



**HAL**  
open science

## Validation and intercomparison of Persistent Scatterers interferometry: PSIC4 project results

Daniel Raucoules, Bernard Bourguine, Marcello de Michele, Gonéri Le Cozannet, Luc Closset, C. Bremmer, H. Veldkamp, D. Tragheim, L. Bateson, M. Crosetto, et al.

### ► To cite this version:

Daniel Raucoules, Bernard Bourguine, Marcello de Michele, Gonéri Le Cozannet, Luc Closset, et al.. Validation and intercomparison of Persistent Scatterers interferometry: PSIC4 project results. *Journal of Applied Geophysics*, 2009, 68 (3), pp.335-347. 10.1016/j.jappgeo.2009.02.003 . hal-00510374

**HAL Id: hal-00510374**

**<https://brgm.hal.science/hal-00510374>**

Submitted on 18 Aug 2010

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1                   **Validation and Intercomparison of Persistent Scatterers Interferometry:**

2                                   **PSIC4 project results.**

3  
4  
5                   Keywords: Persistent Scatterers Interferometry, validation, mining subsidence

6  
7                   Authors: Raucoules D.<sup>1</sup>(corresponding author), Bourgine B.<sup>1</sup>, de Michele M.<sup>1</sup>, Le Cozannet G.<sup>1</sup>,  
8                   Closset L.<sup>1</sup>, Bremmer C.<sup>2</sup>, Veldkamp, H.<sup>2</sup>, Tragheim D.<sup>3</sup>, Bateson L.<sup>3</sup>, Crosetto M.<sup>4</sup>, Agudo  
9                   M.<sup>4</sup> Engdahl M.<sup>5</sup>

10  
11                   <sup>1</sup> BRGM (French Geological Survey), 3 avenue C. Guillemin, BP 6009, 45060 Orléans cedex 2,  
12                   France, [d.raucoules@brgm.fr](mailto:d.raucoules@brgm.fr), phone: (33) 2 38 64 30 86; fax : (33) 2 38 64 36 89

13  
14                   <sup>2</sup> TNO - Built Environment and Geosciences (Geological Survey of the Netherlands) Princetonlaan 6,  
15                   PO Box 80015, 3508 TA Utrecht, The Netherlands

16  
17                   <sup>3</sup> BGS (British Geological Survey)  
18                   Kingsley Dunham Centre, Keyworth, Nottingham, NG12 5GG, United Kingdom

19  
20                   <sup>4</sup> IG (Institute of Geomatics)

21  
22                   Parc Mediterrani de la Tecnologia, Avinguda del Canal Olímpic, s/n , E-08860 Castelldefels,  
23                   (Barcelona), Spain

24  
25                   <sup>5</sup> ESA/ESRIN (European Space Agency)  
26                   Via Galileo Galilei, Casella Postale 64, 00044 Frascati, Italy

28

29 **1. Abstract:**

30

31

32 This article presents the main results of the *Persistent Scatterer Interferometry Codes Cross*  
33 *Comparison and Certification for long term differential interferometry (PSIC4)* project. The project  
34 was based on the validation of the PSI (*Persistent Scatterer Interferometry*) data with respect to  
35 levelling data on a subsiding mining area near Gardanne, in the South of France. Eight PSI participant  
36 teams processed the SAR data without *any a priori* information, as a blind test. Intercomparison of  
37 the different teams' results was then carried out in order to assess any similarities and discrepancies.  
38 The subsidence velocity intercomparison results obtained from the PSI data showed a standard  
39 deviation between 0.6 and 1.9 mm/yr between the teams. The velocity validation against rates  
40 measured on the ground showed a standard deviation between 5 and 7 mm/yr. A comparison of the  
41 PSI time series and levelling time series shows that if the displacement is larger than about 2 cm in  
42 between two consecutive SAR-images, PS-InSAR starts to seriously deviate from the levelling time  
43 series. Non-linear deformation rates up to several cm/yr appear to be the main reason for these reduced  
44 performances, as no prior information was used to adjust the processing parameters. Under such  
45 testing conditions and without good ground-truth information, the phase-unwrapping errors for this  
46 type of work are a major issue. This point illustrates the importance of having ground truth  
47 information and a strong interaction with the end-user of the data, in order to properly understand the  
48 type and speed of the deformation that is to be measured, and thus determine the applicability of the  
49 technique.

50

51

52 **2. Introduction**

53

54 Repeat spaceborne Interferometry is a well known technique to assess the displacement of the ground  
55 surface. It measures the displacement in the sensor's Line-Of-Sight direction from the phase of the

56 signal measured by Synthetic Aperture Radar (SAR) onboard a satellite (Goldstein et al., 1988,  
57 Massonnet and Rabaute, 1993).

58 In order to understand the limitations of the technique and to separate the different components  
59 (deformation, height, atmosphere) of the interferometric signal, a group of methods named Persistent  
60 Scatterers Interferometry (PSI) was developed in the late 1990's (the PSInSAR<sup>TM</sup> algorithm, Ferretti et  
61 al., 2001). The method aims to extract the information from radar interferometry on a large set of  
62 images by taking advantage of the large data archive acquired over the last decade (mainly with the  
63 ERS satellites, Envisat and RadarSat satellites). The main idea behind the method is that some bright  
64 radar targets retain their phase and amplitude stability for a period of months or years. The phase  
65 information of these targets (here denoted Persistent Scatterers or PS) can be exploited to interpret  
66 otherwise uncorrelated long time-scale interferograms. Other alternative methods, which use a quality  
67 indicator based on the temporal coherence on sub-sampled SAR data are restricted to small baseline  
68 pairs (Small BASelines subset technique, Bernardino et al., 2002) or to points showing both a spatial  
69 and temporal coherence (Coherent Target Monitoring, Van der Kooij, 2005). In all the cases, the  
70 objective is to statistically separate the different components of the interferometric phase of points,  
71 where the phase is reliable for the whole data set, to provide a precise assessment of the ground  
72 surface deformation.

73 End-users outside the radar community have little experience in utilizing the products delivered with  
74 PSI methods. Thus, some concerns about PS data quality, trustworthiness and how to interpret them  
75 arose. This issue was clearly identified during the 2003 FRINGE Workshop  
76 ([http://earth.esa.int/fringe03/Fringe\\_reco3.pdf](http://earth.esa.int/fringe03/Fringe_reco3.pdf)). ESA (European Space Agency) which decided to  
77 initiate this validation project, named PSIC4 (Persistent Scatterer Interferometry Codes Cross  
78 Comparison and Certification for long term differential interferometry), in order to assess the  
79 performances of PSI for land deformation monitoring.

80 Eight PSI teams, from business or academic organizations, participated in the project: Altamira  
81 Information (Crosetto et al., 2008), DLR (German Space Agency, Adam et al., 2003), Gamma-RS  
82 (Werner et al., 2003), IREA-CNR (Institute for Electromagnetic Sensing of the Environment National  
83 Research Council of Italy, Bernardino et al., 2002), TRE (TeleRilevamento Europa, Ferretti et al.,

84 2007), TUDelft (Delft University of Technology, Kampes, 2005), UPC (Catalonia Polytechnics  
85 University, Mora et al, 2003), and Vexcel (Van der Kooij et al., 2005). 107 ERS and 10 Envisat ASAR  
86 images were delivered to each of the teams by ESA for them to produce their own PSI deformation  
87 products. Their results were then analysed by an independent validation consortium (BRGM, BGS,  
88 TNO, IG) as an anonymous procedure so the validation consortium received the data without any  
89 identification, except for a random number assigned to each one of the teams.

90 Past validation tests (e.g. Ferretti et al. 2007) were based on controlled displacements of corner  
91 reflector points perfectly identifiable as a PS on the SAR data set, which proved the indisputable  
92 precision (better than the millimetre) of the technique in ideal conditions. In contrast, this test aimed to  
93 address issues more oriented towards the end-user who plans to use PSI products for helping in the  
94 management of deformation hazards in a real operational situation. In particular, there is no a priori  
95 coincidence between the ground measurement points and the PS. A similar context is presented in  
96 Casu et al. 2006. The issue is addressed in section 4.1 based on the spatial characteristics of the  
97 deformation field.

98 This current work aims to give an insight on the information that the end user can expect from PSI.  
99 For this purpose, the project has to address some of the following key questions:

100 - How well does PSI describe the land deformation field, spatially and temporally?

101 - How accurately, and how precisely, does PSI describe the land deformation field?

102 - How consistent are the PSI results between the different teams?

103

104 These questions are addressed by the validation and intercomparison activities, where the results of  
105 each team are compared both against reference ground-based data and against each other in a series of  
106 tests designed to give an extended view of the performance of the PSI methods.

107

108 This paper describes the work performed during PSIC4 activities. In this project, six of the eight PSI  
109 teams used different implementations of the Persistent Scatterers technique, while the other two teams  
110 used alternative concepts (basically two coherent-based approaches). In this paper, the eight PSI teams

111 were asked to process the SAR data according to their own PSI methodology without any *a priori*  
112 information about the type and location of deformation in the chosen test area. In addition to this  
113 paper, an extensive description of the project is available in the PSIC4 final report (Raucoules et al.,  
114 2007).

115

### 116 **3. Test site**

117

118 The test area was selected by ESA from several proposed by the validation group, who considered not  
119 only the characteristics of the deformation (size, rate) but most importantly, the availability and quality  
120 of levelling data for the 1992-2004 period.

121 The area of interest is located near the town of Gardanne (Bouches-du-Rhone, France) in the  
122 sedimentary basin of “l’Arc” between the Sainte-Victoire Mountain and the cities of Marseille and  
123 Aix-en-Provence (figure 1). The area of Marseille-Aix is the second biggest urbanized area in France  
124 (with more than 1.4M inhabitants).

125

126 **[insert figure 1]**

127

128 Figure 1: Gardanne mining area. The location of the exploited panels is shown.

129

130

131

132

133

134 The coal field (lignite) here has been mined since the Middle Ages Exploitation stopped in February  
135 2003. Different mining techniques were used in the area. For the period of interest (1991-2004), the  
136 observed ground subsidence is associated with the coal mining exploitation using long wall mining  
137 technique at nearly 1000 m below the surface (figure 2). This technique has particular consequences in  
138 terms of size (the width of a panel is 250m) and evolution (almost immediate) of the deformation

139 compared to older mining techniques. The older methods include chamber and pillar extraction with  
140 localised subsidence occurring over a longer time-scale.

141 Typically, when a long-wall panel is mined, half of the displacement (hundreds of metres wide) is  
142 produced in the first 2 months with the residual deformation occurring over the following 3 years  
143 (Arcamone, 1980). A point on the surface may be influenced by several mined panels extracted at  
144 different periods causing the deformation to last for longer periods.

145

146

147

148

149

150 [insert figure 2]

151

152 Figure 2: Detailed plan of the underground mine panels. Isolines indicate the depth of the exploitation  
153 panel.

154

155

156

157

158

159

160 The monitoring of land deformation effects associated with coal mining exploitation of the Gardanne  
161 area was performed through spirit levelling surveys by the French coal mining authority (CDF –  
162 *Charbonnages de France*).

