

On the applicability of Persistent Scatterers interferometry (PSI) analysis for long term CO2 storage monitoring

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 interferometry (PSI) analysis for long term CO₂ storage monitoring Jeremy Rohmer^{*1}, Daniel Raucoules[*] *BRGM, Natural Risks Division, 3, av. Claude Guillemin BP 36009, 45 060 Orléans Cedex 2, FRANCE 	1	On the applicability of Persistent Scatterers
3 storage monitoring 4 Jeremy Rohmer ^{*1} , Daniel Raucoules [*] 5 *BRGM, Natural Risks Division, 3, av. Claude Guillemin BP 36009, 45 060 Orléans Cedex 2, FRANCE 6	2	interferometry (PSI) analysis for long term CO_2
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7 Abstract

1

8 Persistent scatterers synthetic aperture radar interferometry (PSI) analysis is viewed as a 9 promising monitoring technique for multiyear detection of ground surface displacements 10 during long term CO_2 injection into deep aquifers (of the order of 25 to 50 years). 11 This technique was successfully utilized at the CO₂ storage site of In-Salah (Algeria). 12 Nevertheless, this success may not be generalized to future onshore industrial-scale CO₂ 13 storage sites, which are planned in far more complex contexts mostly presenting natural 14 terrains, either agricultural or vegetated areas, i.e. characterized by poor PS density, which 15 might locally drop to zero, hence hampering the deployment of the PSI technique. The 16 objective of the present paper is to discuss the constraints on the PSI application regarding the 17 specificities (magnitude, spatial or temporal scales and evolution) of the deformation 18 phenomenon associated with the CO₂ injection into storage aquifers. This discussion is 19 supported by the estimates of the expected spatio-temporal evolution of surface vertical 20 displacements caused by CO₂ injection in deep aquifers by combining the approximate 21 solutions for pressure build-up and for the associated vertical displacements in an elastic half-

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space. A possible option to enhance the performance of the PSI technique is explored, namely
the deployment of a network of corner reflectors CR corresponding to artificial devices
installed on ground to complement the existing "natural" PS network. A methodology is then
proposed to optimize the CR network (number and position) in an iterative strategy (e.g.,
every 10 years) relying on the prediction of surface deformation. The methodology is
illustrated by a CO₂ injection scenario in the Paris basin context.

Keywords: CO₂ geological storage; Monitoring protocol; InSAR remote sensing; Persistent
Scatterers; geomechanical model; Corner Reflectors.

31

32 Introduction

33 CO₂ capture and geological storage (denoted CCS) is seen as a promising technology in the 34 portfolio of measures required to mitigate the effects of anthropogenic greenhouse gas 35 emissions (IPCC, 2005). Among the available geological targets, deep aquifers are recognized 36 to offer very large potential storage capacity with a broad distribution throughout the world in 37 all sedimentary basins (e.g., IPCC, 2005). Therefore, the present paper will primarily focus on this option. The International Energy Agency (IEA) recently evaluated the contribution of 38 39 CCS to emissions reductions by 2050 to one fifth in order to achieve the most cost-effective 40 stabilisation of greenhouse gas concentrations (IEA, 2009). On this basis, the IEA concludes 41 that the implementation of the technology should reach 100 projects in 2020, and more than 42 3000 in 2050 representing an amount of 10 Gt of CO₂ stored per year. These figures highlight 43 the large number of storage sites that shall be operated. Nevertheless, a prerequisite to any 44 CCS large scale industrial development is the demonstration that the storage is both efficient 45 and safe as outlined by the European directive on geological storage of carbon dioxide

(directive 2009/31/EC, April 2009). In this context, monitoring techniques constitute a key
component of any CO₂ storage risk management strategy aiming at providing information on
the CO₂ fate (migration and trapping) within the reservoir and demonstrating that CO₂
containment is effective (e.g., Benson, 2006).

50 Among the possible monitoring techniques, the space-borne Differential SAR (Synthetic 51 Aperture Radar) interferometry may constitute a promising and cost-effective option, as it can 52 provide valuable information on surface deformation induced by volumetric changes through 53 time and space resulting from CO_2 injection in deep aquifers (Tamburini et al., 2010). This 54 technique uses the phase of the radar signal in order to derive an estimation of the 55 displacement of the ground surface between two passes of the satellite over the area. Further details can be found for instance in Raucoules et al. (2007). The basic idea is that (once 56 57 removed flat earth and topographic components) the phase difference is proportional (modulo 58 2π) to the Line of Sight displacement occurred between the two acquisitions (assuming that 59 we can identify a stable reference area on the image. If not, the displacement is obtained with 60 an unknown constant value offset). A major interest of the technique lies on the fact that the 61 spatial sensor is able to image the ground surface regularly (presently, repetitiveness varies 62 depending on characteristics of the available satellites between few days to more than one 63 month).

Yet, one major limitation results from the temporal decorrelation phenomena (Zebker and Villasenor, 1992) stemming from changes in the physical (e.g. dielectric characteristics) or geometric (e.g. the distribution inside of the image elementary cell) characteristics of the backscattering objects. Multi-temporal processing of interferometric series is generally needed to cope with this shortcoming. Among the possible multi-temporal techniques, PSI (Persistent Scatterers Interferometry – e.g. Ferreti et al., 2000) are advanced techniques based on the processing of numerous, typically more than 30, images to identify a network of 71 persistent, temporally stable, highly reflective ground features -persistent scatterers (denoted 72 PS). These scatterers typically are cultural features of the developed landscape such as buildings, utility poles, roadways, etc., but also natural stable elements with stable and strong 73 74 backscattering characteristics (such as stable rocks with specific orientation). 75 To date, such space-borne techniques have been successfully used in a variety of domains 76 involving subsurface fluid withdrawal / injection and surface deformation response, for 77 instance groundwater flow characterization in complex aquifers systems (Bell et al., 2008), oil 78 / gas fields characterization (e.g., Vasco et al., 2008a),-monitoring of seasonal gas storage 79 (e.g., Teatini et al., 2011). More recently, surface deformation have been measured at In Salah 80 Gas Project in Algeria, where 0.5–1 million tonnes CO₂ per year have been injected over the 81 past 5 years (e.g. see site context described by Rutqvist et al., 2010). This analysis revealed an 82 average ground uplift of 5 mm/year and allowed getting better insight not only in the CO_2 83 migration at depth (e.g., Mathieson et al., 2009), but also in the behaviour of the storage 84 reservoir (e.g., Vasco et al., 2008b) and of the storage complex during injection operations 85 (e.g., Rutqvist et al., 2010). 86 However, such a successful result might not be easily generalised to any onshore CO₂ storage 87 sites, because the In Salah case is characterized by very favourable conditions for the 88 implementation of the PSI technique, the site being located in an arid region so that the 89 Permanent Scatterer PS density is very high, exceeding 1,000 per km² (300,000 PS used by 90 Tamburini et al., 2010 over an area of approximately 10 km by 15 km). 91 The objective of the present paper is to get a better insight in the applicability of the Persistent 92 Scatterers Interferometry implementation regarding the conditions of the future storage sites 93 and the specificities of the expected ground deformation phenomenon, i.e. the expected

94 spatio-temporal evolution of surface displacements caused by CO₂ injection operations at

95 depth.

