radiocesium concentrations 591 deposited behind a reservoir in the upper Ota River. Kitamura et al. (2014) indicated that 10^4 t of 592 sediment and 10 TBq of ¹³⁷Cs were transferred into the Ogaki Dam on the Ukedo River which 593 buffered the downstream transfer of radiocesium. Kurikami et al. (2014) reported that the sand 594 particle size fraction was completely contained within the Ogaki Dam, with only 5% of the silt and 595 50% of the clay fractions transferred downstream. Similarly, Yamada et al. (2015) found that 596 practically all of the silt and sand are deposited within the Ogaki Dam reservoir, with the sand being 597 deposited near the entrance of the reservoir and the silt being more uniformly distributed. 598 Approximately a third of the clay fraction was modelled to be exported downstream from this 599 reservoir (Yamada et al., 2015).

600 Management of the dams and reservoirs may have a significant impact on radiocesium migration 601 downstream. Mouri et al. (2014), Yamada et al. (2015) and Kurikami et al. (2013) demonstrated that 602 reservoir management could assist in the control of radiocesium fluxes. The challenge with the 603 management of these reservoirs in the temperate/monsoonal climate in the Fukushima region is that 604 occasional water releases are required for their safe operation. Evrard et al. (2014a) indicated these 605 dam releases could enhance the natural migration of radiocesium downstream and may become a 606 significant factor controlling the downstream transfer of sediment and radiocesium in these 607 anthropogenically modified catchments. In the Fukushima Prefecture, there are 10 significant dams 608 and over a 1,000 reservoirs for both agricultural and surface water management (Yamada et al., 609 2015). Understanding the radiocesium deposition and migration within these water storages is 610 important for understanding potential contamination of downstream regions, particularly cultivated 611 rice paddy systems.

To quantify radiocesium transfer from hillslopes to the Pacific Ocean, it is necessary to characterize riverine radiocesium dynamics within the climatic context. Radiocesium may be transferred in particulate or dissolved fractions, and the particle size distribution of particulate radiocesium will either facilitate its transfer or potential downstream storage during rainfall and snowmelt runoff events that drive the majority of radiocesium migration. Understanding these dynamics, particularly in the catchments with the highest mean radiocesium activities will be of increasing importance

618 moving forward. In particular, characterizing interrelationships between typhoons and other 619 rainfall/runoff events along with radiocesium particle size and distribution coefficients allows 620 researchers to determine whether radiocesium may be stored within a catchment after its hillslope 621 mobilization or whether it is ultimately transferred to the Pacific Ocean.

622

4. Quantifying radiocesium transfers from hillslopes to the Pacific Ocean

623 Radiocesium dynamics within riverine environments are inherently complex. Accordingly, there are a 624 variety of factors governing radiocesium transfer in the region affected by the FDNPP accident. 625 Quantifying radiocesium transfers requires in-stream sampling and monitoring stations or complex 626 desktop sediment modelling. Publications reviewed for this section focus on catchment scale 627 radiocesium transfers and will be subdivided into monitoring and modelling research. Table 4 lists 628 these two approaches to quantifying radiocesium transfers in the Fukushima coastal catchments 629 while Fig. 8 includes the locations of the monitoring stations. Similarly to the research on the factors 630 influencing riverine radiocesium transfers above, the majority of modelling research occurs within 631 the Abukuma River catchment. There is limited monitoring and modelling research in the smaller 632 coastal catchments, particularly the catchments to the south and in close proximity to the FDNPP.