163 The *in situ* data of the Gardanne test site have been acquired at more than 1000 points over the past  
164 decades in order to monitor the deformation effects on the surface associated with the coal mining.

165

166

167 [insert figure 3]

168

169 Figure 3: Gardanne mining area, levelling network . This map shows the names and locations of the  
170 levelling lines and the positions of the mining works (in gray).

171

172

173 Since 1990, levelling surveys have been carried out with an automatic electronic level (Leica Wild  
174 NA3003) whose bar code specified precision performances (Dommanget J.M., 2004) are:

175

176 Standard deviation of height measure error, measured point-to-point =  $\pm 0.7$  mm

177 Standard deviation of height error on a 1 km one-way levelling =  $\pm 1.5$  mm

178

#### 179 **4. Pre-processing**

180

##### 181 4.1 Previous corrections of the data

182

183 Although we tried hard to not modify the PSI data (in order not to add involuntary biases to the test),  
184 we observed two main problems with the data needing modifications: shifts in the geolocation and  
185 biases on the estimated velocities.

186 The only way to carry out the data comparison was to adjust those shifts to make the data comparable.

187 We observed geocoding shifts by overlaying the PSI data on an ortho-rectified (in the local Lambert  
188 III Sud cartographic projection) aerial photograph. The geocoding discrepancies ranged between 5m

189 and 80m depending on the team. After correction, the residual geocoding shifts estimated on other

190 control points were between 3m and 23m. It is important to note that these residual errors include the

191 error made in the identification of the radar features on the ground. A linearly varying shift would

192 probably have provided a better final geolocation. However, we assumed that the constant shift we

193 applied to the data was sufficient for the purpose of this study.



194 More generally, we emphasize the fact that for many applications where the deformation is of very  
195 small extent (e.g. shrink-swelling, cavity collapses, small landslides etc.), the geocoding issue is very  
196 important for a potential end-user. It is essential to produce data that can be overlaid with sufficient  
197 accuracy on reference maps to allow a correct interpretation. During the processing an interaction with  
198 the end-user is recommended.

199 The fact that the different teams did not chose the same reference point for deformation computation  
200 produced biases on the assessment of the deformation, making direct comparison impossible. We  
201 carried out a stable area adjustment based on the levelling information and the knowledge of the end-  
202 user (CDF) and chose the stable area outside the influence of the mining works. The objective is to be  
203 sure that all the PSI datasets are referenced against the same stable area as the levelling network in  
204 order to have comparable data. In a general case, the choice of a reference stable area can be critical  
205 for the measurement. In operational use we think that the reference should be selected in agreement  
206 with the end-user in order to better respond to his needs.

207 Once we selected the stable area (figure 4), we calculated the average deformation rate value for those  
208 PS included in the stable area and for all the teams. In this area, where the velocity should be zero and  
209 the time series should be flat, PS average velocity and time series highlighted either a subsidence or an  
210 uplift. We assume that the non-zero velocity values can be considered as a bias on the velocities  
211 affecting the whole dataset due to bad reference choices. We therefore used the computed average  
212 deformation rate values to force both the PS average velocities to be zero and time series to be flat  
213 within the stable area. In practice, we first corrected the average velocity estimations on the whole  
214 PSI dataset by subtracting the calculated stable area values from the dataset. Secondly, we corrected  
215 the time series by removing a linear trend derived from the previous velocity correction estimation.

216

217

218 [insert Figure 4]

219 Figure 4: The selected stable area (white frame) is located around the five stable levelling points  
220 identified by their codes. The triangles correspond to the levelling points, circles to PS provided by the  
221 team T4.

222

223

#### 224 4.1 Interpolation of the data

225

226 The major issue for the comparison between levelling data and the PSI results comes from the fact that  
227 the spatial and temporal sampling of the PS-field and Time Series and the levelling lines do not  
228 coincide. The levelling frequency is about twice a year whereas the PSI data correspond to the sensor  
229 acquisition rate which is on average every 40 days (taking into account a gap in the acquisitions during  
230 1994). The number of levelling points are fewer than the number of PSI points. The objective of  
231 interpolation is to define which value derived from the SAR data can be compared with a given value  
232 derived from the levelling.

233 The option we applied in this study was the following. We interpolated temporally the levelling data to  
234 the SAR acquisition dates and we interpolated spatially the PSI deformation values to the levelling  
235 point positions (using a limited radius of 50m around the levelling points).

236 The basic justification is that the studied phenomena have sufficient spatial extents to justify a spatial  
237 interpolation in a 50m radius. In fact, in this area, the subsidence bowls due to the mining works  
238 spread over hundreds of metres (Arcamone, 1980).

239 Now, with the selected procedure, each levelling point with given geographic coordinates can be  
240 associated to the deformation Time Series corresponding to the eight participant teams and to the  
241 levelling measures.

242 Our choice for the temporal interpolation of the levelling instead of the SAR data was decided in order  
243 to avoid excessive smoothing of the time series. Smoothing could hamper the assessment of the  
244 quality of the PSI time series. An additional reason is that, due to the nature of the levelling dataset,  
245 which is the compilation of 17 lines measured independently at different dates and frequencies, we  
246 were not able to define a unique temporal sampling for the levelling.

247 All the Time Series (SAR and interpolated levelling) were set to '0 deformation' using 15/07/1992 as  
248 the reference date. This date corresponds to the oldest image used by all the teams.

249 We interpolated the levelling by ordinary kriging. This allowed us to retrieve an interpolation error.  
250 So, for the PSIC4 exercise, we kept only the values with an interpolation error lower than 3.5mm.  
251 For the spatial interpolation, because the PSI results were massively oversampled with respect to  
252 levelling, a spatial kriging computation was performed for each image with a 50 metres search  
253 neighbourhood in order to reject possible variability on longer distances.

254

255

## 256 **5. Intercomparison**

257

258 The intercomparison activities aimed to identify relative differences between the teams. The initial  
259 objective of the exercise was to check if the different methods provided equivalent results and to  
260 assess any discrepancies. Among the tested indicators, the most relevant are: 1) the average  
261 deformation rate; 2) the density and distribution of the PS.

262

263 1) Average deformation rates.

264

265 To estimate the discrepancies between the teams' results in terms of velocity maps, we resampled the  
266 data included in a common area to a 50m by 50m grid containing the velocities of each of the teams.

267 Using the ISATIS<sup>TM</sup> software, we carried out a comparison by pairs of teams: for each pair of  
268 produced PSI sets, we assessed the mean of the differences, the correlation value and the standard  
269 deviation of the differences on the cell occupied by the two compared teams.

270 The main indicator is the standard deviation that reflects the discrepancies between teams. It ranges  
271 from 0.6 mm/yr to 1.86 mm/yr.

272 The mean of the differences can show possible biases, although these values have been partially  
273 affected by the stable area correction and are therefore less relevant as a performance indicator. These  
274 values range from -0.84 mm/yr to +0.44 mm/yr. The following table shows the full set of velocity  
275 intercomparison values.

276



277

278

<b>Team</b>	<b>T1</b> (889pts)	<b>T2</b> (5100pts)	<b>T3</b> (4169pts)	<b>T4</b> (6552pts)	<b>T5</b> (2511pts)	<b>T6</b> (1269pts)	<b>T7</b> (6360pts)	<b>T8</b> (17081pts)
<b>T1</b>		<b>1.57 (0.79)</b>	<b>0.85 (0.85)</b>	<b>0.88 (0.85)</b>	<b>1.23 (0.83)</b>	<b>0.84 (0.72)</b>	<b>0.96 (0.83)</b>	<b>1.44 (0.85)</b>
<b>T2</b>	<b>-0.26</b>		<b>0.87 (0.89)</b>	<b>1.06 (0.85)</b>	<b>1.19 (0.86)</b>	<b>0.99 (0.71)</b>	<b>1.01 (0.85)</b>	<b>1.86 (0.73)</b>
<b>T3</b>	<b>-0.14</b>	<b>0.12</b>		<b>0.63 (0.94)</b>	<b>0.87 (0.91)</b>	<b>0.73 (0.83)</b>	<b>0.71 (0.90)</b>	<b>1.40 (0.72)</b>
<b>T4</b>	<b>-0.46</b>	<b>-0.13</b>	<b>-0.24</b>		<b>1.01 (0.89)</b>	<b>0.74 (0.81)</b>	<b>0.76 (0.89)</b>	<b>1.46 (0.72)</b>
<b>T5</b>	<b>0.11</b>	<b>0.27</b>	<b>0.20</b>	<b>0.44</b>		<b>1.08 (0.71)</b>	<b>0.99 (0.87)</b>	<b>1.71 (0.73)</b>
<b>T6</b>	<b>-0.34</b>	<b>0.06</b>	<b>-0.05</b>	<b>0.20</b>	<b>-0.21</b>		<b>0.81 (0.77)</b>	<b>1.06 (0.64)</b>
<b>T7</b>	<b>-0.32</b>	<b>0.02</b>	<b>-0.1</b>	<b>0.14</b>	<b>-0.32</b>	<b>-0.10</b>		<b>1.33 (0.72)</b>
<b>T8</b>	<b>-0.82</b>	<b>-0.41</b>	<b>-0.56</b>	<b>-0.25</b>	<b>-0.84</b>	<b>-0.47</b>	<b>-0.40</b>	

279

280 Table 1: intercomparison of velocities. The upper right triangle contains for the teams corresponding  
281 to the rows and columns, the standard deviation of the velocity differences and the correlation (in  
282 parenthesis). The lower triangle contains the mean of the velocity differences. The number of cells  
283 occupied by each team is indicated in the first row of the table.

284

285

286 2) Density and spatial distribution of the PS

287 PS distributions show large variations. For example, some teams have no PS selected in the  
288 deformation area whereas team T8 seems to have a larger density in those areas. Such differences are  
289 probably the consequence of the use of different coherence thresholds on the selection of the points.

290 The figures 5 and 6 show the two extreme cases observed: team T6 rejected practically all the points  
291 of the deformed area and team T8 kept a large density of points.

292 [insert figure 5]

293 Figure 5: density of PS for Team 6. The location of the main deformation (derived from  
294 conventional interferometry) is showed (blue contour).

295 [insert figure 6 ]

296

297 Figure 6: density of PS for Team 8. the location of the main deformation (derived from conventional  
298 interferometry) is showed (blue contour).

299

## 300 **6. Validation results**

301

302 The validation in this study is the comparison of the interpolated PSI data with the corresponding  
303 levelling measurement.

304 The first test is a semi-quantative comparison. We selected the more representative levelling line  
305 (“AXE” – located on figure 3 - it crosses the main deformed area) and visually compared the PSI data  
306 versus the levelling data along it. Figure 7 shows the spatial variation of the cumulative deformation  
307 within the period from 1992 to 1998 for each of the teams.

308 Most of the teams (except team T6) spatially localised the deformation. We observed dissimilarities  
309 with the levelling and between PSI teams. In particular, the PSI results seem to underestimate the  
310 higher deformations.