96 The remainder of the paper is organized as follows. We estimate in a first section the expected 97 surface displacements by means of an analytical model. Though inherently based on 98 simplifications, the analytical results allow investigating a large variety of injection scenarios 99 (e.g. depth, injection rate, storage aquifer characteristics, etc.) useful to support the discussion 100 in section 2. Given the discussed limitations for the CO₂ storage context, we propose to 101 explore, in a third section, a possible strategy to enhance the performance of the PSI technique 102 for CO₂ storage monitoring by deploying a network of "artificial" PS, i.e. corner reflectors CR 103 (e.g., Novali et al., 2005) in an iterative design approach over the whole injection duration. 104

105 **1. Estimates of the expected spatio-temporal evolution ground**

106 vertical displacements

In this section, we briefly describe an analytical solution for assessing ground displacement assessment during CO_2 injection in deep aquifers (section 1.1). Details of the analytical model are provided in Appendix A and B. This model is then applied to a typical injection scenario (section 1.2) and its applicability is discussed (section 1.3).

111 **1.1.** Analytical solution for ground displacement assessment

In the present article, the expected spatio-temporal evolution of ground vertical displacements
due to CO₂ injection at depth are estimated using an analytical model developed by
combining :

- 115 the approximate solutions for pressure build-up (i.e. the pressure increase due to
- 116 injection) during CO₂ injection in brine aquifers of very large spatial extent (seen as
- 117 open systems) based on the studies of Mathias et al. (2009a&b);

the poro-elastic solution for vertical surface displacements estimate in an elastic halfspace, originally developed in the field of hydrocarbon and geothermal activities by
considering an axi-symmetric thin disk reservoir with general reservoir pressure
distribution (Segall, 1992; Segall and Fitzgerald, 1998).
For sake of clarity, the main assumptions and formulations of both analytical solutions are

respectively described in Appendix A and B. Noteworthy, each of these solutions have been
validated through numerical simulations (Mathias et al., 2009a) or against in situ observations
(at Lacq in southern France, see Segall, 1992).

126 1.2. Analysis of a typical injection scenario

127 We assess the spatio-temporal evolution of ground vertical displacements resulting from a 128 typical scenario corresponding to 1 Mt/y CO₂ injection during 25 years into a saline aquifer 129 formation at 1500m depth. Considering a pressure and geothermal gradient of respectively 130 $\pm 10^5$ Pa/m and 0.03 °C/m, the initial pressure and temperature are of 15 MPa and 45 °C at 131 1500m depth. The homogeneous properties of the aquifer formation are typical of sedimentary formations suitable for CO₂ storage, with high-enough permeability and porosity 132 133 (see for instance Birkholzer et al., 2009 for typical hydrogeological properties and Rohmer 134 and Seyedi, 2010 for typical geomechanical properties) and are summarized in Table 1. 135

136

[Table 1 about here]

137

Figure 1 depicts the spatial distribution of changes in vertical ground displacements estimated using the equations described in Appendix A and B. Different injection durations are considered and vary from 1 month to 25 years. The pressure build-up reaches a maximum value within the injector near-zone and decreases deeply in an exponential fashion within the CO₂ plume region (the extent being marked by the steep decrease of the pressure build-up in

143 Figure 1A) and linearly within the brine-filled region. The vertical surface displacements 144 follow the same trend (Figure 1B), but with a smoother form of the spatial distribution so that 145 these are approximately constant over approximately 1 km and steeply decrease reaching (and 146 even exceeding) the lateral extent of the pressurized zone at depth (i.e. the zone for which the 147 pressure build-up is non-nil). For instance, after 2 years of injection, the vertical surface 148 displacements exceed 4 mm over the lateral distance range of 0 to 1 km and decrease reaching 149 ~10 km given that the extent of the pressurized zone reaches ~8 km. The smooth spatial 150 distribution is explained by cumulative contribution of all the rings of dilation with magnitude 151 proportional to the pressure build-up (i.e. decreasing with the lateral distance) as shown in 152 Appendix (Figure 9). 153 Let us consider the gradient (in absolute value) of displacements corresponding to the difference between vertical displacements respectively observed at positions \vec{r} and of $\vec{r} + \vec{\Delta}$ 154 divided by the distance $\|\vec{\Delta}\|$, where $\|\vec{\Delta}\|$ is the characteristic length of a satellite image pixel 155 156 (typically between 10 and 50 m depending on sensor and processing). This gradient is largely 157 smaller than the maximum allowed value corresponding to half a fringe per pixel, *i.e.* of about $1.5 \text{ cm}/10\text{m}=1.5\text{x}10^{-3}$ to $1.5\text{cm}/50\text{m}=3.10^{-4}$ for vertical movements with ERS data (Raucoules 158 159 et al., 2007). This value would be modified for sensors with different resolutions and 160 wavelengths, the maximum measurable gradient being proportional to wavelength and 161 inversely proportional to pixel size. Even in case of a point-based method (such as PSI techniques) with larger $\|\vec{A}\|$, typically ranging 500m to 1000m in rural zones (in this case $\|\vec{A}\|$ 162 163 does not correspond to the pixel size, but to the distance between PS), this maximum 164 measurable gradient should not be theoretically a limitation, because the expected total 165 deformation of the order of ~1 cm, as shown in Figure 1B, corresponds to less than one C-166 band sensor fringe (equivalent to ~3 cm of vertical displacement).

[Figure 1 about here]

168 Considering the temporal evolution of the maximal surface displacement within the injector 169 near-zone, Figure 2 shows that it reaches a maximum value of ~11 mm over a time span of 25 170 years and evolves following a logarithmical function over time. This trend is partly explained 171 by the temporal evolution of the maximum pressure build-up within the injector near-zone. 172 But, given its rapid evolution (after 1 month of injection, it reaches more than 85 % of the 173 maximal pressure build-up), the temporal evolution is in great part explained by the 174 cumulative contributions of rings of dilation, as shown in Figure 9, (Appendix A), whose 175 number increases over time as the pressure front propagates within the aquifer formation. 176 [Figure 2 about here]

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177 **1.3.** Applicability of the analytical solution

178 The described analytical model is proposed as a quick-assessment tool, which should be used 179 to capture the general trend of surface deformation due to CO₂ injection, i.e. to estimate 180 orders of magnitude of the expected vertical displacements either to assess the likelihood of a 181 risk event associated with ground surface perturbations in a preliminary risk analysis of the 182 storage site (e.g., Bouc et al., 2011), or, as proposed in the present paper (section 2 & 3), to 183 investigate the applicability of the PSI analysis considering a broad range of different 184 scenarios. Noteworthy, the application of the model to the In-Salah's case considering rock 185 properties and injection scenario described by Vasco et al. (2010) produces a measurable 186 surface displacement of approximately 4 mm/y, i.e. in the order of what has been actually 187 measured. 188 An in-depth understanding of the surface deformation process due to CO₂ injection at depth

189 should consider several aspects, which the analytical model hardly accounts for, most of them 190 being related to the complexity of both the deformation processes and the geological 191 subsurface structure. The described model assumes poro-elastic rock behaviour, hence implying that uplift instantaneously follows linearly the reservoir volumetric increase. Though confirmed through in situ compaction measurements (see e.g. the Groningen field, studied by Mobach and Gussinklo, 1994), a nonlinear dependence may also appear in some cases. See detailed discussion by Hettema et al. (2002). Furthermore, the ground deformation processes may be complicated by the presence of soils and unconsolidated rocks near the surface (e.g., Singh, 198).