633 **5.1 Monitoring radiocesium catchment scale transfers**

Direct monitoring of radiocesium sampled within rivers over several months or years provides data required to quantify radiocesium exports from catchments. For example, relationships between total suspended solid and radiocesium quantities within river water/sediment samples may be quantified and extrapolated over longer temporal periods. Although there are limitations, this monitoring

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					Radionuclides		Catchment
			Area		(Particulate (P)-		Inventory
Author(s)	Туре	Catchment(s)	(km²)	Period	Dissolved (D))	Export (TBq)	Exported (%)
Iwasaki et al. (2015)	Modelling	Abukuma (downstream reach)	31	September 2011	¹³⁷ Cs (P)	3.29	n/a
Kinouchi et al. (2015)	Modelling	Abukuma (Kuchibuto subcatchment)	140	June 2011 - December 2012	¹³⁷ Cs (P)	0.274	0.8
Kitamura et al. (2014)	Modelling	Abukuma, Ukedo, Niida, Maeda, Kuma, Ota, Mano, Kido, Odaka, Tomioka. Natsui. Same. Ide. Uda	8,352	Annually for the initial years after the accident	¹³⁷ Cs (P)	8.4	1.0
Kurikami et al. (2014)	Modelling	Ukedo (Okagi Dam)	n/a	September 2013	¹³⁷ Cs (D&P)	0.86 ^ª	n/a
Mori et al. (2014)	Modelling	Un-named reservoir	15	March 2011 - December 2013	¹³⁷ Cs (P)	n/a	<5.0
Mouri et al. (2014)	Modelling	Tone (Kusaki Dam)	254	2010 - 2090	¹³⁷ Cs (P)	900 Bq/kg ^b	n/a
Nagao et al. (2013)	Monitoring	Natsui, Same	749/600	March 2011 - December. 2011	¹³⁴ Cs, ¹³⁷ Cs (D&P)	0.04	n/a
Pratama et al. (2015)	Modelling	Abukuma	5,172	August 2011-May 2012 + future	¹³⁴ Cs' ¹³⁷ Cs (D&P)	~12	2%
Shinomiya et al. (2014)	Monitoring	Experimental Catchment (Abukuma)	0.01	scenarios June 2012	¹³⁴ Cs, ¹³⁷ Cs	115 (Bq m ²)	0.07
Ueda et al. (2013)	Monitoring	Niida (Wasriki and Hiso Rivers)	< 5	July-November 2011	¹³⁴ Cs, ¹³⁷ Cs (D&P)	0.03	0.4
Yamada et al. (2015)	Modelling	Ukedo (Okagi Dam)	n/a	September 2013 & average flood	¹³⁷ Cs (D&P)	0.07 & 0.35	n/a
Yamaguchi et al. (2014)	Modelling	Abukuma, Ukedo, Niida, Maeda, Kuma, Ota, Mano, Kido, Odaka, Tomioka, Natsui, Same, Ida, Uda	8,352	March 2011 - March 2012 + future scenarios	¹³⁷ Cs (P)	02.3 - 6.5 ^c	0.5 - 7.5 ^c
Yamashiki et al. (2014)	Monitoring	Abukuma	5,172	June 2011 - May 2012	¹³⁴ Cs, ¹³⁷ Cs (P)	10.08	1.1

^adeposited in the Okagi Dam ^bcurrent annual deposition in the Kusaki Dam modelled to be ~900 Bq/kg

 $^{\rm e}$ minimum and maximum values modelled over 30 y

approach is a direct method to quantify radiocesium transfer from catchments and ultimately to thePacific Ocean.

641 Most of the direct radiocesium monitoring within catchments occurs at the smaller subcatchment 642 scale. For example, Shinomiya et al. (2014) monitored radiocesium discharge from a small forested catchment (0.012km²) in Abukuma River basin. The authors found that maximum rainfall intensity 643 and runoff coincided with the maximum radiocesium (3.9 Bq L⁻¹) and suspended solid export. During 644 the receding limb of the hydrograph, radiocesium transfers concentrations dropped sharply. In total, 645 115 Bq m⁻² was exported from this experimental catchment or 0.07% of the initial radiocesium 646 647 inventory, indicative of a negligible radiocesium transfer from this forested catchment during a 648 typhoon event.