311

312

313 [insert figure 7]

314

315 Figure 7a): cumulative deformation between 06/05/1992 and 31/10/1998 along the “AXE” levelling  
316 seen by the different PS Teams and the levelling versus distance to first point. The frame corresponds  
317 to section magnified in figure 7b). We can observe important discrepancies in the area of higher  
318 deformation.

319 Figure 7b): Zoom of the profiles along “AXE” line in a section with moderate (less than 3cm)  
320 deformation. This gives a first insight on the relative variability between the different teams. The PSI  
321 results globally follow the levelling profile but with fluctuations (respect to levelling and other teams)  
322 of about 10-15 mm; a more precise estimation is given below.

323

324

325

326 The RMSE (root mean square error) of the PSI time series against the levelling time series are  
 327 reported in table 2. The values are represented by average velocity classes (estimated on levelling).  
 328 We can observe a dependence of the RMSE with respect to the velocity value (figure 8). The values  
 329 range on average from 1.5 cm for the lower velocity class (less than 5 mm/yr) to 10 cm (for the larger  
 330 movements (more than 15 mm/yr).

331

332

<b>velocity (mm/yr)</b>	<b>1</b>	<b>2</b>	<b>3</b>	<b>4</b>	<b>5</b>	<b>6</b>	<b>7</b>	<b>8</b>
>0	12	10	9	11	14	14	9	20
[0; -5]	18	14	14	18	14	22	9	22
] -5; -10]	38	41	38	46	32	58	25	27
] -10; - 15]	108	97	90	85	86	86	75	69
<-15	93	97	151	112	106	85	99	91

333

334 Table 2. Average RMSE (mm) per velocity class and per processing chain

335

336

337

338

339 [insert figure 8]

340 Figure 8: RMSE per velocity class and per processing chain

341

342 A last indicator consists of comparing the estimation of the velocity obtained by PSI versus the  
 343 levelling. The following figures show the velocity values derived by linear regression on the  
 344 deformation time series, from both the levelling and the PSI on the location of the levelling points.



345 This allows a visual examination of the discrepancies in the area of major deformation (in agreement  
346 with the previous observations).

347

348

349 [insert figure 9]

350 Figure 9: Vertical velocities estimated from levelling data resampled at SAR acquisition dates

351

352

353

354

355 [insert figure 10]

356 Figure 10. Vertical velocities derived from PSI resampled at levelling points location. Teams T1-T4.

357

358

359 [insert figure 11]

360 Figure 11. Vertical velocities derived from PSI resampled at levelling points location. Teams T5-T8.

361

362 Table 3 gives a quantitative assessment of the discrepancies. In the conditions of the experiment  
363 (characteristics of deformation and land-use) the standard deviation of the velocity differences is about  
364 6 mm/yr.

365

366

367

368

<b>Team</b>	<b>Number of points</b>	<b>Mean of</b>	<b>Standard deviation</b>
-------------	-------------------------	----------------	---------------------------

		( Ti – Levelling )	of (Ti – Levelling)
<b>T1</b>	158	2.8	5.3
<b>T2</b>	478	4.4	7.2
<b>T3</b>	328	3.9	6.6
<b>T4</b>	447	3.7	6.3
<b>T5</b>	348	3.6	6.8
<b>T6</b>	136	3.2	5.1
<b>T7</b>	417	2.6	5.7
<b>T8</b>	817	4.7	6.7

369 **Table 3:** Overall differences between PS and levelling velocities (in mm/yr): mean and standard  
370 deviation of difference between PSI and levelling.

371

372 To complete this comparison we have to highlight that the number of points associated with levelling  
373 is very variable between the teams (158 to 817 in the same area). For instance, team T8 produced  
374 points on most of the levelling line lengths. Therefore, the discrepancy with levelling is more affected  
375 by the effect of the underestimation of the higher deformations. The quantitative comparison is then  
376 less favourable to those who provided points in the higher deformation area, but they provided a better  
377 qualitative description of the deformation.

378

379

## 380 **7. Discussion**

381

382

383 • The PSIC4 exercise was conceived as an inter-comparison and validation of PSI data computed by  
384 eight different teams. The aim of this exercise was to evaluate the performance of eight different PSI  
385 methodologies by comparing the results with ground based observations and by inter-comparing the  
386 results from the eight processing chains. The results of this study provide an assessment of the

387 absolute (validation) and relative (inter-comparison) performance of the PSI techniques. The PSIC4  
388 exercise is a blind test, carried out on a mining site characterised by different magnitudes and  
389 evolutions of deformation and variable land cover. It is worth emphasising that the PSIC4 project  
390 represents a unique experiment for the number of PSI teams involved and for the quality of the  
391 available ground truth, which involve more than 1000 levelling benchmarks measured once or twice  
392 per year.

393 However, it is important to underline that the deformations of the PSIC4 test site, especially those of  
394 the mining area of Gardanne, where the majority of the ground truth are located, represent a difficult  
395 case for the PSI techniques. In fact, the deformations range from a few centimetres up to some  
396 decimetres, and most of the deformation occurs in a few months, a rather short period for the (at most)  
397 monthly SAR acquisitions.

398 Therefore, in PSIC4 the performances of PSI techniques were tested at the very edge of their  
399 capability, as the PS interferometry processing is performed in the less favourable conditions and is  
400 evaluated upon the strongest criteria.

401

402 When going through the results of the PSIC4 exercise, one should keep in mind the following issues in  
403 order to properly interpret the outcome of the study:

404

405 1. The eight teams had no knowledge of the type of deformation occurring on the test site, i.e.  
406 information such as the linearity/non-linearity of the deformation, the driving mechanism, the location  
407 of the deformation, when it started and when it ended, and the expected deformation magnitude. In  
408 general, the teams received no information about the deformation signal of interest and. the goal of the  
409 PSI analysis was not clearly specified. In contrast, the validation was focused on a specific  
410 deformation phenomenon, i.e. the deformation associated with the mining area of Gardanne. This  
411 point is important because PSI data processing has different parameters that can be adjusted for a  
412 specific application goal. For instance, the processing can be modified to take into account a priori  
413 knowledge of fast displacements.

414

415 2. The PS measurements were evaluated against the strongest criteria. For the first time the validation  
416 of PSI against levelling data was performed quantitatively to the millimetre level. Some of these  
417 criteria can be non relevant in specific applications.

418

419 The results show that the PS technique is not invalidated, but the outcomes of the PSIC4 project  
420 should be used to improve PS interferometry performances for the critical application cases.

421

422 • The main indicators investigated during the project were the following:

423

424 *Time Series validations.* A comparison of the spatio-temporal profiles of the levelling data and PSI  
425 data along two levelling transects show that, for this case, the stretch along the line experiencing most  
426 subsidence was not well sampled by seven out of eight teams and an underestimation of subsidence  
427 velocity was shown by all teams.

428 A comparison of the PSI time series and levelling time series shows that if the displacement is larger  
429 than about 2 cm in between two consecutive SAR-images, PS-InSAR starts to seriously deviate from  
430 the levelling time series. Since, for the Gardanne site, a large number of the levelling time series show  
431 large displacements for two consecutive SAR-acquisition dates (35% in excess of 2.8 cm and 70% in  
432 excess of 1.4 cm) validation results are therefore negatively biased. This also explains the low number  
433 of cases for which PS-InSAR time series and levelling time series could be tested to belong to the  
434 same population. If one only compares those time series having a maximum displacement of 1.4 cm or  
435 less for two consecutive SAR-acquisition dates, then the current study *shows average RMSE between*  
436 *levelling time series and PSI time series of 7 mm to 25 mm.* One has to compare this with validation  
437 studies performed on artificial scatterers which show standard errors of the time series of 2 mm. From  
438 this it can be concluded that for those locations for which phase unwrapping ambiguities do not exist,  
439 at least some of the processing chains obtain results in line with previous studies, which mainly took  
440 place under controlled circumstances. In all cases reviewed, the team's results did underestimate the  
441 subsidence rate in areas showing moderate to fast subsidence. The main reason suggested for this is  
442 the character of the subsidence process in the study area. As a result of mining activities in this area,

443 the subsidence takes place over a relatively short time-span. The strong correlation between RMSE  
444 and magnitude of displacement for two consecutive SAR-acquisition dates, suggests that the results  
445 have certainly been affected by phase unwrapping ambiguities, leading to a systematic  
446 underestimation of the subsidence rates.

447

448 *Velocity validation.* The comparison of the PS-InSAR velocity with the ones derived from levelling  
449 shows an average absolute difference with standard deviations between 5 and 7 mm/yr. Again, the  
450 standard deviations are strongly dependent on the absolute value of the actual displacement of the  
451 measured point. It can reach more than 15 mm/yr on the main deformed area but generally less than 2  
452 mm/yr on stable levelling points.

453

454 *Spatial distribution intercomparison.* The highest densities related to urban areas, where many  
455 scattering objects exist. Surprisingly, some teams were able to find points in areas of forest or  
456 agriculture (Team T8 in particular). Some teams (such as Teams T1 and T8) did succeed in finding PS  
457 within a rapid ground motion zone in urban areas. Other teams were not able to identify as many  
458 points in these areas. The PSIC4 exercise shows that for the case under consideration, the main area of  
459 subsidence could not, or could only partly, be assessed and identified by seven out of eight teams. The  
460 main reason for this has been the low density of Persistent Scatterers in the area of interest thus  
461 masking the actual subsidence bowl. Nonetheless, improvements are possible taking into account that  
462 one team did find a high distribution of Persistent Scatterers within the subsidence area.

463

464 *Velocity intercomparison.* Velocity is the basic deformation parameter derived from PSI techniques, as  
465 it is obtained by assessing a linear regression on the phase history. A very high precision was therefore  
466 expected. However, the inter-comparison results show discrepancies, in terms of standard deviation,  
467 between 0.6 to 1.9 mm/yr.

468

469 *Geocoding comparison.* Significant differences occur between teams, with magnitudes of ‘average  
470 geocoding difference’ between 6 and 80 metres before correction. Improvements can be considered  
471 and a better use of prior cartographic information (such as high resolution ortho-photos) might help.

472

473 To conclude this discussion, we will present some key outcomes resulting from the responses of the  
474 participant teams to open questions addressed after the analysis of the PSI products.

475

476 1) One of the most important conclusions of the project concerns the characteristics of the mining test  
477 site of PSIC4, which include abrupt nonlinear movements with magnitudes that range from a few  
478 centimetres up to some decimetres. These are severe characteristics from the viewpoint of C-band PSI  
479 with the temporal acquisition capabilities of ERS and Envisat. Why are the deformation characteristics  
480 so important? Because in principle, PSI can measure surface displacements with millimetric precision,  
481 but this can only be achieved under the following conditions:

- 482     ▪ The right model to describe the deformation is adopted. This is difficult to accomplish with  
483     abrupt nonlinear movements.
- 484     ▪ The aliasing due to low PS density and/or low temporal sampling with respect to deformation,  
485     which may cause phase unwrapping errors, is controlled. This is difficult or impossible with  
486     strong deformation magnitudes.
- 487     ▪ The assumptions to separate the atmospheric contribution from deformation are correct. This  
488     typically fails in presence of nonlinear motion.