199 The proposed model considers an idealized view of the subsurface structure so that both the 200 reservoir and the surrounding rock are considered homogeneous and having the same material 201 properties. Yet, aquifer formations targeted for CO₂ storage are mostly located within large 202 sedimentary basins, which can be viewed as multi-layered systems (see e.g., Birkholzer et al., 203 2009). Analytical models have been proposed to account for this aspect (Fokker and Orlic, 204 2006) and can be extended as a future direction of research to the CO₂ storage case in the 205 similar approach as proposed in section 1. Furthermore, the presence of geologic 206 discontinuities (i.e. the existence of faults) may increase ground movement potential, but on a 207 localized spatial scale, as numerically demonstrated by Rutqvist et al., (2010).

208 2. Constraints on the implementation of PSI technique

In turn, we examine constraints to the implementation of the PSI technique regarding the specificities of the deformation phenomenon associated with the CO_2 injection at depth, in terms of magnitude (section 2.1), spatial and temporal scales (section 2.2 and 2.3) and evolution (section 2.4). Table 2 summarizes the constraints on the PSI implementation and the needs in terms of sensor characteristics in order to address the issues resulting from the expected motion characteristics.

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[Table 2 about here]
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216 **2.1. Magnitude of the deformation**

It is difficult to determine beforehand whether the specific deformation phenomenon associated with the CO_2 injection at depth (as illustrated by the typical injection scenario described in section 1.3), may be monitored, because the feasibility and accuracy of PSI strongly depend on the number of satellite acquisitions and the temporal baselines (e.g. Hanssen, 2005).

222 The value of the precision PSI techniques has been widely discussed in the literature and can 223 vary from less than 1 mm to centimeters. Colesanti et al. (2003) suggested that in an 224 operational context under very good conditions (large data sets -i.e. >30 images, dense PS 225 network, quasi linear deformation regimes), the precision could reach 0.5 mm/y for the annual 226 deformation rate and 3 mm for the displacement values at given acquisition dates for PSI 227 based on data from ERS, Envisat/ASAR or equivalent (C-band, resolution ~10m, 228 repetitiveness of ~ 1 month). However, in less favourable cases (deformation rates higher than 229 some cm/y, sparse PS, non-linear deformation) results could be worse. Studies based on inter-230 comparison of different PSI techniques and validation respect to ground measurements (see 231 for instance Raucoules et al. 2009) showed that in unfavorable cases, the precision value can 232 drop to 1.5 mm/y and the accuracy can be much worse (>5 mm/y). Furthermore, deformation 233 superior to ~cm/y tends to be under-estimated if no prior assumption on the deformation is 234 used in the processing (Raucoules et al., 2009). This issue is not expected for CO₂ storage 235 applications as shown by the estimates of ground displacements resulting from a typical CO_2 236 injection scenario (see Figure 1B).

237 On this basis, a detection threshold for an individual deformation value of 5 mm (i.e. below

which the deformation phenomenon is very unlikely to be detected) is considered in the

present study. For the considered injection case in section 1.3, we show that deformation

240 phenomenon associated with the CO₂ injection is likely to be detected only after the first year

241	of injection. Different injection scenarios are then considered with injection rates and depths		
242	respectively ranging from 0.5 to 2.5 Mt/y and 1000 to 2500 m and assuming the aquifer		
243	formation properties constant as described in Table 1. Figure 3A) shows the maximum		
244	vertical displacement reached at surface in the injector near-zone after respectively 1 and 10		
245	years of injection (top and bottom left panels). Several aspects can be outlined:		
246	• The maximal vertical displacements is proportional to the mass injection rates,		
247	whereas it is non-linear with depth mainly due to the non-linear form of the		
248	relationship between the fluid properties of CO ₂ and brine with depth;		
249	• The depth of the aquifer storage formation by 500m implies increasing the necessary		
250	injection rate by 0.5 Mt/y for the surface deformation to be detectable. For instance,		
251	ground displacements after 1 year caused by CO_2 injection at 0.5Mt/y within a 1000m		
252	deep aquifer is equivalently detectable as those caused by CO_2 injection at 1.0 Mt/y		
253	within a 1500m deep aquifer;		
254	• For the whole range of depths and time durations, the deformation associated with the		
255	0.5 Mt/y injection scenario is hardly detectable ;		
256	• Only the deformation associated with the 1000m deep injection may be detectable		
257	over the whole range of injection rates.		
258	[Figure 3 about here]		
259			

260 **2.2.** Spatial scale

The basic idea of PSI techniques is that there exist points with stable backscatter characteristics during the period covered by the data series. In other terms, those points (or PS) have high (respect to other pixels of the images) backscattered amplitude that does not change significantly from one image to other. They are generally considered as objects with extents smaller than the pixels where they are located. Such PS are little affected by geometric and temporal decorrelation and allow (using statistical criteria) to compensate the atmospheric phase screens in order to improve the precision of the measurement. In urban area, buildings and other man-made structures often act as PS due to their corner reflector-like scattering behavior and high radar reflectivity so that we can obtain a very high density (hundreds per km²) of PS, even exceeding thousands per km² in arid regions (e.g., Bell et al., 2008).

271 However, the total area at ground surface to be affected by the deformation phenomenon 272 caused by CO₂ injection at depth is expected to be very large in relation with the large over-273 pressurized zone at depth. Several authors have recently shown that the latter can reach of the 274 order of several to more than a few tens of kilometers (e.g. Birkholzer et al., 2009). The 275 detectable zone of deformation at surface will then be characterized by a lateral extent of the 276 order of a few kilometers in the early stage of injection (ranging between 0.5 and less than 5 277 km after 1 year of injection, Figure 3B, top) and of more than ten of kilometers (ranging 278 between ~5 km and less than 15 km after 10 years of injection, Figure 3B, bottom). Such a 279 large size of the monitored area implies that the size of the area covered by the satellite 280 acquisitions (i.e. swath) should ideally be larger than 50 km. Yet, the larger the swath, the 281 lower the resolution of the sensor. For instance, TerraSAR-X (mode scanSAR) presents a 282 very large swath of 100km, but with a resolution of ~16m, which may be insufficient (see 283 discussion below in section 2.3 and 2.4). Conversely, very high resolution sensors (~1m, such 284 as TerraSAR-X or Cosmo-Skymed) offer insufficient coverage (~10km) and their use implies 285 acquiring and processing a much larger amount of image data, which may become cost and 286 resource intensive.