Ueda et al. (2013) conducted regression analyses to relate radiocesium to river discharge at two sampling locations in the upper Niida catchment between March 15 and December 31, 2011. They estimated that 0.045 TBq and 0.020 TBq were exported from the Hiso and Wariki Rivers catchments, respectively. This equated to 0.5% of the radiocesium inventory of the Hiso catchment and 0.3% of the Wariki River catchment. Although each catchment had different land uses, their proportional radiocesium export were similar.

Nagao et al. (2013) calculated radiocesium export from the Natsui and Same Rivers catchments, south of the FDNPP. They quantified radiocesium flux by multiplying mean monthly discharge by the ¹³⁷Cs concentrations and time. They determined that 0.003 and 0.005 TBq y⁻¹ of ¹³⁷Cs were exported from the Natsui and Same Rivers, respectively. They also indicated that 30-50% of their estimated 10 month ¹³⁷Cs export occurred during Typhoon Roke.

To our knowledge, there has only been one current study published quantifying the transfer of radiocesium from a large catchment to the Pacific Ocean. Yamashiki et al. (2014) used in-stream samplers to collect suspended sediment along with turbidity meters and pressure transducers to calculate suspended sediment fluxes from the Abukuma River catchment to the Pacific Ocean. The authors estimated that ~10 TBq of radiocesium was exported to the Pacific between June 2011 and May 2012. This amounted to ~1% initial estimated catchment inventory. This direct transfer of sediment from hillslopes through the Abukuma catchment to the Pacific Ocean was within the same order of magnitude of the direct leakage from the FDNPP to the ocean on August 21, 2013 (24 TBq) (Yamashiki et al., 2014) clearly demonstrating the importance of quantifying the riverine transfer of radiocesium to the Pacific Ocean.

670 **5.2 Modelling radiocesium catchment scale transfers**

671 Quantifying catchment scale radiocesium transfers with modelling approaches requires a variety of 672 model inputs along with calibration and validation data (often obtained from direct river monitoring). 673 The challenge for researchers when quantifying catchment radiocesium fluxes is incorporating the 674 fundamental radiocesium transfer dynamics required to populate models without over 675 parameterizing radiocesium dynamics. Often, the scale of the research catchment controls the 676 breadth of radiocesium dynamics incorporated into individual models. Similarly to the monitoring 677 studies, desktop models have been performed at various scales, from upstream catchments to large 678 river basins to quantify radiocesium transfer directly to the Pacific Ocean.

679 For an upstream catchment of the Abukuma River system, Kinouchi et al. (2015) developed a 680 complex model incorporating evaporation, evapotranspiration, infiltration, surface and subsurface 681 runoff along with groundwater, overland and river flow. The authors coupled hydrological processes and energy transfer within a catchment along with interactions between surface water and 682 683 groundwater. Importantly, they modelled the hydrological processes for paddy fields, including 684 irrigation, ponding, drainage and seepage to nearby channels. Their model also incorporated season 685 variation in vegetation. The authors reported that after 30 y, the catchment inventory of ¹³⁷Cs will be reduced to 39% of the initial inventory (Kinouchi et al., 2015). They reported that between 686 September 2011 and December 2012, 0.274 TBq of ¹³⁷Cs was exported from the Kuchibuto River 687 catchment or less than 1% of the initial ¹³⁷Cs fallout. 688

689 Several studies modelled sediment and radiocesium deposition in reservoirs of the region. Mouri et 690 al. (2014) examined radiocesium inputs to the Kusaki Dam reservoir on the edge of the fallout 691 affected region. The authors incorporated mixed-particle size distributions into a sediment transport 692 equation to determine sedimentation and concomitant ¹³⁷Cs deposition in the reservoir. The authors 693 included 10 different particle size fractions into their model. They reported that the wash load 694 (particle size ≤0.2mm) contributed little to reservoir siltation and predicted that 3.4 Mt/y of sediment were deposited in the catchment with 900 Bq/kg of ¹³⁷Cs. For the future, they predicted that 695 sediment deposition would increase to ~3.65 Mt/y with ¹³⁷Cs activities of 120 Bq/kg. 696