489 Most of the results of PSIC4 can be understood in the context of the above conditions: none of them is  
490 fully accomplished in the mining area of Gardanne in the context of this study.

491 2) It is worth underlining that the above consideration of “strong deformations” holds for C-band PSI  
492 with the current temporal acquisition capabilities of ERS and Envisat. They cannot in principle be  
493 extended to other types of SAR missions based on different bands and more frequent SAR acquisition  
494 capabilities.

495 3) The PSIC4 project was conceived in a specific framework, where the teams worked under “blind  
496 conditions”, with no a priori information on deformation type, driving mechanism, deformation  
497 magnitude, etc. Furthermore, they had no information about the exact deformation signal of interest,  
498 i.e. the goal of the PSI analysis. By contrast, the validation was focused on a specific deformation  
499 phenomenon, *i.e.* the deformation associated with the mining area of Gardanne. This point is important  
500 because it played a key role in the PSI processing. In fact, instead of tailoring the processing to a  
501 specific objective of the analysis, the teams used a standard approach and processing which is feasible  
502 with reasonable resources. It is worth emphasising that the area covered by most of the teams is  
503 considerably larger than the 100 km<sup>2</sup> area used for the validation. None of the PSI teams has  
504 performed an advanced or refined PSI analysis, because neither the area of interest nor the goal of the  
505 refinement was defined. This again explains most of the PSIC4 results, e.g. the lack of PS in the  
506 mining area “of interest”.

507 4) It is worth analysing a specific consequence of the above point, which explains the different  
508 densities achieved by the teams. The PS densities are different because there is no “definition” of what  
509 exactly is a PS. The teams used their standard PSI approach (instead of an advanced or tailored one)  
510 and delivered the PS only where both velocity and time series could be extracted with reasonable  
511 reliability. Unfortunately the validation area represents a difficult area, where phase unwrapping errors  
512 represent the main problem. Due to the high probability of this type of error, many teams did not  
513 deliver reliable information. Note that this did not occur outside the mining area, *i.e.* in the great  
514 majority of the covered areas, see point six.

515 5) Considering the above points and the results achieved in the Gardanne mining area we can say that  
516 PSIC4 has clearly demonstrated that the PSI limitations are real, *i.e.* that PSI is not applicable  
517 everywhere. Though this was already clear to many PSI specialists, this evidence has now been widely  
518 documented.

519 6) To conclude, the limitations of PSI over deformation areas with similar characteristics to Gardanne  
520 open a series of important issues for the future. The first one is the importance of a feasibility study  
521 before running a PSI analysis. This may help in avoiding false expectations and disappointing results.

522 A second issue concerns the appropriate ways to inform the PSI users of the limitations of the  
523 technique, especially in those cases where PSI is employed under the non-ideal conditions. Then it is  
524 interesting to investigate the possibility of using alternative techniques to PSI, like DInSAR which  
525 could provide useful information in difficult applications like mining areas.

526

## 527 **8. Acknowledgement**

528

529 This study was carried out in the frame of the PSIC-4 project funded by the European Space Agency.

530 The authors wish to thank Charbonnages de France for their help in this project.

531 For D. Tragheim and L. Bateson, this paper is published with permission of the Executive Director,

532 British Geological Survey.



533 **9. References**

534

535 ADAM N., KAMPES, B. M., EINEDER, M., WORAWATTANAMATEEKUL, J. and KIRCHER,  
536 M., 2003, The development of a scientific permanent scatterer system, Proceedings of ISPRS  
537 Workshop on High Resolution Mapping from Space, Hannover, Germany

538

539 ARCAMONE, J., 1980, Méthodologie d'étude des affaissements miniers en exploitation totale et  
540 partielle – application au cas des houillères de Provence, PhD thesis, Institut National Polytechnique  
541 de Lorraine, France.

542

543 BERARDINO, P., FORNARO, G., LANARI R. and SANSOSTI E., 2002, A new algorithm for  
544 surface deformation monitoring based on small baseline differential SAR interferograms, IEEE  
545 Transactions on Geoscience and Remote Sensing, 40, 2375-2383

546

547 CASU, F., MANZO, M. and LANARI, R., 2006, A quantitative assessment of the SBAS algorithm  
548 performance for surface deformation retrieval from DInSAR data, *Remote Sensing of Environment*, 1,  
549 20-28

550

551 CROSETTO M., BIESCAS E. and DURO J., 2008, Generation of advanced ERS and Envisat  
552 interferometric SAR products using the stable point network technique, *Photogrammetric Engineering*  
553 *And Remote Sensing*, 4, 443-450

554

555 DOMMANGET J.M., 2004, Gardanne levelling data specifications, personal communication

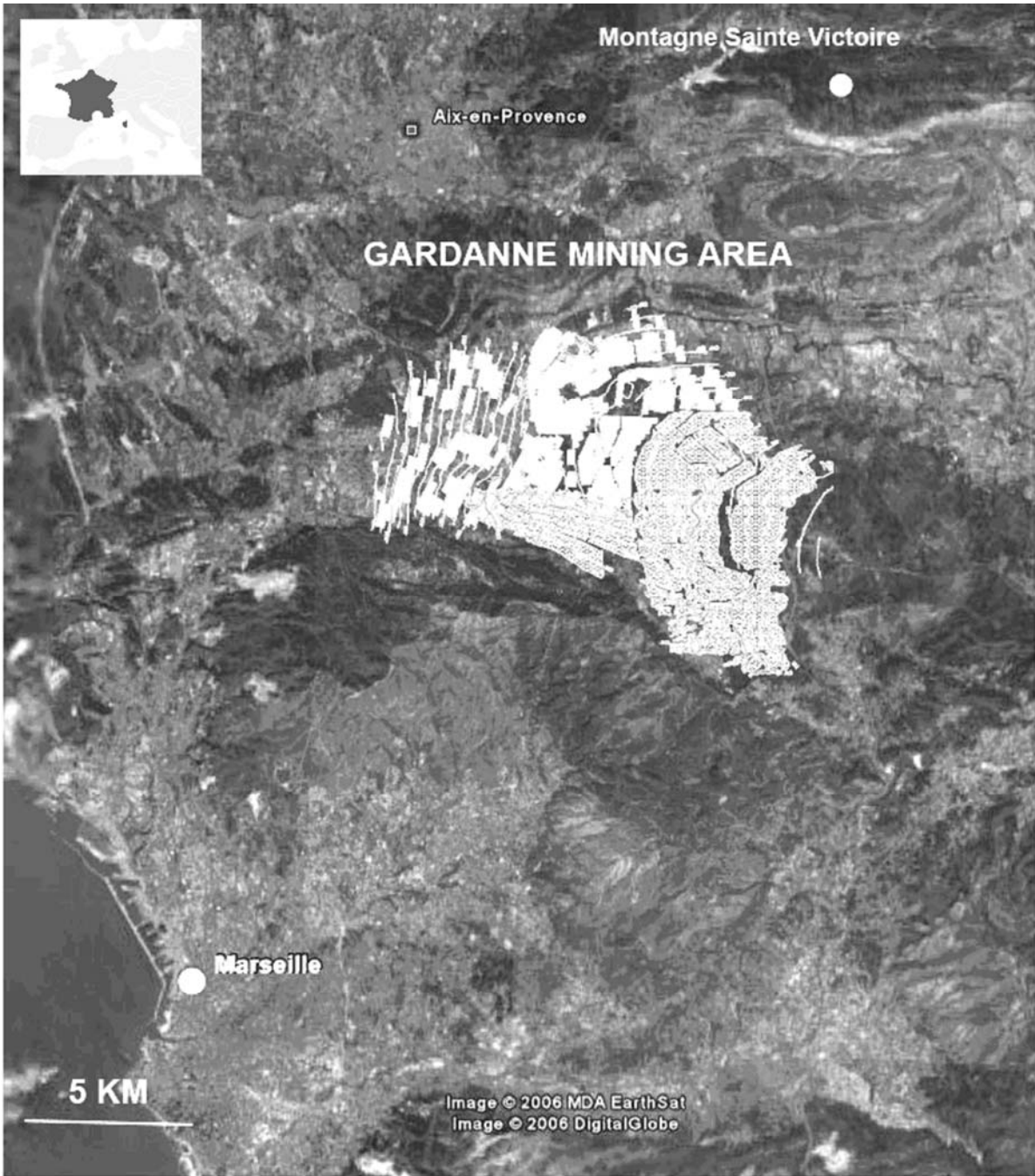
556

557 FERRETTI, A., PRATI, C. and ROCCA, F., 2001, Permanent scatterers in SAR interferometry, IEEE  
558 Transactions on Geoscience and Remote Sensing, 39, 8-20