Furthermore, over such large areas, we can reasonably expect a vegetated cover (fields, forests, etc.) on a significant part of the storage site. For illustration purpose, the PICOREF sector in the Paris basin (of 3,455 km², see location in Figure 6), where the potential implementation of CO_2 storage sites have been investigated (e.g. Grataloup et al., 2008), is characterized by a land use composed of less than 5 % of artificial / urban areas and more than 95 % of natural (mainly forests) and agricultural areas. In such natural terrains, the absence of bright man-made structures might lead to sparse PS network (with density ~10 PS every km² to less than 1/km²). This constitutes a major obstacle to the phase unwrapping stage, where the solutions are directly dependent on the PS density and makes reliable estimation of deformation using PS techniques a challenging task.

The use of L-band data may contribute to improve the processing given that the vegetation is better penetrated using signal with short wavelength (e.g. Raucoules et al., 2007). Unfortunately there is currently, to the authors' knowledge, (and probably for the next 2 decades) no planned satellite mission fulfilling the requirements in terms of resolution and/or acquisition repetitiveness for the proposed monitoring application as discussed below.

302 2.3. Temporal scales

303 CO₂ injection operations are typically planned to last between 25 and 50 years. This "long
 304 term" monitoring implies considering several issues.

305 A first difficulty is the availability of a space-borne sensor, whose mission durations are 306 typically of the order of a ten of years (e.g. see Hanssen, 2005, Table 2). Thus, any long term 307 monitoring of the injection operations should be designed potentially considering a minimum 308 of 2 to 3 satellites, whose characteristics (orientation of the orbital trajectory, center 309 frequency, C-, L- band, etc.) might change from a satellite to another and not be necessarily 310 suited for interferometric analysis. For instance, the currently-used ENVISAT mission, which 311 was deeply modified in October 2010 (change in the orbital characteristics, affecting the 312 possibility of carrying out reliable PSI processing after this event), will come to an end in the 313 coming years. A transition period with another space-borne sensor should be accounted for in 314 the design of long term monitoring plan. For instance, the successor of ENVISAT, namely 315 Sentinel 1, is to be launched in 2013. This 7-12 years mission will consist in a space-borne SAR in C-band (and therefore providing data similar to ASAR) with a repetitiveness of 12
days. Considering an ideal data-set for PSI baseline of >30 images the transition period
should last ~10 months, i.e. before the planned commercial-scale deployment of storage sites
by 2020 (IEA, 2009).

A second difficulty arises from the decorrelation of PS over time, i.e. the "life expectancy" of the PS over such long time durations. Ferreti et al. (2003) tackled this issue by identifying points corresponding to PS only on sub-sets of the full ESA-ERS data set over the period 1992-2003 on the Milano urban area. This consisted in identifying PS, which have appeared or have disappeared during a given time span between the first and last acquisition. The number of such points is usually 10% of the number of PS over a period of a ten of years. This can be considered an order of magnitude of the life expectancy of a PS.

Though requiring further investigations, the effect of the land cover nature (urban / rural) is expected to have little influence on the life expectancy of PS. It is true that on vegetated area the temporal decorrelation (on none-PS points) is globally higher than on urban areas, but pure point target PS are little affected by temporal decorrelation (as they are identified by the temporal stability of their signal) whatever on urban and on non-urban areas.

332 The use of high resolution data is expected to improve the life expectancy of PS. As shown by 333 Crosetto et al. (2010), higher PS densities are expected with higher resolution sensors. 334 Assuming an equivalent PS decay rate for different resolutions, the total number of PS is 335 expected to be larger with high resolution sensors considering an isolated small area with high 336 PS density located in a larger area depleted of PS (e.g. small groups of buildings in a larger 337 agricultural zone). Following this reasoning, the probability of having at least a few PS lasting 338 during the whole observation period is therefore higher with high resolution sensors (provided 339 that a maximum of isolated building groups will keep some PS during the observation 340 period). Yet, as mentioned in section 2.2, the high resolution requirement is limited by the large swath requirement. Current sensor TerraSAR-X (Strip Map Mode) and future Sentinel-1
mission should both fulfill such a trade-off, but with major differences in the data provision
policy.

344 This latter issue should be carefully accounted for, because long term monitoring possibly 345 implies constituting a large data base (depending on the acquisition and processing frequency; 346 which is further discussed in section 2.4). To remain cost-effective regarding ground-based 347 geodetic measurement techniques (e.g., optical levelling, GPS, tiltmeters, etc.), a possible 348 strategy can be similar to the provision policy associated to the Global Monitoring for 349 Environment and Security (http://www.gmes.info/), which allows obtaining the data at low 350 price as envisaged for the future Sentinel-1 mission. The interested reader can refer to (Rucci 351 et al., 2012) for additional discussions on the use of future Sentinel-1 and new X-band sensors 352 for reservoir applications.

353 2.4. Spatio- temporal evolution

354 The monitoring of CO₂ storage sites has two major objectives (see e.g., Benson, 2006). A first 355 objective, categorized as "verification", focuses on the monitoring of the expected behaviour 356 of the storage complex, in particular detecting the location and quantifying the size of CO_2 357 plume within the reservoir and tracking the pressure front propagation. The second objective, 358 categorized as "early-warning", focuses on the detection of any "abnormal behaviour", i.e. 359 deviations from the expected behaviour of the storage complex. This latter case corresponds 360 to the occurrence of unwanted events such as the sudden (i.e. over a short time span) multi-361 millimetric uplift localized on a fault zone, which may be related to the fault reactivation at 362 depth (see e.g., Rutqvist et al., 2010).

363 Regarding the "verification" objective, the acquisition and processing rate should be chosen

364 based on the simulations of the expected spatio- temporal evolution of the deformation

365 phenomenon. In this view, section 1 provides useful orders of magnitude for a typical

injection scenario. At a given location, Figure 2 shows that the deformation phenomenon is 366 367 characterized by a highly non-linear evolution over time. This might limit the performance, 368 because the PSI technique is generally better applied for rather conservative types of 369 deformation (e.g. Hanssen, 2005), e.g., characterized with (approximately) uniform 370 deformation velocity relatively linear temporal evolution of deformation velocity or periodic 371 (such as seasonal natural gas storage, Teatini et al., 2011). To tackle this issue, an alternative 372 would be to focus on the discrepancies of the measured displacement field respect to the 373 proposed evolution model, i.e. deviation to the predictive model. By subtracting the model-374 based simulation of the phase to the interferometric data, an important part of the (non-linear) 375 temporal signal would be removed, facilitating the PSI processing (in particular by reducing 376 the risk of unwrapping errors). Besides, the important information is the (unforeseen) 377 discrepancy respect to the model that is needed for adjusting the representation of the 378 phenomenon. The proposed iterative procedure described in section 3.3 that revisits 379 periodically the knowledge of the phenomenon, could partly deal with this issue. 380 Figure 1B provides information on the spatio- temporal evolution of the "deformation front" 381 given a typical injection scenario (see details in section 1.2). An individual measurement point 382 located at a given lateral distance from the injection zone, requires a major and more 383 frequently sampled monitoring during a specific period of the order of \sim 2-4 years, the 384 beginning of this period depending on the distance to the injector. In other terms, a 385 "deformation front", as schematically depicted in Figure 4, affects the location during this 386 short time span $\Delta t=2-4$ years. Before the front has passed no deformation affects the point. 387 After, the point has smaller displacement (see Fig. 4). Such a spatio- temporal behaviour has 388 implications on both the choice of the frequency of image data acquisition (i.e. sampling 389 procedure) and of the processing rate (i.e. frequency of PSI analysis).