Mori et al. (2014) modelled ¹³⁷Cs dynamics in an unnamed reservoir catchment between March 2011 697 698 and December 2013. In particular, they developed an integrated model incorporating fully coupled 699 radionuclide behaviour in a channel, overland and subsurface three dimensional (3D), physics-based 700 model. The authors demonstrated that the majority of radiocesium was transported with the 701 particulate fraction (>95%). Their 3D model indicated that radiocesium was predominantly eroded 702 from steep hillslopes in close proximity to the stream network with the majority of deposition occurring at the entrance of the reservoir. The authors determined that the total ¹³⁷Cs discharge with 703 704 both sediment and dissolved fractions was <5% of the initial fallout inventory indicating the majority 705 of the fallout radiocesium remained in the dammed catchment.

706 Kitamura et al. (2014) reported the results from the application of the Time-dependent One-707 dimensional Degradation and Migration (TODAM) model in the Ukedo Catchment. This model is also 708 outlined in Kurikami et al. (2014) and discussed in Kitamura et al. (2015). The TODAM model 709 incorporates the deposition and re-suspension of sediment along with the potential 710 desorption/adsorption of radiocesium within sub-models. It simulates the movement of three sediment particle size fractions (sand/silt/clay). The authors determined that ¹³⁷Cs concentrations 711 were lower than the drinking water limit after approximately 4 d following a significant rainfall event 712 (Kurikami et al., 2014). The authors also reported that 0.86 TBq of ¹³⁷Cs was deposited in the Okagi 713

Dam Reservoir (Kitamura et al., 2014), which is only one order of magnitude less than the total ¹³⁷Cs
 export to the Pacific Ocean from 14 coastal catchments.

Yamada et al. (2015) also modelled radiocesium dynamics for the Ogaki Dam with the Nay2D model (Shimizu, 2003) to examine the influence of countermeasures such as controlling the reservoir volume on ¹³⁷Cs export. This paper was also discussed in Kitamura et al. (2015). The authors compared a September 2013 flood event to the long-term average flood event and calculated ¹³⁷Cs deposition and export. For the September 2013 flood, they modelled that 0.16 TBq was deposited in the Ogaki Dam with 0.07 TBq exported downstream. These transfers were low compared to the average yearly flood which was modelled to deposit 0.55 TBq and export 0.35 TBq.

723 Finally, several studies quantified the export of radiocesium from catchments to the Pacific Ocean by focusing on lower river reaches or by modelling major catchments affected by fallout. Iwasaki et al. 724 (2015) investigated ¹³⁷Cs dynamics in a downstream reach of the Abukuma River. They incorporated 725 726 a Rouse number threshold within their complex model to determine whether sediments are 727 exported from the reach or deposited. The authors reported there is significant deposition of 728 contaminated sediments on river floodplains. The authors highlighted the potential of using a Rouse 729 number threshold to determine the location of radiocesium deposition within the catchment to 730 target decontamination activities. In particular, they suggested that the floodplains and inner areas of river bends can be contaminated at the end of significant rainfall events. Even whilst incorporating 731 deposition within the 31 km reach, the authors reported that 98% of ¹³⁷Cs (or 3.2 TBq) was 732 733 transferred downstream towards the Pacific Ocean.

For the entire Abukuma River catchment, Pratama et al. (2015) developed a multiple compartmental model that incorporated the complexity of radiocesium export from forest, urban and agricultural land uses. The authors' model accurately predicted seasonal patterns of radiocesium fluxes in the Abukuma catchment. The model was calibrated with monitoring data between August 2011 and May 2012 and then forecasted radiocesium migrations for the next 100 y. The authors found that in total

739 155 TBq of radiocesium will be exported from the Abukuma catchment over the next century. Based 740 on their results, approximately 12 TBq of radiocesium was modelled to be exported from the 741 Abukuma basin in the first year after the accident representing on average 2% of the initial 742 catchment inventory.