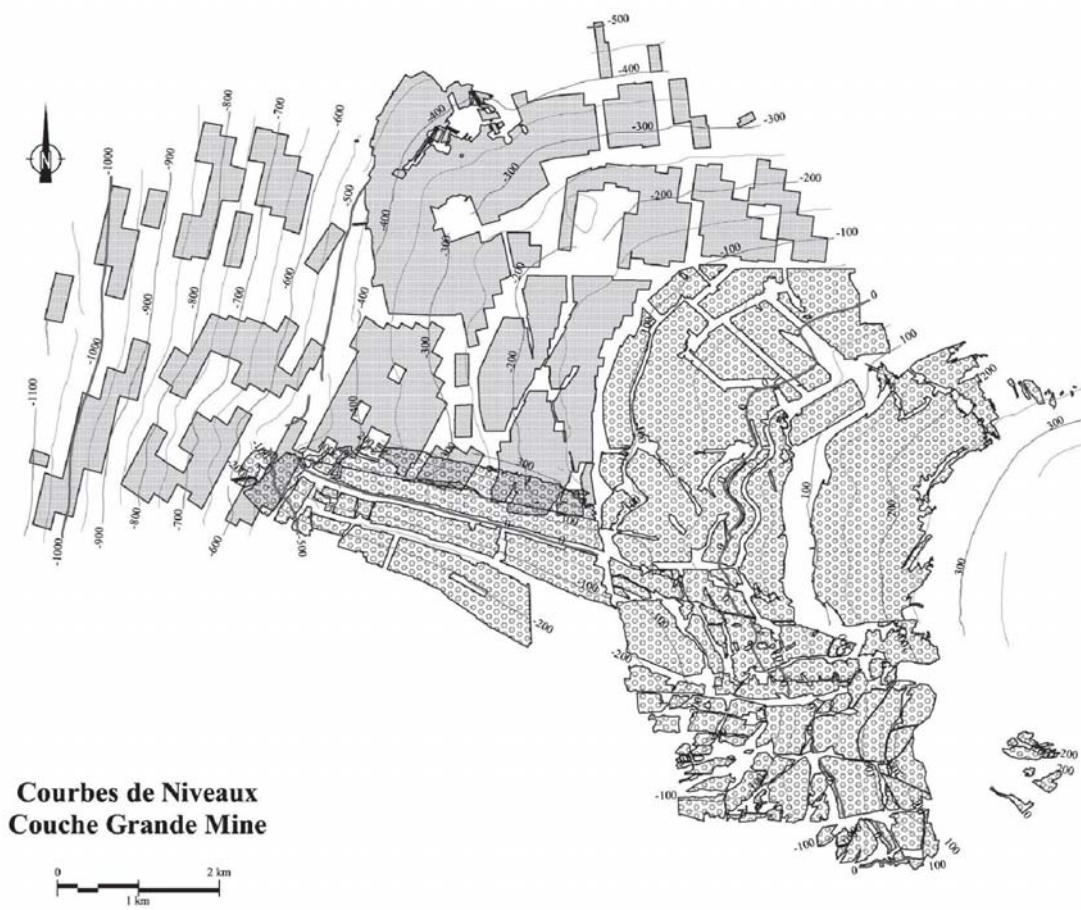
559

560 FERRETTI, A. ,SAVIO, G., BARZAGHI, R. , BORGHI, A., MUSAZZI, S. , NOVALI, F. , PRATI,  
561 C. and ROCCA, F., 2007, Submillimeter accuracy of InSAR time series: Experimental validation,  
562 IEEE Transactions on Geoscience and Remote Sensing, **45**, 1142-1153  
563  
564 GOLDSTEIN, R., ZEBKER, H. and WERNER, C., 1988, Satellite radar interferometry two-  
565 dimensional phase unwrapping, Radio Science, *23*, 713-720  
566  
567 KAMPES B. M., 2005, Deformation parameter estimation using permanent scatterer interferometry.  
568 PhD thesis. Delft University of Technology, Delft  
569  
570  
571 LANARI, R., MORA, O. AND MANUNTA, 2004, A small-baseline approach for investigating  
572 deformations on full-resolution differential SAR interferograms. IEEE Transactions on Geoscience  
573 and Remote Sensing, *7*, 1377-1386  
574  
575  
576 MASSONNET, D. and RABAUTE, T., 1993, Radar interferometry, Limits and potential  
577 IEEE Transactions on Geoscience and Remote Sensing, *31*, 455-464.  
578  
579 MORA O., MALLORQUI J.J. and BROQUETAS A., 2003, Linear and nonlinear terrain deformation  
580 maps from a reduced set of interferometric SAR images, IEEE Transactions On Geoscience And  
581 Remote Sensing, *10*, 2243-2253  
582  
583 RAUCOULES, D., BOURGINE, B., DE MICHELE, M., LE COZANET, G., CLOSSET, L.,  
584 BREMMER, C., VELDKAMP, H., TRAGHEIM, D., BATESON, L., CROSETTO, M., AGUDO, M.,  
585 and ENGDAHL, M., 2007, available online at [http://earth.esa.int/psic4/PSIC4\\_D9\\_final\\_report.pdf](http://earth.esa.int/psic4/PSIC4_D9_final_report.pdf)  
586 (accessed August 2007) .  
587

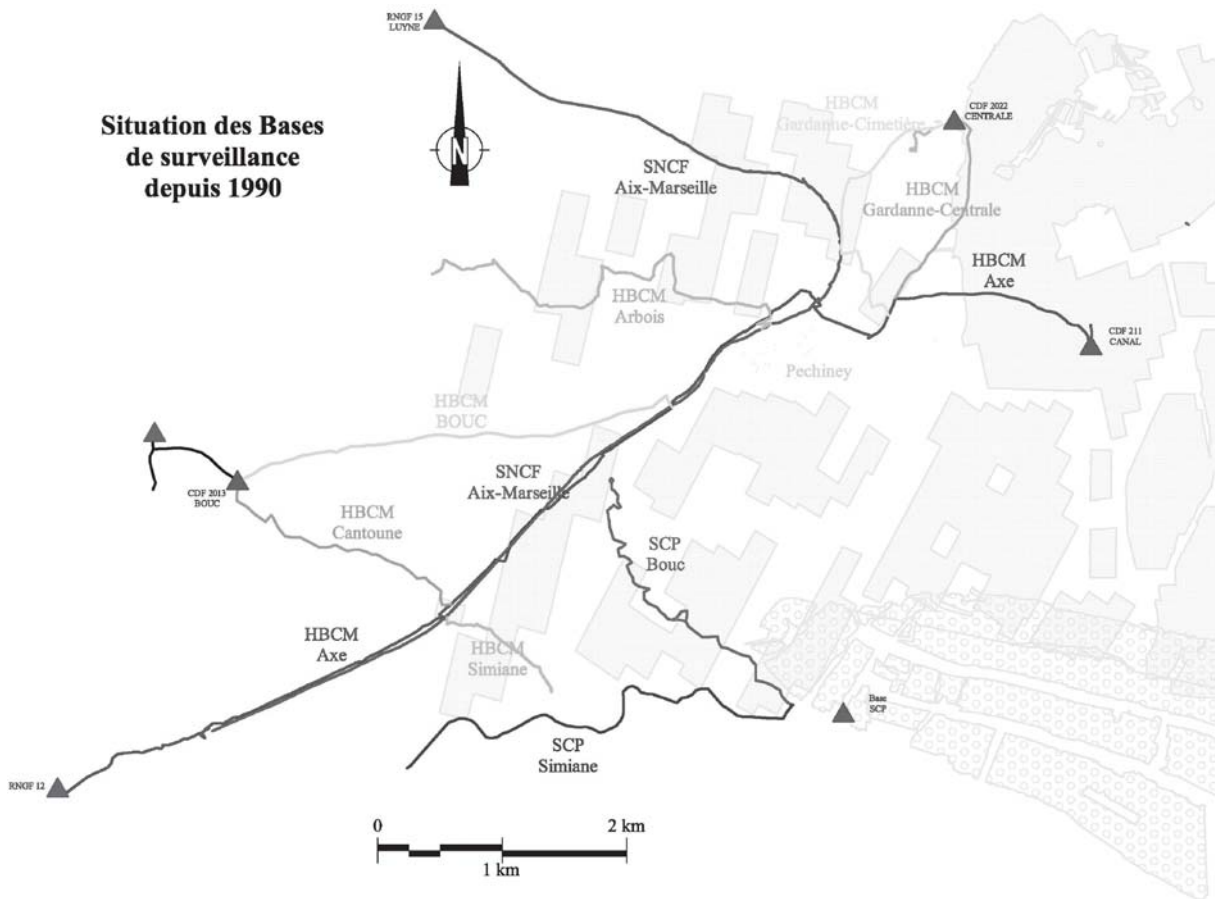
588  
589 VAN DER KOOIJ, M., HUGHES, W., SATO, S. and V. PONCOS, 2005, Coherent Target  
590 Monitoring at High Spatial Density, Examples of Validation Results, Fringe 2005, ESA, Frascati,  
591 Italy, November 28 – December 2, 2005  
592  
593 WERNER, C., WEGMULLER, U., STROZZI, T., AND WIESMANN, A. 2003, "Interferometric  
594 point target analysis for deformation mapping", Proceedings of IGARSS'03, vol. 7, pp. 4362-4364,  
595  
596