At a given location, the monitoring should be highly-rated sampled during the short period of the deformation front passage. As indicated in section 2.1, >30 images is an ideal data set size for precision reasons. The orbit cycle of ASAR-like sensors (about 10 acquisitions per year) may be therefore insufficient to carry out such a measurement at the limit of the sensitivity of the method. Current sensor TerraSAR-X (Strip Map Mode) and future Sentinel-1 mission should fulfill such a requirement (with both repetitiveness of the order of a ten of days), but with differences in the data provision policy (see section 2.3).

397 Besides, changes in the velocities values of the differential displacements are expected to be 398 detectable only after a few months (except for the case of multi- millimetric sudden ground 399 movements, which is discussed in the following). The processing of the data image with very 400 high frequency (e.g. every day) may not be appropriate given the described physical evolution 401 of the deformation, but also considering cost and resource constraints. Considering the 402 expected spatio- temporal evolution of the deformation phenomenon as calculated for the 403 typical injection scenario described in the section 1.2, one PSI analyse every 3-6 months can 404 be proposed to fulfil the "verification" objective of monitoring. Noteworthy, this processing 405 rate is intermediate between quasi- real time (e.g. bottom-hole fluid pressure) and pluri-406 annual monitoring techniques (e.g. seismic campaign).

407 Regarding the "early- warning" objective of the monitoring, two cases can be considered. On 408 the one hand, potential zones at risk (e.g. fault zones, injection zone, etc.) can be identified 409 prior the injection operations based on a risk assessment procedure (see Bouc et al., 2011). On 410 these localized zones, a specific PSI analysis can be envisaged using high resolution/high 411 repetitiveness (but reduced swath) data equivalent to those provided by the current X-band 412 sensors TerraSAR-X or COSMO-SkyMed (see e.g., Ferretti et al. 2011b), to complement the 413 PSI analysis for the "verification" monitoring. Noteworthy, other ground-based geodetic 414 techniques would require systematic on-site measurements, such an approach being time415 consuming and resource intensive.

416 On the other hand, if any "abnormal behavior" is detected in an "un-expected" zone either 417 after processing the image data or based on other indicators (induced seismicity recorded at 418 the surface due to possible fault reactivation; increase of the fluid pressure recorded in the 419 monitoring aquifer indicating a possible leakage, etc.), the considerations for data acquisition 420 and processing rate as discussed for the verification monitoring should be adjusted. The great 421 advantage of PSI analysis compared to ground-based geodetic measurement is the possibility 422 to assess deformation after the unwanted event has occurred, because image data would have 423 been archived without user's intervention (e.g., every12 days using the future Sentinel-1 424 sensor). Punctual and local studies on short time spans can also be proposed to focus on the 425 "un-expected" zone at risk to complement the PSI analysis for the verification monitoring. 426 [Figure 4 about here]

427 **3. Planning a network of artificial reflectors**

428 **3.1.** Description of corner reflector devices

429

430 One major difficulty to the implementation of PSI in the CO₂ storage context stems from the 431 potentially low PS density over such large areas as previously discussed. Though improved 432 co-registration and use of more efficient computational techniques (the interested reader can 433 refer to Sousa et al., 2011) can provide valuable support, a potential solution to enhance the 434 performance in regions where bright man-made structures are merely absent can rely on the 435 installation of artificial PS, i.e. corner reflectors CR (e.g., Novali et al., 2005). These artificial 436 metallic devices (of generally trihedral or cubic geometry characterized by a few tens of 437 centimeters long) are installed on ground providing high backscatter to the radar signal to

complement the existing "natural" PS network, i.e. to fill the gap between natural persistent
reflectors. This approach has been proposed in other contexts (mine, volcano, coastland and
marshland environments, etc.), where the ground movements' detection has proved to be very
challenging. See for instance the recent application case in the Venice lagoon (Teatini et al.,
2010).

443 **3.2.** Optimizing the position of the CR

444 To the authors' knowledge, very few works have handled the issue of planning the position of 445 CR in a predictive fashion, i.e. before any ground deformation measurement has been carried 446 out. For instance, in the recent interferometric analysis of the land subsidence in the Venice 447 lagoon (Teatini et al., 2010), a network of 55 CR has been planned taking into account a fixed 448 distance of 1 km between them of any adjacent natural reflectors. Yet, these were installed 449 only once the full interferometric analysis had been carried out for a ten of years. 450 Planning the CR network using an "one-size-fits-all" approach, i.e. using a fixed distance 451 value may rapidly become impracticable considering the large area expected at the end of 452 injection (of ~325 km² after 25 years for the case modeled in section 1.3). Optimizing the 453 location of the CR should result from a trade-off between the issues related to the installation 454 of such devices (costs, authorizations, maintenance, etc.) and the requirements in terms of 455 coverage and precision for the PSI results. Furthermore, the time-scales of CO₂ storage 456 constitute a major constraint for the monitoring plan that should be designed to follow the 457 long term evolution of the ground movement over a time span ranging from 25 to 50 years. 458 An example of a ground-based network, i.e. of devices fixed at the surface that exist over such 459 a long time duration, is provided by the networks of benchmarks used for national spatial reference systems (e.g. in France, the currently-used network has been installed since the 19th 460 461 century), but dimensions of such ground-based devices have nothing comparable to the ones 462 of a CR.

In this context, a crude design approach based on the expected surface deformation calculated
regarding the whole time duration of the storage project may be highly limited by both
maintenance issues of the CR network given such large spatio-temporal scales and by the
changes in the sensors characteristics (typically each 7-10 years, see section 2.2) over the full
lifecycle of the storage injection phase. Therefore, we advocate a design strategy thought in
an iterative fashion:

469

470 Step1: estimate the expected spatio-temporal evolution of ground surface 471 displacements induced by CO₂ injection operations at different epochs for the full 472 period of injection. The duration of each epoch should be no longer than a ten of years 473 to handle the possible change in the sensors characteristics. Let us consider the 474 estimates of the surface deformation after 5 epochs of 10 years each. This step should 475 be based on large-scale geomechanical analysis of the complex site. This can rely on 476 analytical models such as the multilayered visco-elastic solution developed by Fokker 477 and Orlic (2006) or in a first approach, the one described in section 1 presenting the 478 attractive feature of a rapid estimation, but limited by the simplifications inherent to 479 the analytical nature of the model of such a simplified model (as discussed in section 480 1.3). A more sophisticated analysis can rely on large-scale coupled numerical models 481 as demonstrated by Rutqvist et al. (2010) in the In Salah's context. The advantage of 482 such analysis are the ability to take into account complex injection scenarios, complex 483 reservoir pore pressure spatio-temporal evolution, complex geological structures (e.g., 484 multi-layered system as present in sedimentary basin suitable for CO₂ storage, 485 Birkholzer et al., 2009;