Kitamura et al. (2014) outlined an USLE-based model to estimate the ¹³⁷Cs discharge from 14 743 744 catchments that received fallout which are all outlined in Fig. 3. In particular, they focused on land-745 use as they argued it may have the greatest potential impact on soil erosion. This model is also 746 described in Yamaguchi et al. (2014) and discussed in Kitamura et al. (2015). These authors reported that 8.4 TBq of ¹³⁷Cs will be exported to the Pacific Ocean in the immediate years following the 747 748 FDNPP accident (Kitamura et al., 2014). This transfer of radiocesium to the Pacific Ocean can be 749 compared to estimates of direct radiocesium discharge from the FDNPP between June 2011 and 750 September 2012 (17 TBq), and a leakage event in August 2013 (24 TBq)(Kitamura et al., 2014; 751 Yamaguchi et al., 2014). The model also estimates that only 35% of eroded soils are transported to 752 the Pacific Ocean along with less than 1% of the initial radiocesium catchment inventory (Kitamura et 753 al., 2014).

754 **5.3 Monitoring and modelling summary**

755 A convenient approach to examine radiocesium transfer results is to compare the percentage of 756 initial radiocesium catchment inventory exported or remaining. At the plot or hillslope scale, Yoshimura et al. (2015a) determined that between 0.07 and 3.2% of the catchment ¹³⁷Cs inventory 757 758 was exported with wash-off, with a mean annual export of 0.7% for their 7 USLE plots. For a forested 759 experimental catchment, Shinomiya et al. (2014) reported that only 0.07% of the initial radiocesium 760 inventory was exported during a typhoon event. In an outline of Fukushima related research projects, Saito and Onda (2015) indicated that less than 1% of the radiocesium inventory is 761 762 transferred from hillslopes and discharged from catchments.

763 This annual wash-off rate of less than 1% is similar to the catchment modelling results obtained at the subcatchment and catchment scales. In the Kuchibuto subcatchment of the Abukuma River, 764 765 Kinouchi et al. (2015) reported that 0.8% of the initial fallout (or 0.28 TBq) was exported from the 766 catchment. For the entire Abukuma catchment, Yamashiki et al. (2014) found that 1.1% of the 767 radiocesium inventory was exported to the Pacific Ocean compared to an average of 2% reported by 768 Pratama et al. (2015). In the upper Niida catchment, Ueda et al. (2013) reported that 0.3 to 0.5% of 769 the initial catchment radiocesium fallout was exported. For all 14 coastal catchments, Kitamura et al. 770 (2014) reported 99% of the catchment inventory remained with only 1% of the original fallout being exported. The highest ¹³⁷Cs export compared to catchment inventory of <5% was modelled by Mori 771 772 et al. (2014). Although there is a significant amount of radiocesium exported from hillslopes and 773 catchments, this amount represents only a small fraction of the initial fallout inventory. Importantly, 774 comparing research results through using the percentages of catchment inventories remaining and 775 exported normalizes the results for comparison by removing the influences of catchment area, scale 776 of research, or even the simply different approaches to quantifying catchment inventories.

777 Although they did not estimate discharge, Chartin et al. (2013) clearly demonstrated that the 778 catchments that received fallout will become a perennial supply of radiocesium to the Pacific Ocean. 779 This was confirmed by the modelling research reviewed above with significant volumes of 780 radiocesium modelled to be exported annually from the fallout-affected catchments to the Pacific 781 Ocean. In total, for the entire region, this amount was estimated to be 6.9-10 TBq by both sample-782 based (Yamashiki et al., 2014) and desktop modelling (Kitamura et al., 2014; Yamaguchi et al., 2014). Iwasaki et al. (2015) modelled that 3.3 TBq of the ¹³⁷Cs export migrated through the lower reaches of 783 the Abukuma River alone during Typhoon Roke. After the Abukuma River, (Kitamura et al., 2014) 784 785 found that the Ukedo and Niida River catchments were the only catchments to export >1 TBq of 786 ¹³⁷Cs. Indeed, these coastal catchments will be significant long-term source of radiocesium to the 787 Pacific Ocean.