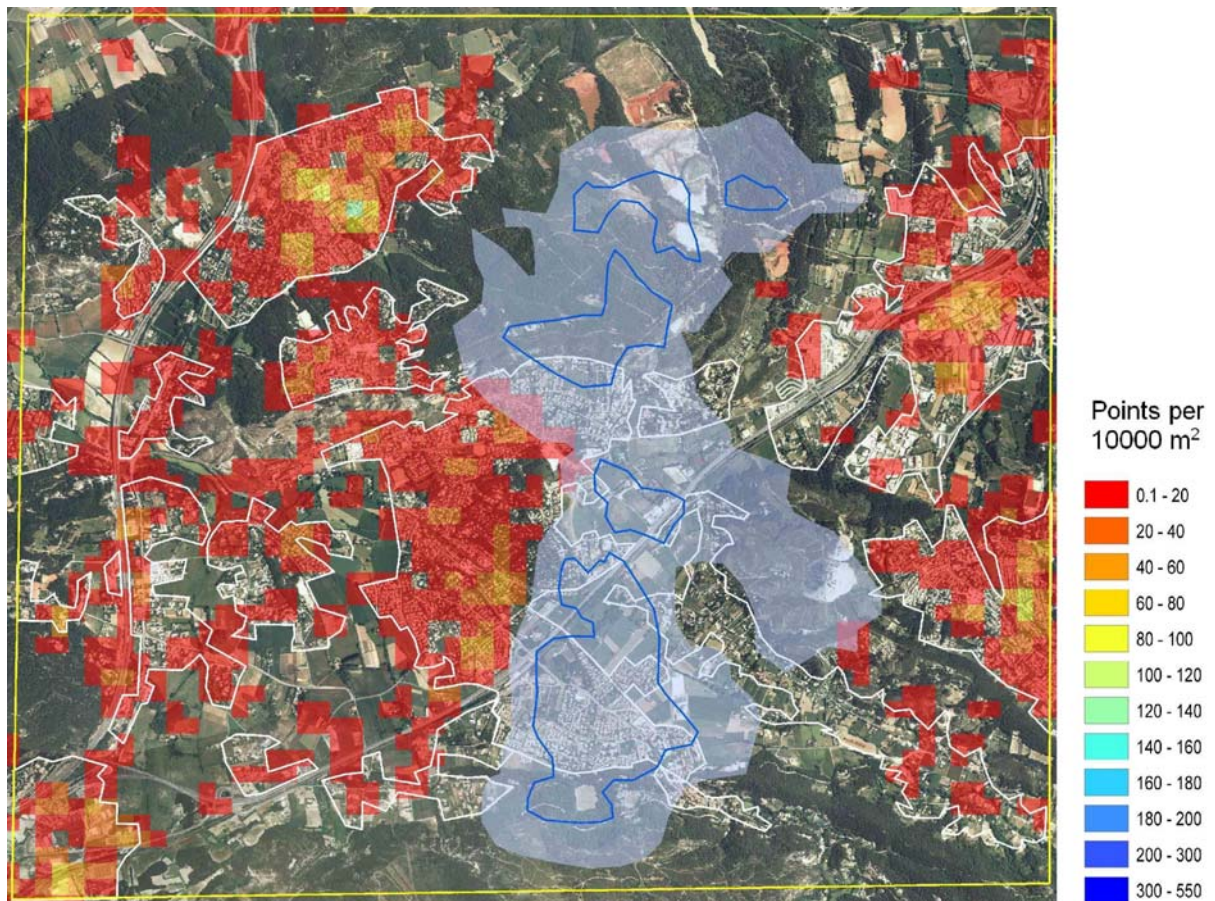
597



598



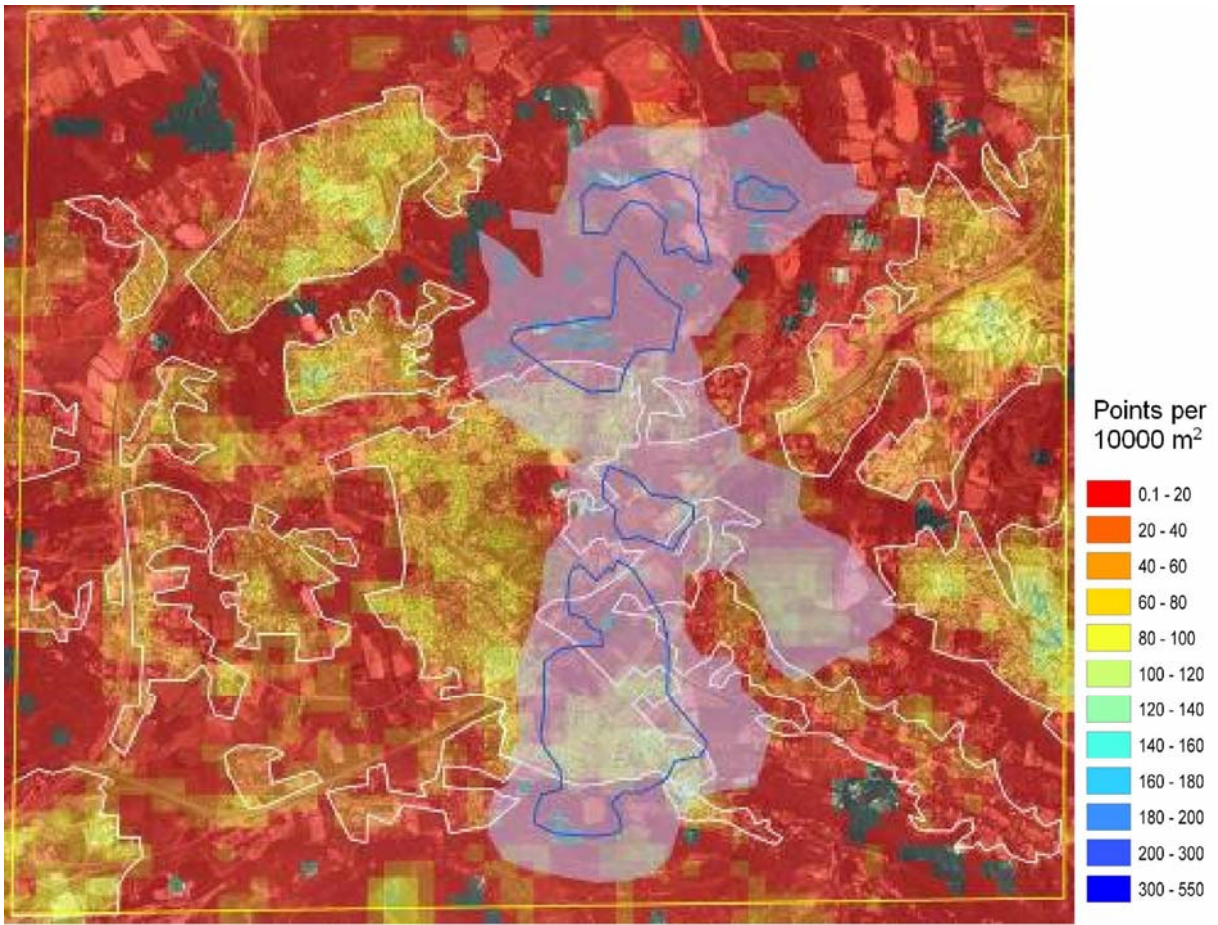
599



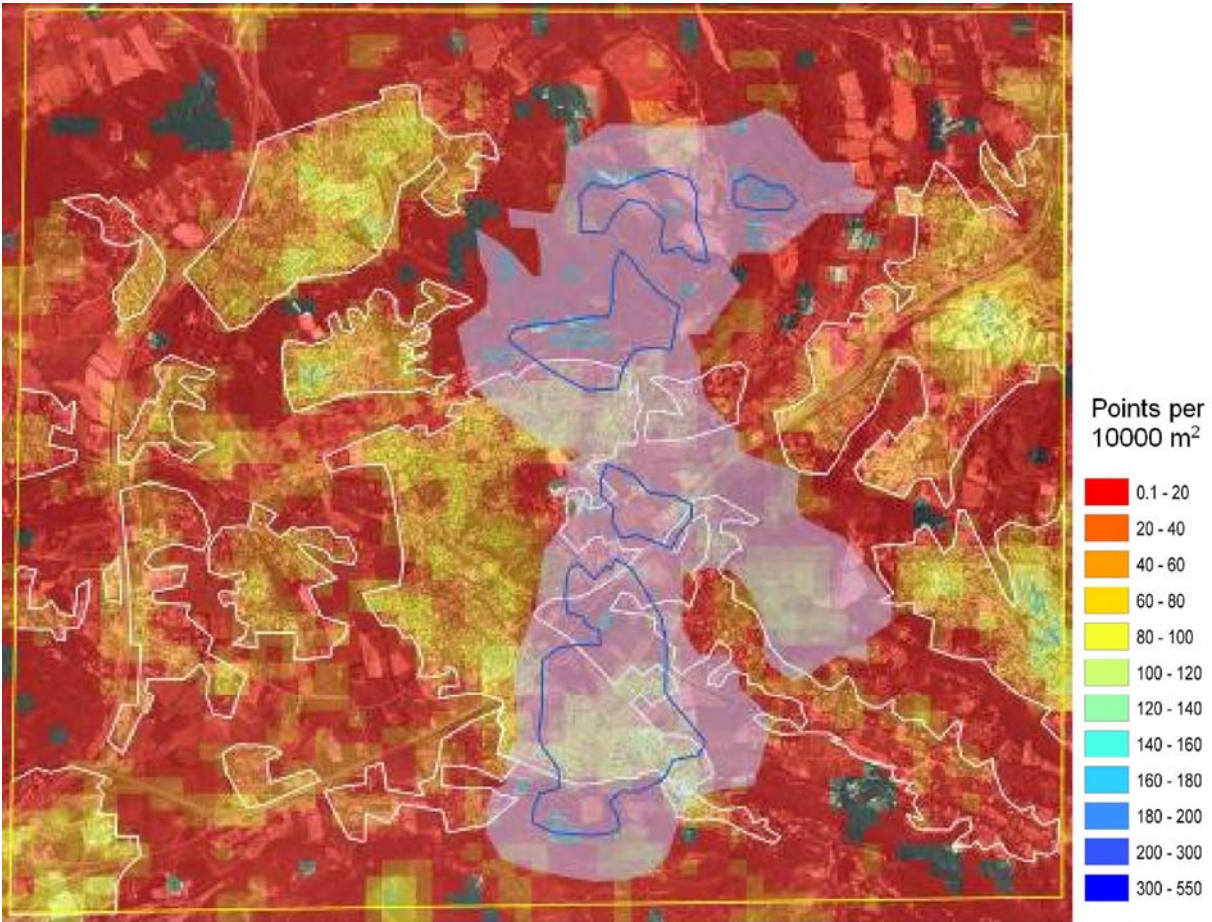
600

601

602



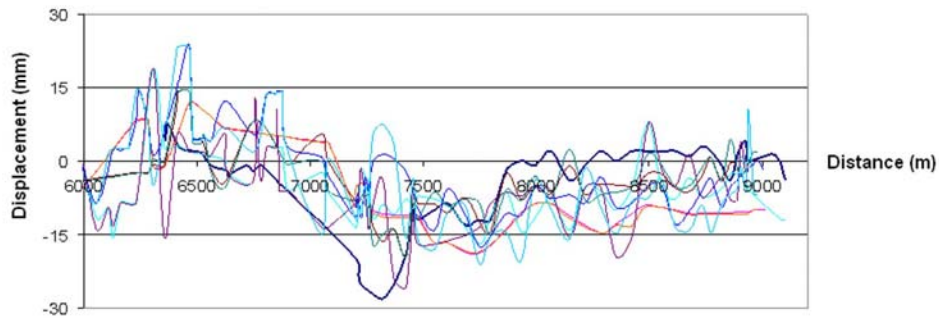
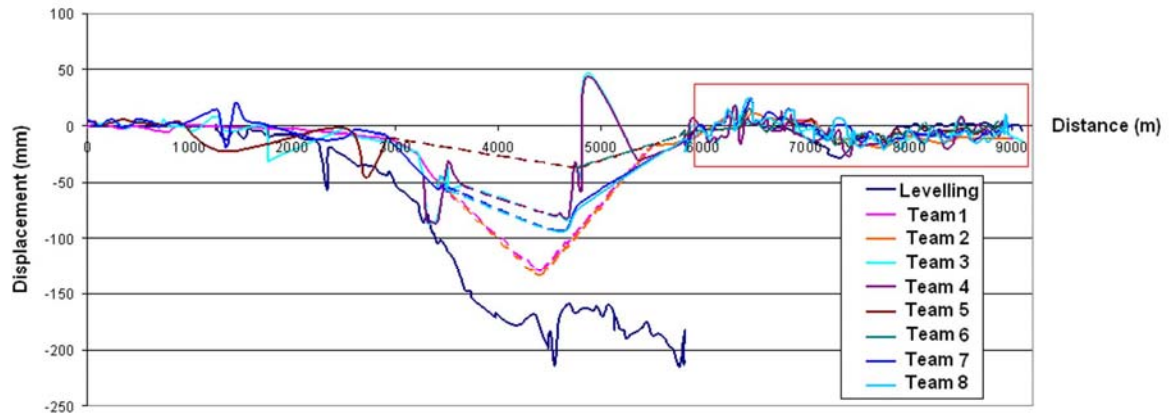
603



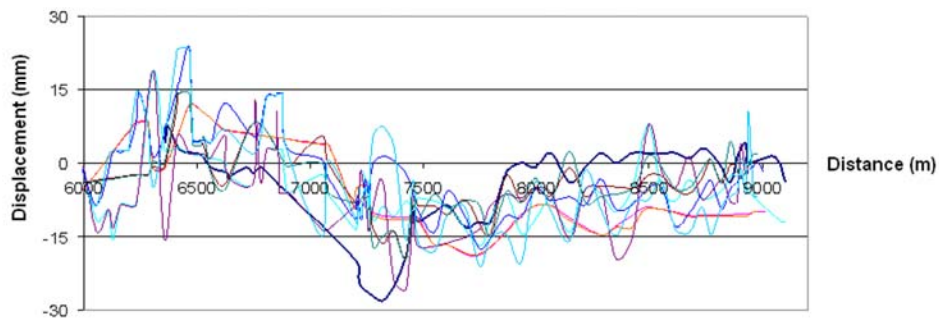
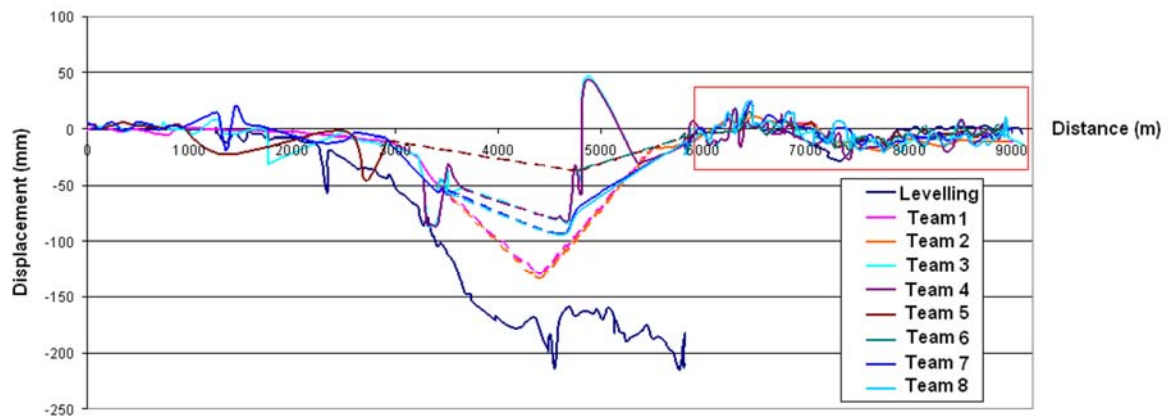
604