Step2: based on the expected extent of detectable deformation surface, localize the a
priori good candidate for PS, i.e. spatial points that appear to be persistent in the

present, but above all, that shall remain persistent over the full duration of the
injection. Long term "natural" PS can constitute man-made structures within urban
zones (historical monuments, churches, industrial plants, roads, etc.). We can consider
that a large number of these spatial points identified as PS will keep their
characteristics (reflectivity) over the time duration of the considered epoch (of a ten of
years). This assumption can be considered valid regarding the study of Ferreti et al.
(2003), as discussed in section 2.2.≒

495 Step3: optimize the position of the corner reflectors CR considering the expected 496 detectable surface deformation for the whole duration of the injection. We advocate 497 choosing the distance between CR and natural PS based on a criterion dependent on 498 the initial network of natural PS. The idea is to maximize the spatial coverage of the 499 area where ground deformation is expected to be detected considering the location of 500 identified potential "natural" PS. Formally, this can be achieved by treating the 501 "natural" persistent scatters as Voronoi sites and by affecting to each of them a 502 Dirichlet cell (also named Voronoi cell) consisting of all points closer to PS than to 503 any other site as schematically depicted in Figure 6A. The segments of Voronoi cell 504 are all the points in the space that are equidistant to the two nearest sites PS. The 505 potential positions of the CR are then chosen as the Voronoi nodes corresponding to 506 the equidistant to three (or more) PS sites. The positions of the CR merely derive from 507 a pure mathematical space-filling criterion, but in real-case applications, adjustments 508 will certainly be necessary if the Voronoi-based method indicates installing the CR in 509 densely vegetated cover (forests, wetlands, etc.), which may highly prevent the 510 detection or the installation of the CR, or too close to another CR (e.g., with a distance 511 <1 km). Then, the CR only lying within the expected detectable zone for the

considered 10-years epoch, can be selected. These CR should be installed to help 512 513 monitoring the expected surface deformation over the duration of one epoch. 514 Based on the discussions in section 2, the PSI analysis to fulfil the "verification" objective of 515 the monitoring can conducted using data characterized by a large swath, a rather good 516 resolution and repetitiveness of a ten of days. The processing rate of the image data should be 517 chosen depending on the simulation predictions for the spatio- temporal evolution of the 518 deformation phenomenon associated with CO₂ injection. The simulations for the typical 519 injections scenario (section 1.2) show that a processing of 3-6 months can be envisaged 520 during the 10-year epoch. This rate can be adjusted especially depending on the detection of a 521 sudden "abnormal behaviour" on a localized zone (e.g., sudden multi- millimetric uplift due 522 to possible fault reactivation). A complementary PSI analysis using high resolution / high 523 repetitiveness (but reduced swath) data can be proposed on these zones at risk (see discussion 524 in section 2.4). Such an iterative approach allows a frequent re-interpretation of the 525 discrepancies between model and measurements in terms of evolution predictions (i.e. 526 deviation to the predictive model). Valuable information on the storage complex behaviour 527 can be extracted in an inversion procedure as proposed more specifically in the field of CO₂ 528 storage by Vasco et al. (2008a&b). The combination of both information (correction of the 529 model based on the measurements and inversion of the complex behavior) can serve to 530 increase the quality and robustness of the predictions. 531 The procedure is then repeated for the next 10-years epochs (step 1). Using the initial PS 532 network enriched with the CR identified for the previous epoch (step 2), the next CR (i.e. 533 necessary for monitoring the next epoch) can be identified using the updated expected 534 detectable surface deformation region (step 3), and the procedure is then repeated.

535 Noteworthy, Before the end of each 10-years observation epoch, a transition period should be

536 envisaged between two successive satellite missions, with sensors of possibly different

characteristics. During this period, a sufficient amount of SAR data (> 30 images) should be
acquired for the PSI baseline using the new sensor. Considering a repetitiveness of a ten of
days, this period should last ~1 year. Figure 5 summarizes the sequence of the steps
considering two epochs and two different sensors.

541

[Figure 5 about here]

542 In case of sensors with different characteristics, a change in the configurations (e.g. 543 orientation) of the CR, which have been installed during the first epoch, should be envisaged 544 to make these CR "visible" by the new sensor. In such a case, a gap in the produced 545 deformation time series may be expected of the order of few months, but only partly on the 546 ground surface covered by the SAR data, i.e. corresponding to the region whose CR were 547 adapted to the first 10-years epoch. To the authors' knowledge, the transition between two 548 sensors during the implementation of SAR interferometry combined with CR has not been 549 experienced yet and requires further investigations.

Finally, considering such a long term monitoring, the possibility that during certain periods, the lack of satellite mission with suitable characteristics cannot be fully discarded. In such a case, the information provided before and after the information gap could still provide useful information on the evolution of the injection processes.

3.3. Application to a potential storage site in the Paris basin

To illustrate the main steps of the proposed methodology, we apply the described
methodology to a potential injection site selected within the PICOREF sector in the Paris

basin (Figure 6). This potential storage site was identified based on the combination of criteria

558 in terms of capacities and injectivity, environmental constraints, existing land-use and

underground-use, economic and social aspects and based on a comprehensive risks analysis

560 (Grataloup et al., 2008; Bouc et al., 2011). Figure 6 shows the location of the potential

561 injection well (outlined by a red cross marker). We consider 40 years of CO₂ injection at a

562	rate of 2 Mt/y into the 1500m deep carbonate aquifer of the Dogger formation. The aquifer
563	properties are: intrinsic permeability of 1.10 ⁻¹³ m ² ; porosity of 12 %; Young's modulus of 24
564	GPa; Poisson's ratio of 0.29 (Rohmer and Seyedi, 2010 and references therein).
565	[Figure 6 about here]
566	• Step 1: in this example, the analytical model described in section 1 is used to estimate
567	the lateral extent of the expected detectable zone of vertical displacements
568	(considering a 5 mm detection threshold) induced by a whole CO ₂ injection duration
569	of 40 years and considering four 10-years long epochs. The final zone represented by a
570	black circle in Figure 6 covers a total area of ~220 km ² and corresponds to the zone of
571	maximum extent that should be monitored by the PSI technique over the whole life
572	span of the potential storage site;
573	• <i>Step 2:</i> we consider each centroids of the urban zone present within the detectable
574	zone for the whole injection duration (total number of 24 zones). These represent
575	localized zones where good candidates for natural PS are highly likely to be detected.
576	In a preliminary approach, we only consider these centroids as natural PS (outlined in
577	Figure 7 as pink diamond markers);
578	• <i>Step 3</i> : on this basis, we assign to each of the centroids a Voronoi cell (Figure 7A) and
579	choose the potential CR as the Voronoi cells (outlined in Figure 7A as blue triangular
580	markers); the CR lying within forest zones or distant from each other of less than 1km
581	are detected and are re- adjusted (deleted or re-localized). A total number of 23 CR are
582	in this manner identified within the detectable region for the whole duration of the
583	injection (here 40 years, black -coloured circle in Figure 7B).
584	[Figure 7 about here]
585	Using the region of detectable deformation for the first 10 years of injection (light grey-
586	coloured circle in Figure 7B), a total of 6 CR are identified through the Voronoi analysis.