788 Confirming the terrestrial origin of radionuclides detected in sediment collected in the Pacific Ocean 789 is particularly difficult. Several origins are possible: previous deposits associated with atmospheric 790 bomb testing in the 1960s; direct releases from FDNPP to the ocean; and exports from coastal rivers 791 of Eastern Japan. In order to quantify the respective contributions of these sources, the 792 measurement of additional radionuclides in marine sediment may provide indications on their origin. 793 For instance, a study showed that material exported by rivers draining the Fukushima radioactive pollution plume was characterised by a different ²³⁹Pu/²⁴¹Pu atom ratio than the material originating 794 795 from areas of Japan devoid of Fukushima fallout (Evrard et al., 2014b). Alternatively, the detection of the shorter-lived ¹³⁴Cs isotope in marine sediment indicates the presence of Fukushima fallout at this 796 location, as ¹³⁴Cs fallout originating from the 1960s atmospheric bomb fallout fully decayed (Bu et al., 797 798 2013).

799 5. Management and research implications

Characterizing the transfer of sediment and particle-borne radiocesium is required to predict future radiocesium dispersion in the Fukushima Prefecture and design and implement effective countermeasures. Reviewing the processes and dynamics governing radiocesium transfer from hillslopes to the Pacific Ocean provides not only a comprehensive depiction of the fundamental factors driving radiocesium transfer, it also highlights future potential radiocesium management challenges on one hand, and future research perspectives on the other.

In the fallout-affected areas, there has been a massive decontamination effort. Thousands of workers are decontaminating rice paddy fields, grasslands, and entire villages. For soil decontamination, the vegetation and litter are first removed followed by the removal of the topsoil. The impact of this decontamination on sediment modelling and radiocesium transfer research is unknown. The decontamination may temporarily increase the supply of radiocesium as the significant devegetation/soil management interventions is likely to result in an increase in the supply of sediments to riverine systems, as reported by Evrard et al. (2014a) and Lepage et al. (2014a) following

preliminary decontamination works in the Upper Niida River catchment. The question is whether the elevated suspended sediment loads mobilized by the decontamination will have elevated radiocesium quantities, be equivalent to pre-decontaminated levels, or be lowered through dilution with non-contaminated sediment? Further, will the decontamination efforts reduce the transfer of radiocesium from the hillslopes to the Pacific Ocean?

818 Examining the remaining catchment inventory from multiple studies provides a direct method to 819 compare multi-scale research in the region. Although a significant quantity of radiocesium is 820 exported from catchments to the Pacific Ocean, this amount represents only a small fraction of the 821 initial catchment inventory (<5%). Therefore, quantifying sediment (dis)connectivity is fundamental 822 to not only understanding radiocesium migration in these catchments though also for highlighting 823 potential areas that may accumulate and store radiocesium. The coastal plains, that remain 824 inhabited, potentially serve as a large buffer zone which could temporarily store radiocesium 825 transferred from the upstream forested, mountainous piedmont range. Floodplains and riparian 826 zones that were not directly impacted by the FDNPP accident may receive significant pulses of 827 radiocesium during flooding events that transfer contaminated sediment from upstream areas. 828 Accordingly, it is important to determine whether and where are any important storages of 829 radiocesium that need to be managed appropriately. The potential incorporation of Rouse number 830 threshold may help isolate locations of radiocesium deposition within the catchment to target 831 decontamination activities. After major typhoons, it may be necessary to analyse deposited material 832 in populated areas for radioactive contamination.