605

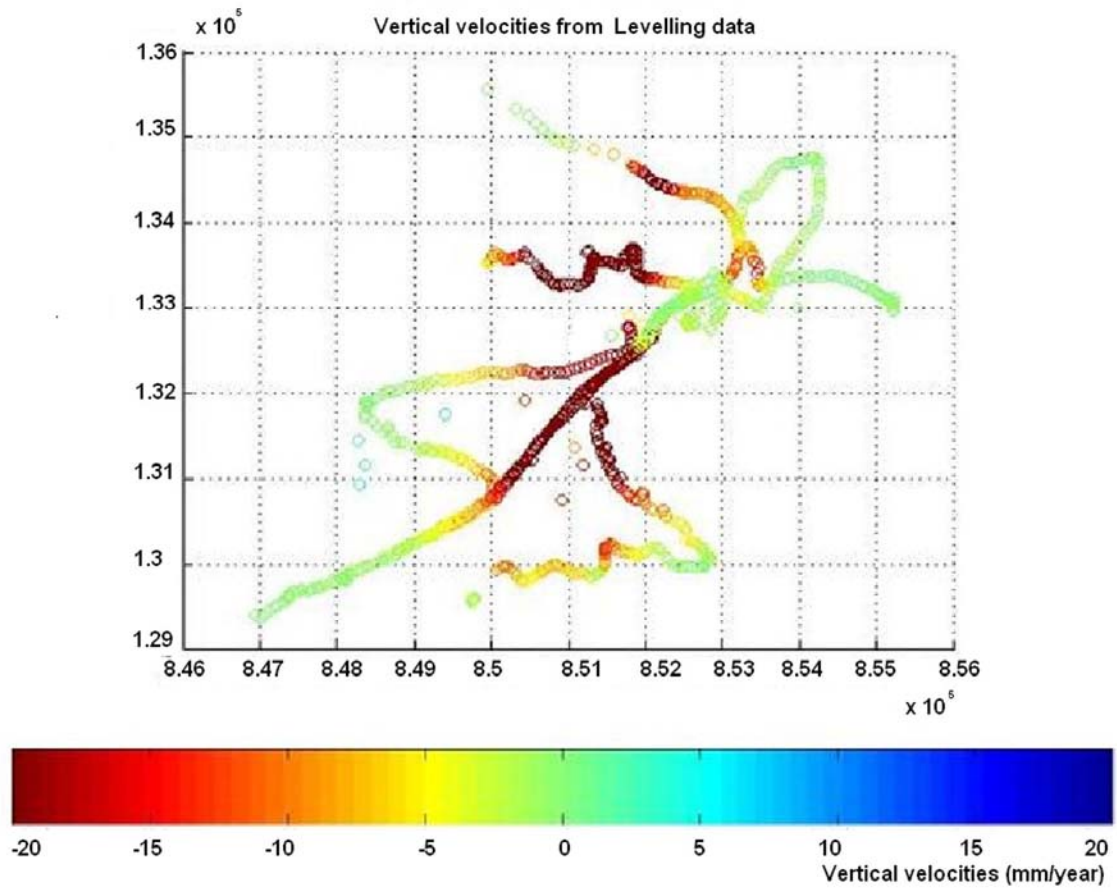




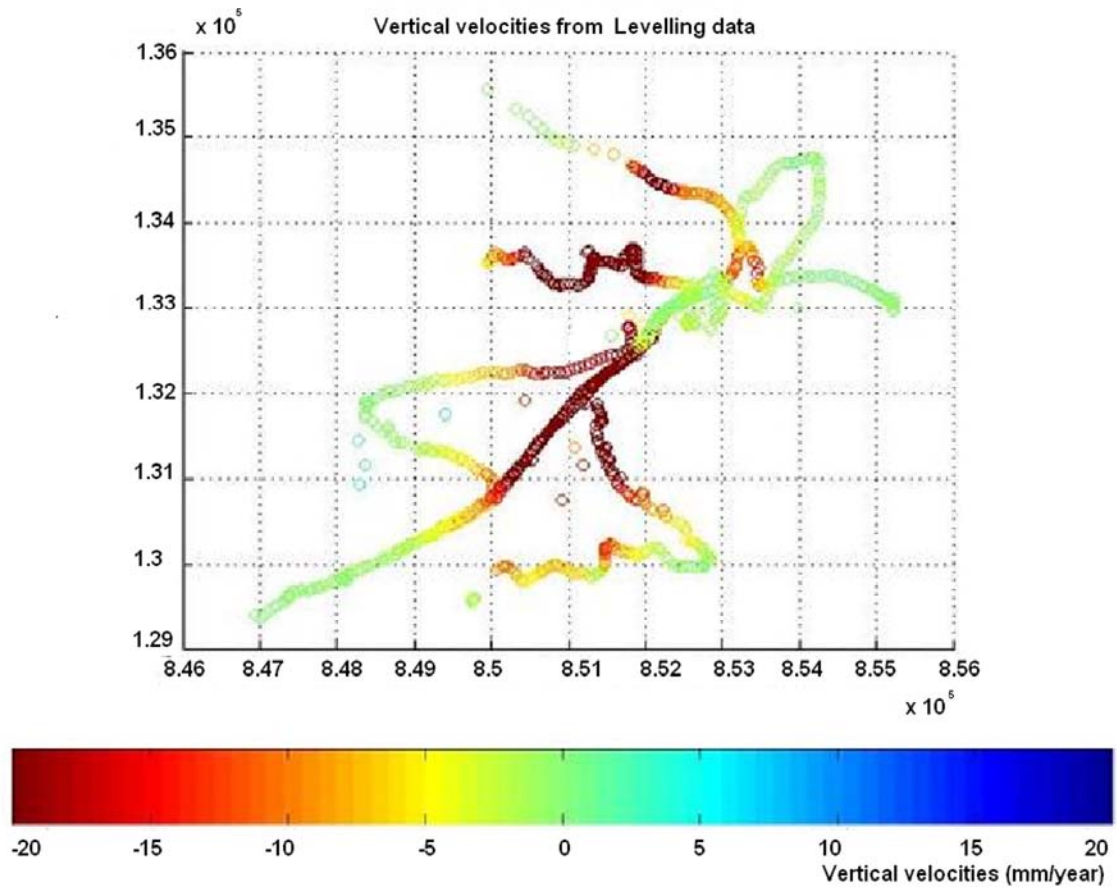
606



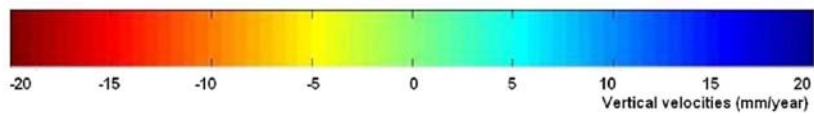
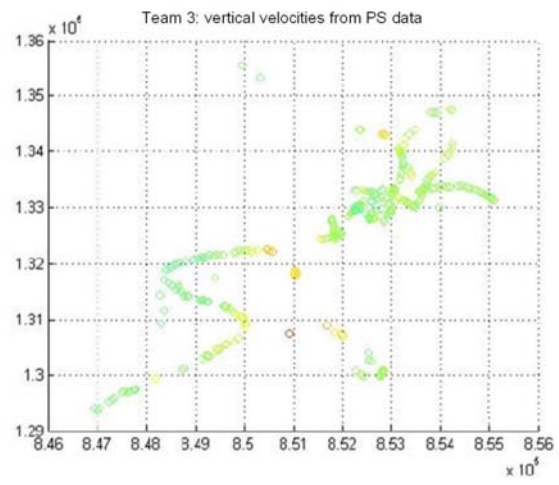
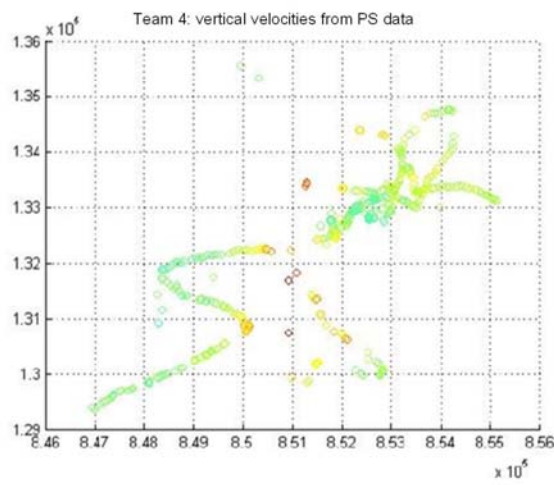
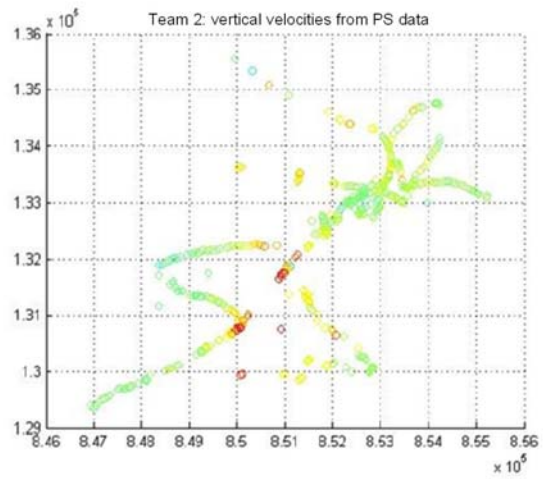
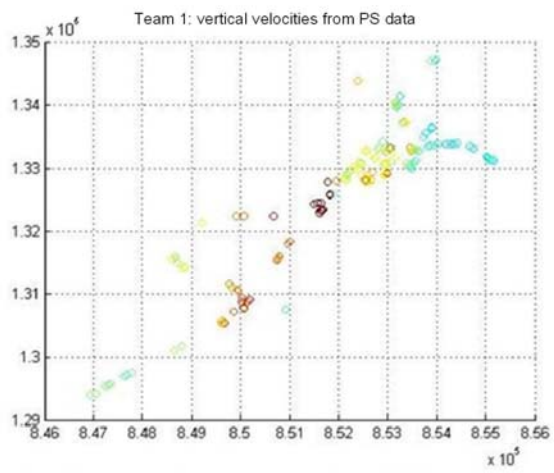
607