587	Among the identified CR, two CR should be re- localized, because they lie within forest areas
588	and one CR should deleted, because it is too close from the others. Without any other
589	consideration (such as legal issues, accessibility of the site,), we are able to propose a
590	monitoring procedure adapted to the expected characteristics of the deformation for the first
591	10 years of injection both based on CR-InSAR and on current PSI techniques.
592	For the next 10-years epoch, Figure 8 provides an illustration of the location of the
593	corresponding CR (i.e. after 10 years of injection) using the original PS network (Figure 7A,
594	pink coloured diamond markers) enriched with the 5 CR selected for the first 10-years epoch.
595	The expected detectable regions for the whole injection duration (at 20, 30 and 40 years) are
596	then re- estimated (step 1), here assumed unchanged for illustration purposes. On this basis,
597	the CR for the whole duration are identified (step 2 & 3). Then, using the expected detectable
598	deformation region for the next 10-years epoch (here for the date 20 years), the CR for this
599	epoch are selected (Figure 8B).

600

[Figure 8 about here]

601 4. Concluding remarks and recommendations

602 The present paper proposes a description of the characteristics of the surface deformation, 603 which will be expected during the injection of CO_2 into deep aquifers during 25 - 50 years. 604 On this basis, the associated constraints for the planning of a monitoring plan relying on PSI 605 techniques are discussed. To overcome such limitations, we focused on the optimisation of a 606 CR (corner reflectors) network. An iterative strategy is proposed, which is justified by the 607 large spatial extent of the expected detectable region and by the changes in the sensors 608 characteristics (typically each 7-10 years) that could result in a need of re-organisation of the 609 network (and even CR characteristics such as orientation). At each period (say, every 10 610 years), the number and location of the CR, necessary for the monitoring of the surface

611 deformation of the next epoch, are estimated based on the prediction of surface deformation 612 using large-scale geomechanical models. Furthermore, we propose to focus on discrepancies 613 between model and measurements, which can be useful in terms of evolution predictions. To 614 fulfil the "verification" objective of the monitoring, the PSI analysis should be frequently 615 conducted during each epoch with a processing rate, which should be chosen depending on 616 the predicted spatio- temporal evolution of the deformation phenomenon. The simulation 617 results for a typical injection scenario show that a processing rate of the order of a few months 618 may be sufficient. Such a frequent re-examination of the situation and model correction is 619 therefore useful in a perspective of forecasting consequences of the injection. To fulfil the 620 "early-warning" objective of the monitoring, a complement should be envisaged so that well 621 definite zones at risk (such as fault zones) should be monitored using high repetitiveness / 622 high resolution (but reduced swath) data. In both cases, a careful attention should be paid to 623 cost and resource issues especially regarding provision policies of image data. 624 Finally, from now to the start of injection and during the 50 years injection, improvement of 625 PSI processing techniques can occur. For instance, Ferretti et al. (2011a), recently proposed a technique (SqueeSARTM) based on the concept of diffuse scatterers, which can improve 626 627 considerably the density of usable "natural" scatterers on certain un-urbanised area. See the 628 recent application of such a technique at In-Salah (Rucci et al., 2012). If such improvement is 629 efficient for surface characteristics such as those existing on the monitored area, an 630 optimisation of the network could be proposed – probably based on a reduced CR number. 631 The periodic re-thought of the CR network will therefore allow the possibility of 632 methodological improvement in addition to taking into account new sensors characteristics. 633

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637

638 Appendix A: General solution to ground displacements

Segall and co-authors (Segall, 1992; Segall and Fitzgerald, 1998) have presented poro-elastic solutions for stress and displacement fields estimate in an elastic half-space, induced by fluid extraction (or injection, the solutions being merely the opposite) for general axi-symmetric geometries with arbitrary pressure radial distribution within the reservoir (given as a function of space and time for instance from field data and/or reservoir simulations). Noteworthy this solution provides both horizontal and vertical surface displacements, but we will focus on the latter one.

The variation in reservoir pore pressure *p* is assumed to mainly govern the volumetric strain in the reservoir. The vertical displacement at the radial distance *r* at the free surface $u_z(r)$ results from the cumulative contribution of all the rings of dilatation at radius \tilde{r} and depth *Z* (as schematically depicted in Figure 9), the magnitude of dilation being proportional to the product of the pressure *p* and of the Biot's coefficient *b*. Based on the Green's function at the free surface $g_z(r)$ for the radially symmetric case (Equation A1b), the solution for $u_z(r)$ holds as follows:

653

$$u_{z}(r) = \frac{b \cdot 2(1+\nu)}{E} \int_{0}^{\infty} \int_{0}^{\infty} p(\tilde{r}, Z) \cdot g_{z}(r, z, \tilde{r}, Z) d\tilde{r} dZ$$
(A1)

656 where:

658 659	$g_{z}(r) = (1 - 2 \cdot \nu) \cdot \tilde{r} \cdot \frac{Z \cdot K^{3} \cdot F(K)}{4\pi \cdot (1 - K^{2}) \cdot (\tilde{r} \cdot r)^{3/2}} $ (A2)		
660	where:		
661			
662	$K^2 = 4 \cdot r \cdot \tilde{r} / (\tilde{r} + r)^2 + Z^2$ and $F(K)$ is the complete elliptic integral of the second kind.		
663	The physical properties are described in Table 1.		
664			
665	[Figure 9 about here]		
666			
667	Appendix B: General solution to the pressure build-up		
668	The radial pore pressure $p(r, t)$ evolution (Equation B1) is assessed based on the analytical		
669	solution developed by Mathias et al. (2009a) using the method of matched asymptotic		
670	expansions for the non-inertial flow of CO ₂ into a slightly compressible brine aquifer		
671	(assumed to be an open system, and the pressure corresponds to the initial one) with a vertical		
672	pressure equilibrium assumption. This solution holds as follows, distinguishing three zones:		
673	- the CO ₂ fully saturated zone (near the injector), for $x \le 2\gamma(x)$ being dependent on <i>r</i> , see		
674	Eq. B2 and γ is the ratio of the viscosity of CO ₂ μ _{CO2} to the one of the brine μ _{brine});		
675	- the diphasic zone (brine + CO ₂), for $2\gamma \le x \le 2/\gamma$;		
676	- the brine-filled zone in the far field for $2/\gamma \le x$.		
677			

$$p(x,t) = \frac{1}{2\gamma} Ei\left(\frac{\alpha \cdot x}{4 \cdot \gamma}\right) + \frac{1}{2\gamma} \left(\ln\left(\frac{\alpha \cdot x}{4 \cdot \gamma}\right) + \kappa\right)$$

$$= \left\{ \begin{array}{l} -\frac{1}{2}\ln\left(\frac{x}{2 \cdot \gamma}\right) - 1 + \frac{1}{\gamma} - \frac{1}{\gamma} \left(\ln\left(\frac{\alpha}{2 \cdot \gamma^{2}}\right) + \kappa\right) & \text{if } x \le 2\gamma \\ + \left\{ -\left(\frac{x}{2 \cdot \gamma}\right)^{0.5} + \frac{1}{\gamma} - \frac{1}{2\gamma} \left(\ln\left(\frac{\alpha}{2 \cdot \gamma^{2}}\right) + \kappa\right) & \text{if } 2\gamma \le x \le 2/\gamma \\ - \frac{1}{2\gamma} \left(\ln\left(\frac{\alpha \cdot x}{2 \cdot \gamma}\right) + \kappa\right) & \text{if } x \ge 2/\gamma \end{array} \right.$$

$$(B1)$$

678 679

680 where:

681

$$x = \frac{\left(r/r_{W}\right)^{2}}{t \cdot Q_{W} / \left(2\pi \cdot \omega \cdot H \cdot \left(r_{W}\right)^{2} \cdot \rho_{CO2}\right)}$$

$$\alpha = \frac{Q_{W} \cdot \mu_{CO2} \cdot \left(c_{rock} + c_{brine}\right)}{2\pi \cdot H \cdot k}$$
(B2)

683

- 684 where:
- 685 $\kappa \approx 0.5772$ is Euler-Mascheroni constant ; $r_{\rm W}$ is the injector well radius (m); $\rho_{\rm CO2}$ is the CO₂
- density (kg/m³); Q_W is the mass injection rate (kg/s); *t* time (s); c_{brine} is the brine
- 687 compressibility (Pa⁻¹); Other parameters are described in Table 1. Fluid properties are derived
- 688 (viscosity, density and compressibility for both brine and CO₂) based on the empirical
- 689 correlations with pressure and temperature supplied by Mathias et al. (2009b).
- 690

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810 Table caption

811

- 812 Table 1: Aquifer homogeneous properties used for the spatio- temporal analysis of the surface
- 813 vertical displacements

Aquifer property	Symbol	Value	Unit
Intrinsic permeability	k	$1.e^{-13}$	m ²
Porosity	ω	15	%
Thickness	Н	50	m
Salinity	sal	150	g/l
Rock matrix	Crock	$4.5e^{-10}$	Pa ⁻¹
compressibility			
Young's modulus	Ε	15	GPa
Poisson's ratio	υ	0.25	-
Biot's coefficient	b	1.0	-

⁸¹⁵

- 817 Table 2: Summary of the constraints on the implementation of the PSI technique and on the
- 818 sensor regarding the specificities of the deformation phenomenon associated with CO₂
- 819 injection at depth

Expected deformation		Corresponding constraints on the implementation of the PSI
characteristics		technique and on the sensor
Large spatial scales tens	_	Large swaths (>50 km) are required;
of km)		
Large vegetated areas	-	Poor PS density is expected. Additional Corner Reflectors
		could be required to complement "natural" PS network.
Long term monitoring	-	Because of limited sensors lifetimes, the use of different
(25 – 50 years)		successive satellites is required;
	_	High sensor repetitiveness (typically, ten of days or better)

⁸¹⁶

could reduce (respect to Envisat/ASAR-like configurations) the transition period between sensors by providing a minimum data archive to start the PSI processing (about 30 images) in the shortest possible duration;

- Lifetime of PS are expected to be improved on small isolated areas with high density PS allowed by the use of high resolution image data;
- Large amount of data are expected to be acquired, hence implying considering flexible and low-price data provision policies.

Spatio– temporalThe acquisition and processing rate should be adapted to theevolutionoftheexpected spatio- temporal evolution of the deformationdeformationphenomenon. Based on simulations on a typical injectionphenomenonscenario:

- the regions, where the deformation front passes, should be highly-rated sampled for a 2-4 years period;
- the changes in the velocities values of the differential displacements are expected to be detectable only after a few months, implying a processing rate of 3 to 6 months.

PotentiallysuddenTo fulfil the "early-warning" objective of the monitoring, aunexpectedgroundcomplement should be envisaged so that well definite zones atmovements on localizedrisk (such as fault zones) should be monitored using highareasrepetitiveness / high resolution (but reduced swath) image data.

820 Figure captions



Figure 1: A) Pressure build-up (Pa) within the storage aquifer for a 1Mt/y CO₂ injection
scenario into a 1500m deep aquifer (considering the aquifer properties in Table 1); B) Vertical
displacement (mm) at the surface induced by the injection at depth. Detection threshold at 5
mm (discussed in section 2.1) is outlined by a grey-coloured horizontal line.



826

Figure 2: Temporal evolution of the maximum vertical displacement u_{max} (mm) within the injector zone for a 1Mt/y CO₂ injection scenario into a 1500m deep aquifer (considering the aquifer properties in Table 1).

Figure 3: A) Maximum vertical displacement reached at the surface respectively after 1 and
10 years (top and bottom panels) considering different injection scenarios (mass injection
rates and depths) and the aquifer properties described in Table 1; B) Corresponding lateral
extent (km) of the detectable zone at the surface given a threshold of 5 mm respectively after
1 and 10 years (top and bottom panels).

- Figure 4: Schematic representation of the deformation to be measured. *t* corresponds to the
- 838 beginning of the front pass on the measurement point, Δt the duration of the front pass. Δu is
- the resulting displacement on the observation point.

841 Figure 5: main steps for the application of the PSI analysis for "verification" monitoring using 842 CR and considering the transition period between successive satellite mission (sensor 1 and 2 843 with possible different characteristics). The duration of the PSI baseline step (outlined by a 844 star-like marker) is estimated considering a minimum of 30 images with a repetitiveness of a 845 ten of days. The processing rate every a few months is estimated based on the spatio-846 temporal evolution of the deformation phenomenon associated with the typical injection 847 scenario described in section 1.2. This should be adjusted in case of the detection of a 848 "abnormal behaviour" (see discussion in section 2.4).

Figure 6: Location of the PICOREF sector (Grataloup et al. 2008) in the Paris basin context (red line). The red cross marker represents the potential injector position selected based on the risk analysis of Bouc et al. (2011). The circles represent the lateral extents of the expected detectable zones (5mm threshold) for different time durations considering a 2Mt/y injection scenario.

 $\begin{array}{c} \text{XL2E (m)} \\ \text{x 10}^5 \end{array}$ Figure 7: A) Voronoi analysis: pink coloured diamond markers represent natural PS located within the detectable zone expected for the whole duration of the injection (black circle, see right figure); blue coloured triangular markers represent potential CR at the edge of the Voronoi cells depicted by blue segments. B) Application to the PICOREF sector for the first

10-years epoch. Light blue coloured triangular markers outline the CR lying within forest
zones or at a distance inferior to 1km to the neighbouring CR. Only the CR located within the
area bounded by the light grey-coloured circle are integrated for the monitoring of the
expected detectable zone for the first 10 years of injection.

- 873 Figure 9: Schematic representation of the cumulative contribution (light blue coloured arrows)
- 874 of the rings of dilation with magnitude proportional to the reservoir pore pressure spatial
- 875 distribution p(r) to the ground displacement (red arrow) at the lateral distance r. The vertical
- 876 ground displacement $u_z(r)$ is represented by a vertical black arrow.
- 877