Perhaps the most critical potential sediment and radiocesium sink are rice paddy fields. Rice paddy fields, depending on the time of year and hydrological environment, may be both sources and sinks of radiocesium. The fertile soils of the Fukushima region make it important agriculturally. In particular, paddy fields occupy approximately 70% of all farming area in the region (Wakahara et al., 2014). Unfortunately, radiocesium may transfer into rice. Contamination of agricultural products

838 mainly occurs by direct absorption (from fallout or irrigation) or by root uptake resulting in long-term 839 contamination of crops. Contamination during irrigation is problematic as radiocesium adsorption 840 from water is more direct than adsorption from soil and many fields are irrigated by rivers 841 contaminated with dissolved and particulate bound radiocesium (Nemoto and Abe, 2013). Transfer 842 factors (TF) from soil to plant depend on vegetation properties. Endo et al. (2013) determined the TF 843 for different parts of the rice (Table 5) and concluded that contamination is principally concentrated 844 in the roots. Furthermore, the TF was significantly higher in samples collected closer to the paddy 845 fields water inlet (Yoshikawa et al., 2014). Quantifying the radiocesium and sediment transfer to and 846 from paddy fields is therefore a fundamental research question with a significant impact for both 847 human health and also for truly understanding the transfer of radiocesium throughout the fallout 848 affected region.

Author(s)	Туре	Number of samples	TF values
Endo et al. (2013)	Root (rice)	3	0.39 ± 0.08
	Straw (rice)	12	0.09± 0.02
	Ear (rice)	3	0.02 ± 0.01
	Brown rice	3	0.02 ± 0.00
	Polished rice	3	0.01 ± 0.00
	Rice bran	3	0.14 ± 0.01
	Chaff (rice)	1	0.05 ± 0.01
Takeda et al. (2014)	Soybean grain	46	0.09 ± 0.11
Tazoe et al. (2012)	Mugwort	6	0.02 ± 0.01
	Clover	1	0.03
	Grass	1	0.71

849 **Table 5**. Radiocesium transfer factor of different products collected after the FDNPP accident.

Understanding the storage of radiocesium within dams and reservoirs is important. As the decontamination efforts remove significant inventories of radiocesium from low-to-medium impacted areas, dams and reservoirs may present significant long-term radiocesium storages. These reservoirs are not only major dams that provide water for the region, they also include hundreds of smaller irrigation and farm dams. These dams could store significant quantities of radiocesium that could be redistributed during a major flood event throughout the landscape potentially contaminating, for example, previously decontaminated paddies. 857 Similarly, forests in the region will likely remain a significant long-term radiocesium reservoir. 858 Characterizing the storage and intra- and inter-forest radiocesium transfers will be necessary to 859 understand the transfer of radiocesium to and from forested hillslopes. This is important as forests 860 received the bulk of medium to high radiocesium fallout. Further, most of the forested communities 861 occur at high elevation on the coastal mountain ranges. Accordingly, they will likely act as a long-862 term source of dissolved and particulate (including sediment bound and organic matter bound) 863 radiocesium. Forests may therefore contaminate downstream rice paddy fields, rural areas and 864 rivers, again including potentially recently decontaminated landscapes. Importantly, it will be 865 necessary to quantify radiocesium transfers from both deciduous and coniferous forests in order to 866 accurately quantify radiocesium transfers from hillslopes to the Pacific Ocean.

In general, more research is needed to quantify radiocesium mobilization and transfer at the hillslope scale. In particular, a more detailed quantitative monitoring of soil erosion runoff under a variety of land uses is necessary to improve our understanding of the generation and transfer of sediment bound radiocesium at the hillslope scale, including the impact of different land uses (e.g. different forests, rice paddy fields, rural residential and other cultivated landscapes). Direct monitoring of hillslope erosion is necessary to ensure catchment scale models accurately reflect hillslope scale erosion dynamics.