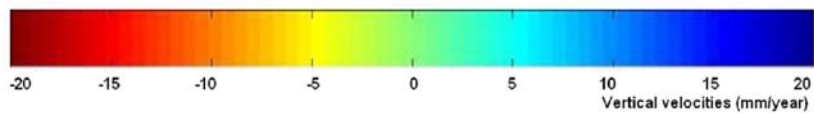
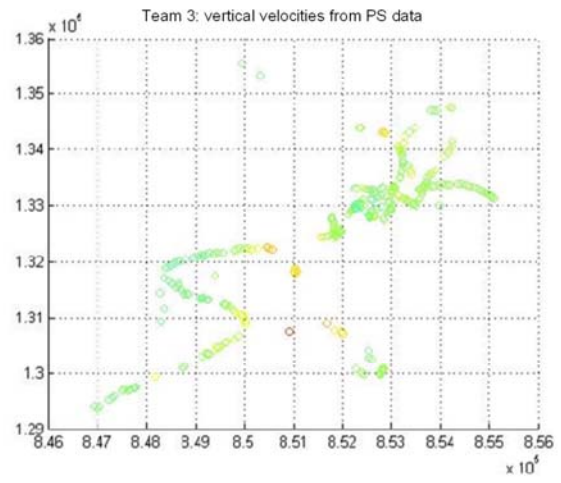
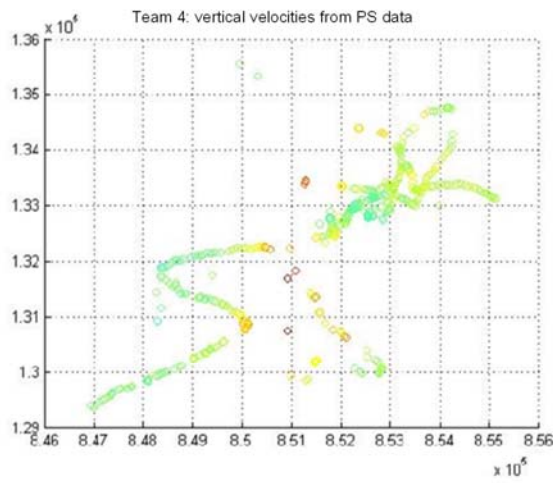
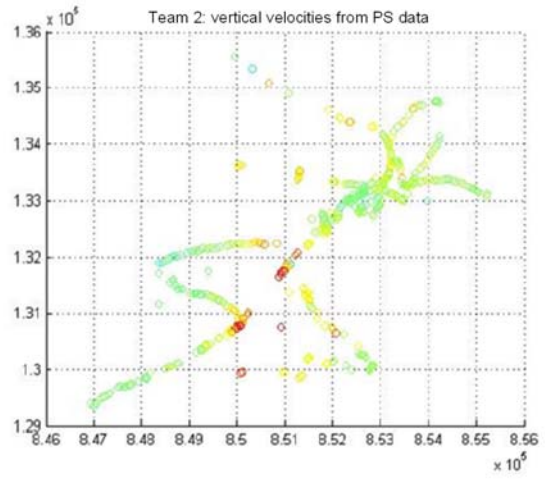
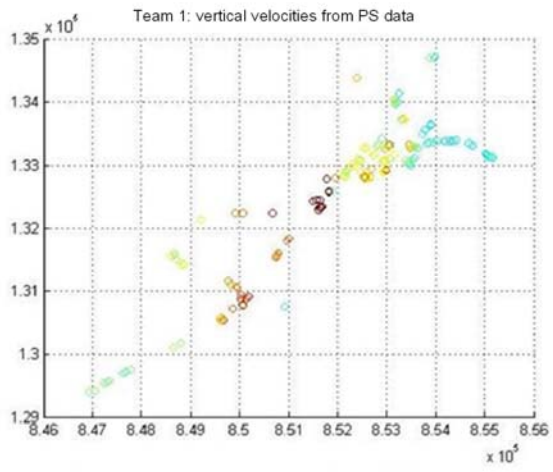


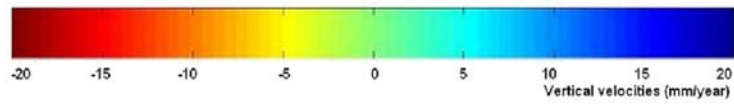
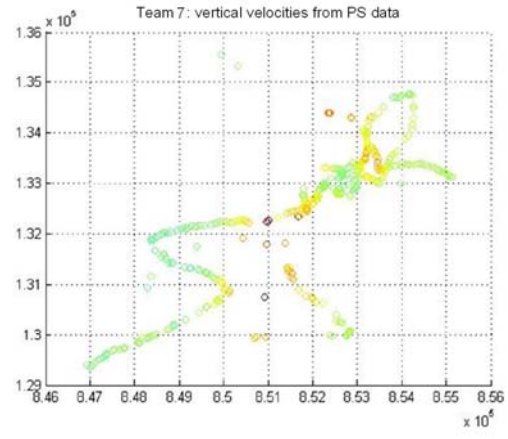
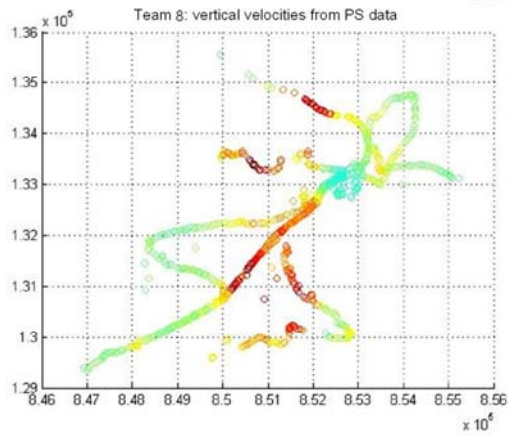
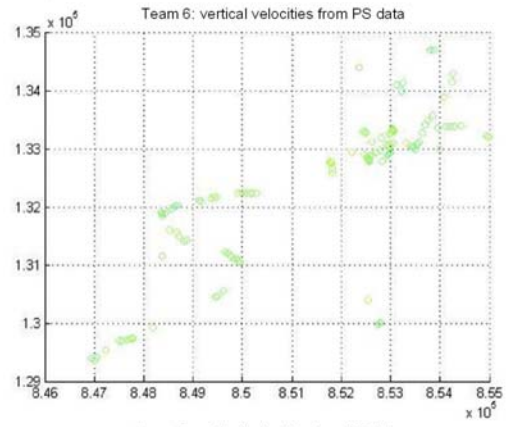
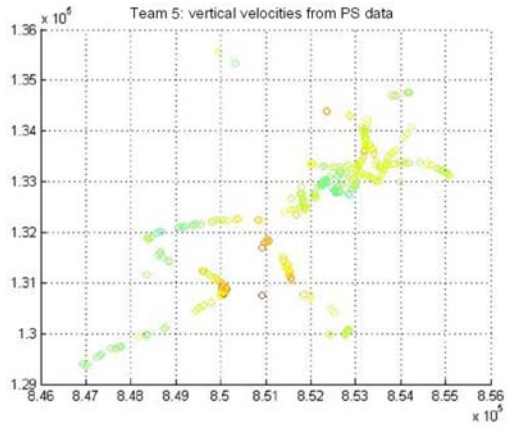
608



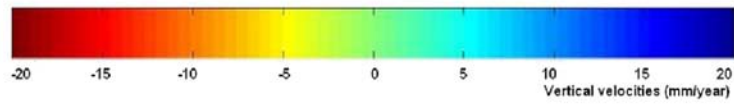
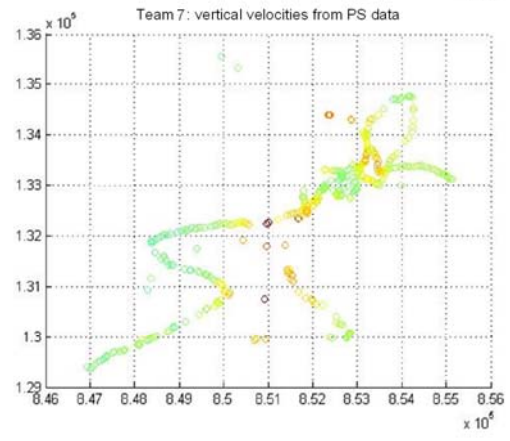
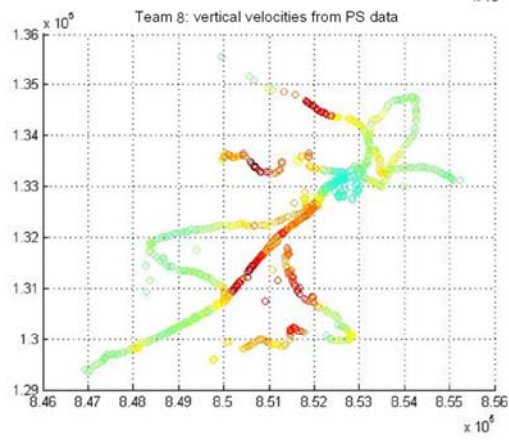
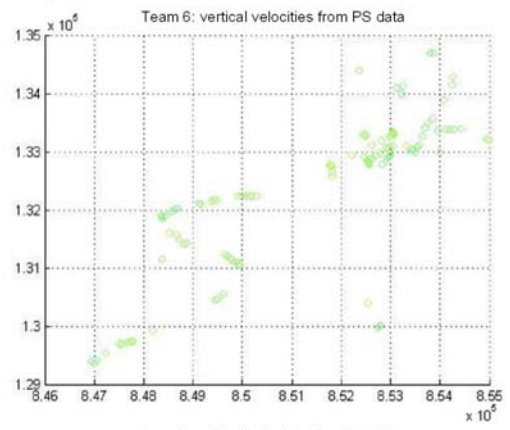
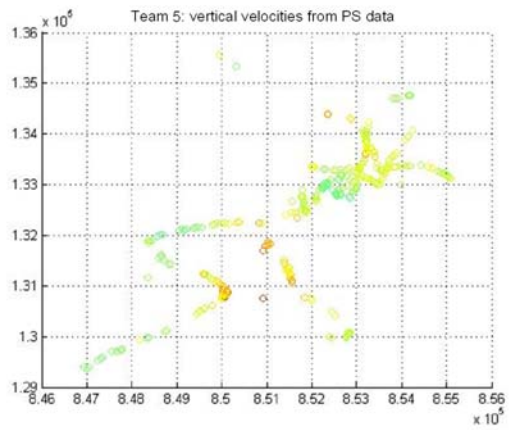
609







612



613