874 Along with a more accurate quantification of hillslope erosion dynamics, catchment scale models 875 require a rigorous incorporation of the regional climate along with the important drivers and 876 dynamics of radiocesium transfer from deposition zones. In particular incorporating the seasonality 877 of flooding events to coincide with changes in vegetation in the forest and rice paddy landscapes may 878 improve the quantification of radiocesium transfer in the region. In the fallout-impacted catchments, 879 most radiocesium is transported in particulate form, mainly during floods generated by typhoons or, 880 to a lesser extent, by snowmelt. Further research into this seasonal cycle is important to understand 881 radiocesium dynamics in the Fukushima region. Indeed, it is necessary to fully comprehend the

882 frequency and magnitude of heavy rainfall events in the past, in order to properly forecast potential 883 rainfall driven radiocesium export in the future. Installing long-term riverine monitoring stations 884 throughout the fallout impacted catchments to monitor flow and suspended sediment information 885 while providing relationships to radiocesium would provide both a consistent long-term monitoring 886 of radiocesium transfers. This long-term monitoring riverine data should be coupled with long-term 887 monitoring of erosion from soils along with forestry radiocesium transfers to provide a concrete 888 foundation for effective and accurate long-term modelling research through the provision of both 889 reliable long-term calibration data and plot-scale transfer factors necessary to populate catchment 890 scale models.

891 Ultimately, this type of holistic approach, incorporating the different fundamental components 892 driving radiocesium transfer, is necessary. For example, research derived from monitoring at the 893 local forest scale regarding the different transfer dynamics within different tree communities is 894 needed in catchment modelling studies along with erosion data from the hillslope scale runoff 895 research. Indeed a transdisciplinary approach, with a team including hydrologists, geomorphologists, 896 soil and forest scientists, and mathematical modellers, will likely provide the best understanding of 897 radiocesium transfer and migration in the fallout affected region. Through working together, a 898 transdisciplinary team of researchers holds the greatest potential to accurately quantify the transfer 899 of radiocesium from hillslopes to the Pacific Ocean. Indeed, such an approach will also greatly 900 increase our understanding of sediment and material transfers in coastal mountainous environments 901 subject to frequent typhoons.

902 6. Conclusions

The majority of radiocesium fallout from the FDNPP accident was deposited on forests covering 66% of the region exposed to the highest fallout levels. Although the majority of this fallout was transported to the soil with throughfall, litterfall may become a dominant pathway of contamination transfer in the future. In soils, most of the radiocesium remains in the upper 5 cm of the soil profile

907 except in tilled fields where radiocesium is homogenized throughout the plough layer. Radiocesium is 908 strongly fixed to the finest mineral particles. Runoff generated from snowmelt and heavy rainfall 909 events generates and transfers sediments and their bound radiocesium. In fact, the major typhoons 910 in the region are known to transport the majority of radiocesium from hillslopes to the Pacific Ocean.

911 The FDNPP accident provides an important opportunity to research sediment transfers in general, 912 and radiocesium transfers in particular. Owing to the affinity of radiocesium to fine sediment 913 particles, there is an opportunity to use the accident to understand more about material transfers in 914 coastal mountainous regions subjected to massive typhoons. This understanding may be particularly 915 useful with the anticipated elevated rainfall intensities expected over the next century. Much can be 916 learnt from the Fukushima accident. The application of this learning is fundamental to manage the 917 fallout from this accident now, fallout from potential future accidents, and also simply increase our 918 understanding about the management of contaminants in rivers and riverine geomorphology in 919 general.

920 Acknowledgements

921 We would like to thank the anonymous reviewers for their comments and feedback on the 922 manuscript. This work has been supported by the French National Research Agency (ANR) in the 923 framework of AMORAD project (ANR-11-RSNR-0002).

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